# A nodal-integration based particle finite element method (N-PFEM) to model cliff recession

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## 1 Abstract

2 Cliff recession poses a significant threat to the built environment, transportation infrastructure and 3 land use. In this paper, a novel computational framework called the Nodal-integration based Particle 4 Finite Element Method (N-PFEM) is developed for modelling the cliff recession resulting from 5 weathering-induced landslides. The N-PFEM combines the nodal-integration technique with the 6 PFEM in second-order cone programming and thus requires no variable mapping operation, which is 7 essential in the classical PFEM for modelling history-dependent materials, for modelling large 8 deformation problems such as landslides in cliff recession processes. To verify the developed N-9 PFEM, a series of benchmarks have been simulated including the cliff recession under both the 10 weathering-limited and transport-limited conditions. Simulation results from the N-PFEM are 11 validated in detail to these from the limit analysis method, well established geomorphologic models 12 and the discrete element method. Additionally, measures for preventing cliff recession such as the 13 construction of retaining wall structures are also investigated using the N-PFEM.

14 Keywords: Cliff recession, Landslide, PFEM, Nodal integration

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#### 19 **1. Introduction**

20 Modelling the progressive retreat of cliffs has recently received considerable attention by the 21 engineering community due to increasing coastal erosive processes caused by climate change and amplified environmental awareness at national and European level (Bray and Hooke, 1997). Also 22 23 the insurance industry needs reliable models for the predictions of the amount of cliff retreat over 24 time for residential buildings located in exposed areas whereas local authorities and decision makers 25 need to know the level of risk faced by residential buildings and public infrastructure (e.g. coastal 26 roads, pedestrian footpaths, car parks, etc.). Along many coastal areas, the recession rate of cliffs is 27 significantly high leading to non-negligible socio-economic impacts (Bird, 2016). A typical example is the soft glacial drift cliff at Holderness coast, Aldbrough, UK, that erodes at about 2 meter per 28 29 year with a maximum recession rate being 3.4 meter per year (Hobbs et al., 2020). This rapid 30 erosion results in critical impacts on human communities including, but are not limited to, the loss of 31 farmland, damage to infrastructure, properties and tourism. In order to engineer adequate mitigation 32 and remediation measures, it is thus of vital importance to develop reliable models to predict cliff recession rates driven by the local environmental conditions. 33

34 The natural drivers responsible for triggering cliff recession can be classified into two categories: factors increasing the driving forces, such as seismic loading (Massey et al., 2017), wave 35 36 loading (Sunamura, 1982) and anthropogenic activities (Xue et al., 2009), and factors decreasing 37 resistance forces, such as weathering (Utili and Crosta, 2011a) and crack formation (Kogure et al., 2006, Voulgari and Utili, 2017). This paper focuses on the latter and on weathering-induced cliff 38 39 recession in particular. Indeed, weathering proceeds in the manner of physical break down and 40 chemical alteration of rocks which weakens the shear strength of rocks and form thick sequences of 41 weathered geo-materials whose engineering properties have been highly altered. The weathering 42 process gradually decreases the stability of the slope, which ultimately leads to subsequent

landslides and the progressive retreat of the cliff crest (Hutchinson, 2001). The earliest effort 43 devoted to forecasting the progressive evolution of cliff erosion lies in (Fisher, 1866) in which a 44 model was proposed to study the disintegration of a coastal chalk cliff considering the accumulation 45 46 of a basal debris apron. This model was then extended by further introducing the effect of different weathering stages (Lehmann, 1933) and the upper and lower sloping sectors (Bakker and Le Heux, 47 1946. Nash, 1981). In these models, the evolution process is controlled based on assumptions 48 49 concerning the geometry of falling blocks, the bulking of geo-materials, the accumulation of debris 50 and the weathering rate. Further consideration of the effects of mechanical properties of geo-51 materials on the evolution process were performed in Andrews and Bucknam (1987). More recently, 52 an analytical method based on an upper bound limit analysis approach was proposed to simulate the cliff recession process (Utili and Crosta, 2011a). The basic idea is to perform limit analysis of slope 53 instability induced by a homogeneous decrease of ground strength with the debris following the 54 landslide occurrence being removed before the onset of the successive landslide. This implies 55 weathering-limited slope conditions. Voulgari and Utili (2017) extended the limit analysis model of 56 57 (Utili and Crosta, 2011a) to account for the effects of seismic actions, the formation of tension cracks and seepage. From their modelling it emerges that although the formation of tension cracks 58 59 affects the geometry of each landslide profile, it bears little influence on the overall geomorphologic 60 evolution of the cliff especially relative to rock strength degradation. Therefore, the onset of tension 61 crack was not considered in the paper.

Using numerical techniques, cliff recession under more complex conditions, e.g. non uniform slope weathering, have been investigated, such as the evolution of natural cliffs subject to weathering using the discrete element method (Utili and Crosta, 2011a), the evolution of an overhanging rock slope using the displacement discontinuity method (Zhang et al., 2016). These studies, except for Utili and Crosta (2011a, 2011b), focus on the factors triggering slope instabilities leading to cliff 67 recession without attempting to predict the complete evolution of cliff recession that results from a68 discrete sequence of landslides.

69 In order to predict the complete evolution process of cliff recession and the resulting 70 geomorphology, the numerical approach adopted has to be capable of predicting not only the failure of a slope but also the post-failure process such as the mass transport and deposition which affects 71 72 the subsequent slope failures. Recently, several methods have been developed and applied to analyse 73 landslide and granular flow problems such as the smoothed particle hydrodynamics (SPH) method (Bui et al., 2011, Pastor et al., 2014), the material point method (MPM) (Andersen and Andersen, 74 2010, Soga et al., 2016, Tran and Sołowski, 2019), the finite volume method (FVM) (Mangeney et 75 al., 2003, 2007a), the discrete element method (DEM) (Staron and Hinch, 2005, Staron, 2008), the 76 77 dedicated numerical model, SHALTOP, that can consider complex 3D topography (Lucas and 78 Mangeney, 2007, Mangeney et al., 2007b, Favreau et al., 2010), the particle finite element method (PFEM) (Zhang et al., 2014, Wang et al., 2019, Zhang et al., 2019b, Mulligan et al., 2020, Yuan et 79 al., 2020), etc. The PFEM is developed based on the idea that mesh nodes are treated as particles 80 81 that can move freely and even separate from the domain they originally belong to (Idelsohn et al., 2004, Oñate et al., 2004). Computational domains are re-identified based on the particles followed 82 83 by mesh generations. It has been shown that the PFEM is particularly suitable for modelling large 84 deformation problems with free-surface evolutions (Oñate et al., 2011). So far, a series of challenging problems, in addition to landslides, in geomechanical and geotechnical problems have 85 86 been tackled using the PFEM including, but are not limited to, ground excavation (Carbonell et al., 87 2010), granular flows (Lagrée et al., 2011, Zhang et al., 2013, Cante et al., 2014, Staron et al., 2014, Franci and Cremonesi, 2019), soil-structure interactions (Oñate et al., 2011, Monforte et al., 2017). 88 89 Nevertheless, it is worth to note that a drawback associated with the PFEM for solving problems 90 with history-dependent materials in geomechanics, or more generally solid mechanics, is the

91 requirement of the variable mapping operation (Zhang et al., 2013). When new meshes are generated in the PFEM procedure, although information of displacements, velocities and 92 accelerations is stored at mesh nodes and requires no mapping operation, field variables such as 93 94 stresses, strains and other internal variables, for example those controlling the strain softening for sensitive clays (Zhang et al., 2017), have to be mapped from the old to the new integration points. 95 Such a mapping operation inevitably leads to errors and increases complexity. To overcome this 96 97 issue, the nodal integration method was recently introduced to the PFEM for geotechnical problems 98 that the integrals are performed over smoothed domains rather than finite elements and all field 99 variables are computed and stores at nodes (Zhang et al., 2018, Yuan et al., 2019). Franci et al. 100 (2020) then explored the possibility of a nodal PFEM formulation for free-surface fluid dynamics problems. It has been shown that the PFEM with the nodal-integration technique not only 101 102 circumvents the requirement of variable mappings as in the classical PFEM but also allows the use 103 of low order elements, such as the 3-node triangle element, without volumetric locking and 104 insensitivity to mesh distortion.

105 In this paper, the nodal-integration technique is implemented in the version of the PFEM developed in Zhang et al. (2013, 2017). Specifically, the nodal-integration technique is introduced to 106 107 the dynamic finite element formulation in mathematical programming to form the N-FEM which is 108 then incorporated in the PFEM framework to form the Nodal-based PFEM (N-PFEM). Compared to 109 the existing PFEM with nodal-integration techniques in geomechanics in which an explicit 110 integration scheme is used, the developed N-PFEM is based on implicit integration so that 111 significantly larger time steps can be used in the simulation. This feature is of great importance for 112 the modelling of cliff recession subject to weathering since the length of time to model is of some 113 years. Additionally, as the finite element formulation developed in the presented N-PFEM is based on second-order cone programming, it inherits some unique merits of the FEM in mathematical 114

115 programming that have been showcased in Krabbenhøft et al. (2007). Among them, a notable 116 advantage for modelling landslide-induced cliff erosion rests with its convergence properties. Indeed, 117 significant changes of stresses or strains may occur in just a very small time interval in the post-118 failure analysis of landslides. Whereas this critical change is likely to result in the nonconvergence of the standard Newton-Raphson based FEM, the adopted advanced optimisation algorithms still 119 120 guaranteed the convergence regardless of whether the solved known states (e.g. the field variables at  $t_n$ ) are close to the new unknown states (e.g. the field variables at  $t_{n+1}$ ) (Zhang et al., 2014). The 121 nodal integration also enables a more straightforward treatment of the cohesive-frictional contacts 122 123 compared to the PFEM developed in Zhang et al. (2013, 2017). Last but not least, low-order 124 elements such as the three node triangular element can be used in the developed version for tackling nearly incompressible problems while maintaining satisfactory accuracy which is not possible for 125 the earlier PFEM version in Zhang et al. (2013, 2017) where six-node triangular elements are 126 127 adopted for modelling geomechanical problems. To demonstrate the robustness and correctness of the developed N-PFEM, cliff recession under both weathering-limited and transport-limited 128 129 conditions are considered. Comparisons between the simulation results from the N-PFEM and these from analytical approach and DEM modelling available in the literature are performed. Furthermore, 130 131 the use of retaining wall for alleviating cliff recession subject to weathering is investigated using the 132 N-PFEM with focus on the recession distance and the evolution of the resistance force from the 133 retaining wall.

The paper is organised as follows. In Section 2, we present the formulations of the nodalintegration based finite element method (N-FEM). The merging of the N-FEM into the PFEM framework to form the N-PFEM is detailed in Section 3. Numerical benchmarks such as a block sliding on a rigid surface, cliff recession under uniform and non-uniform weathering conditions are 138 illustrated in Sections 4 to 6 and the design of retaining wall structure to prevent cliff recession is

139 discussed in Section 7. Conclusions are drawn in Section 8.

140

# 141 **2. Nodal-integration based finite element method (N-FEM)**

Before introducing the nodal-integration based particle finite element method (N-PFEM), the nodalintegration based finite element method (N-FEM) is formulated in this section, which is the core for solving the equations governing cliff erosion processes.

145 2.1 Governing equations

# 146 Considering a continuum medium with volume $\Omega$ , the momentum equations read

147 
$$\begin{cases} \nabla^{\mathrm{T}} \boldsymbol{\sigma} + \boldsymbol{b} = \rho \boldsymbol{v}, \text{ in } \Omega \\ \mathbf{N}^{\mathrm{T}} \boldsymbol{\sigma} = \boldsymbol{t}, \text{ on } \Gamma \end{cases}$$
(1)

where  $\nabla$  is the usual linear strain-displacement differential operator,  $\sigma$  is the stresses, **b** is the body forces,  $\rho$  is the density of the medium, v is the velocity with a superposed dot representing differentiation with respect to time, **N** is the matrix containing the unit outward normal to the boundary  $\Gamma$  and **t** is the traction. It can be discretised in time using the standard  $\theta$ -method:

152 
$$\begin{cases} \nabla^{\mathrm{T}} \Big[ \theta_{1} \boldsymbol{\sigma}_{\mathrm{n+1}} + (1-\theta_{1}) \boldsymbol{\sigma}_{\mathrm{n}} \Big] + \boldsymbol{b} = \rho \frac{\boldsymbol{v}_{\mathrm{n+1}} - \boldsymbol{v}_{\mathrm{n}}}{\Delta t} \text{ and } \theta_{2} \boldsymbol{v}_{\mathrm{n+1}} + (1-\theta_{2}) \boldsymbol{v}_{\mathrm{n}} = \frac{\Delta \boldsymbol{u}}{\Delta t} \\ \mathbf{N}^{\mathrm{T}} \Big[ \theta_{1} \boldsymbol{\sigma}_{\mathrm{n+1}} + (1-\theta_{1}) \boldsymbol{\sigma}_{\mathrm{n}} \Big] = \boldsymbol{t}_{\mathrm{n}}, \text{ on } \Gamma \end{cases}$$
(2)

where the subscripts n and n+1 refer to the known and unknown states in a typical time step,  $\theta_1$  and  $\theta_2$  are parameters taking values in [0, 1] and  $\Delta \mathbf{u} = \mathbf{u}_{n+1} - \mathbf{u}_n$ . The above equation can be reformulated as:

156 
$$\begin{cases} \nabla^{\mathrm{T}} \boldsymbol{\sigma}_{\mathrm{n+1}} + \frac{(1-\theta_{\mathrm{l}})}{\theta_{\mathrm{l}}} \nabla^{\mathrm{T}} \boldsymbol{\sigma}_{\mathrm{n}} + \hat{\mathbf{b}}_{\mathrm{n}} = \mathbf{r}_{\mathrm{n+1}} \\ \theta_{\mathrm{l}} \mathbf{N}^{\mathrm{T}} \boldsymbol{\sigma}_{\mathrm{n+1}} = \hat{\mathbf{t}}_{\mathrm{n}} \end{cases}$$
(3)

157 where

158 
$$\begin{cases} \hat{\mathbf{b}}_{n} = \frac{\mathbf{b}}{\theta_{1}} + \frac{\rho \mathbf{v}_{n}}{\theta_{1} \Delta t \theta_{2}} \\ \mathbf{r}_{n+1} = \frac{\rho}{\theta_{1} \theta_{2}} \frac{\Delta \mathbf{u}}{\Delta t^{2}} \\ \hat{\mathbf{t}}_{n} = -\mathbf{N}^{\mathrm{T}} (1 - \theta_{1}) \boldsymbol{\sigma}_{n} + \mathbf{t}_{n} \end{cases}$$
(4)

Following Zhang et al. (2019a), the generalised Hellinger–Reissner variational principle for the discretised governing equations (3) and (4) with rate-independent materials reads:

$$\min_{\Delta \mathbf{u}} \max_{\boldsymbol{\sigma}_{n+1}, \boldsymbol{r}_{n+1}} \int_{\Omega} \boldsymbol{\sigma}_{n+1}^{\mathrm{T}} \nabla (\Delta \mathbf{u}) d\Omega + \frac{1 - \theta_{1}}{\theta_{1}} \int_{\Omega} \boldsymbol{\sigma}_{n}^{\mathrm{T}} \nabla \Delta \mathbf{u}_{n+1} d\Omega - \int_{\Omega} \hat{\boldsymbol{b}}_{n}^{\mathrm{T}} \Delta \mathbf{u}_{n+1} d\Omega - \int_{\Gamma} \hat{\boldsymbol{t}}_{n}^{\mathrm{T}} \Delta \mathbf{u}_{n+1} d\Gamma$$

$$- \int_{\Omega} \frac{1}{2} (\Delta \boldsymbol{\sigma}_{n+1})^{\mathrm{T}} \mathbb{C} (\Delta \boldsymbol{\sigma}_{n+1}) d\Omega - \frac{1}{2} \Delta t^{2} \int_{\Omega} \boldsymbol{r}_{n+1}^{\mathrm{T}} \hat{\rho}^{-1} \boldsymbol{r}_{n+1} d\Omega + \int_{\Omega} \boldsymbol{r}_{n+1}^{\mathrm{T}} \Delta \mathbf{u}_{n+1} d\Omega$$

$$\text{subject to } F(\boldsymbol{\sigma}_{n+1}) \leq 0$$
(5)

In the above mixed variational principle, displacements, stresses and inertial forces are independent fields. *F* is the yield function and  $\hat{\rho} = \frac{\rho}{\theta_1 \theta_2}$ .  $\mathbb{C}$  is the elastic compliance modulus that, for plane-

164 strain cases, is in the form:

165 
$$\mathbb{C} = \frac{1+\nu}{E} \begin{bmatrix} 1-\nu & -\nu & 0\\ -\nu & 1-\nu & 0\\ 0 & 0 & 2 \end{bmatrix}$$
(6)

166 where *E* and v are elastic modulus and Poisson's ratio, respectively. Its validity has been proved in 167 Zhang et al. (2013, 2019a), and interested reader are referred to these references for details.

#### 168 2.2 FEM with the nodal-integration technique

The discrete form of program (5) needs to be derived by using spatial discretisation methods. In the paper, the nodal-integration based finite element technique is adopted. By first using the standard finite element discretisation, we have:

172 
$$\mathbf{u} \approx \mathbf{N}_{\mathrm{u}} \hat{\mathbf{u}}$$
 (7)

173 where  $\hat{\mathbf{u}}$  is a vector consisting of displacement at mesh nodes and  $\mathbf{N}_{u}$  is a matrix of shape functions. 174 Following the classic FEM, the strain field is then approximated as:

175 
$$\boldsymbol{\varepsilon} = \mathbf{B}_{\mathrm{u}} \hat{\mathbf{u}}$$
 (8)

where  $\mathbf{B}_{u} = \nabla \mathbf{N}_{u}$  is the strain-displacement matrix. Contrary to the standard FEM that integration is 176 177 performed over finite elements, the integration in the proposed method is operated on cells. As 178 shown in Fig.1, a cell (also called a smoothing domain) is a "non-overlap" and "no-gap" domain associated with each mesh node, for example, the cell  $\Omega_k^s$  associated with the kth node (the red zone 179 180 in Fig. 1) that covers several one-third of adjacent elements of a node k. The coloured polygon is 181 bounded by multiple straight boundary segments which connect the midpoint of an edge to a 182 centroid of the triangular elements. Consequently, the strain in the kth cell is the weighted average of 183 the strain (or called smoothed strained) of all the one-third adjacent elements of node k (Zhang et al., 184 2018, Yuan et al., 2019, Franci et al., 2020)

185 
$$\overline{\boldsymbol{\varepsilon}}_{k} = \int_{\Omega_{k}^{s}} \boldsymbol{\Phi}_{k}(\mathbf{x}) \boldsymbol{\varepsilon}(\mathbf{x}) \mathrm{d}\boldsymbol{\Omega} = \int_{\Omega_{k}^{s}} \boldsymbol{\Phi}_{k}(\mathbf{x}) \mathbf{B}_{u} \hat{\mathbf{u}} \mathrm{d}\boldsymbol{\Omega}$$
(9)

186 where  $\Phi_k(\mathbf{x})$  is the smoothing function in the form (Liu et al., 2009, Liu and Trung, 2010):

187 
$$\Phi_k(\mathbf{x}) = \begin{cases} 1/A_k^s, \, \mathbf{x} \in \Omega_k^s \\ 0, \quad \mathbf{x} \notin \Omega_k^s \end{cases}$$

188 in which  $A_k^s$  is the area of the smoothing domain  $\Omega_k^s$ .

189





191 Fig. 1. Node-based cells (also called smoothing domains) based on given triangle mesh.

192

As demonstrated in Fig.1, the cell  $\Omega_k^s$  is comprised of  $N_s$  sub-domains that are one-third of the triangular elements adjacent to node k. For the linear triangular element, the strain is constant inside the elements. Therefore, substituting Eq. (10) into (9), the smoothed strain  $\overline{\epsilon}_k$  is simply:

196 
$$\overline{\boldsymbol{\varepsilon}}_{k} = \frac{1}{\mathbf{A}_{k}^{s}} \sum_{i=1}^{N_{s}} \frac{1}{3} \mathbf{A}_{i}^{e} \boldsymbol{\varepsilon}_{i}^{e} = \frac{1}{\mathbf{A}_{k}^{s}} \sum_{i=1}^{N_{s}} \frac{1}{3} \mathbf{A}_{i}^{e} \mathbf{B}_{i}^{e} \hat{\mathbf{u}}_{i}^{e}$$
(11)

(10)

197 where *i* is the element number and  $A_i^e$ ,  $\boldsymbol{\varepsilon}_i^e$ ,  $\mathbf{B}_i^e$  and  $\hat{\mathbf{u}}_i^e$  are the area, the strain, the strain gradient 198 matrix and the displacement of the *i*th triangular element, respectively. For simplicity, the smoothed 199 strain on the cell  $\Omega_k^s$  is written as:

200 
$$\overline{\boldsymbol{\varepsilon}}_{k} = \overline{\boldsymbol{B}}_{k} \hat{\boldsymbol{u}}_{k}$$
 with  $\overline{\boldsymbol{B}}_{k} = \frac{1}{A_{k}^{s}} \sum_{i=1}^{N_{s}} \frac{1}{3} A_{i}^{e} \boldsymbol{B}_{i}^{e}$  (12)

201 where  $\hat{\mathbf{u}}_k$  is the nodes' displacement of the cell k. Thus for each cell, the strain increment

$$202 \qquad \Delta \boldsymbol{\varepsilon} = \nabla (\Delta \mathbf{u}) = \mathbf{B} \hat{\mathbf{u}} \tag{13}$$

203 where  $\overline{\mathbf{B}}$  consists of the smoothed strain-displacement matrix as shown in Eq (12).

As the mixed variational principle is adopted, state variables such as the displacement, the stress and the dynamic force have to be interpolated independently. Herein for each cell, we have

$$206 \quad \boldsymbol{\sigma} \approx \mathbf{N}_{\sigma} \boldsymbol{\bar{\sigma}} \tag{14}$$

where  $\bar{\sigma}$  is the smoothed stress and  $N_{\sigma}$  is the shape function matrix which are in fact identity matrices since the linear approximation is made for the displacement field, and

$$209 \quad \boldsymbol{r} \approx \mathbf{N}_{\mathrm{r}} \tilde{\boldsymbol{r}} \tag{15}$$

where  $\tilde{r}$  is the dynamic force applied on the nodes, and  $N_r$  is the corresponding shape function matrix. Note that the displacement approximation for elements in the approach is the same as in the classic FEM and therefore shape function matrix for force vectors is constructed in the same way as in the classic FEM. 214 Substituting the approximations in (13), (14) and (15) into the spatially continuous variational

215 principle (5) results in the fully discrete problem:

$$\min_{\Delta \hat{\mathbf{u}}} \max_{\boldsymbol{\sigma}_{n+1}, \boldsymbol{\bar{r}}_{n+1}} \Delta \hat{\mathbf{u}}^{\mathrm{T}} \mathbf{B}^{\mathrm{T}} \boldsymbol{\bar{\sigma}}_{n+1} - \frac{1}{2} \Delta \boldsymbol{\bar{\sigma}}_{n+1}^{\mathrm{T}} \mathbf{C} \Delta \boldsymbol{\bar{\sigma}}_{n+1} - \frac{1}{2} \Delta t^{2} \boldsymbol{\tilde{r}}_{n+1}^{\mathrm{T}} \mathbf{D}_{r} \boldsymbol{\tilde{r}}_{n+1} + \Delta \boldsymbol{\hat{u}}_{n+1}^{\mathrm{T}} \mathbf{A}^{\mathrm{T}} \boldsymbol{\tilde{r}}_{n+1} - \Delta \hat{\mathbf{u}}^{\mathrm{T}} \mathbf{\tilde{f}}$$
216 subject to  $F(\boldsymbol{\bar{\sigma}}_{n+1}) \leq 0$ 
(16)

217 where

$$\mathbf{B}^{\mathrm{T}} = \int_{\Omega} (\mathbf{\bar{B}})^{\mathrm{T}} \mathbf{N}_{\sigma} d\Omega, \qquad \mathbf{C} = \int_{\Omega} \mathbf{N}_{\sigma}^{\mathrm{T}} \mathbb{C} \mathbf{N}_{\sigma} d\Omega,$$
218
$$\mathbf{D}_{r} = \int_{\Omega} (\mathbf{N}_{u})^{\mathrm{T}} \hat{\rho}^{-1} \mathbf{N}_{u} d\Omega, \quad \mathbf{A}^{\mathrm{T}} = \int_{\Omega} (\mathbf{N}_{u})^{\mathrm{T}} \mathbf{N}_{u} d\Omega,$$

$$\tilde{\mathbf{f}} = \int_{\Omega} \mathbf{N}_{u}^{\mathrm{T}} \hat{\mathbf{b}}_{n} \mathbf{N}_{\sigma} d\Omega + \int_{\Gamma} \mathbf{N}_{u}^{\mathrm{T}} \hat{\mathbf{t}}_{n} d\Gamma - \frac{1 - \theta_{1}}{\theta_{1}} \mathbf{\bar{B}}^{\mathrm{T}} \mathbf{\bar{\sigma}}_{n}$$
(17)

219 It is noteworthy that all integration operations for equations in (17) are carried out on cells using the nodal-integration technique rather than on finite elements using Gauss integration technique. Thus 220 221 information on all the variable states, such as displacements, velocities, strains, stresses, etc., is 222 stored on mesh nodes. A proper treatment with boundaries in the numerical model is essential. Inspired by the recently proposed framework for the discrete element method (Krabbenhoft et al., 223 2012, Meng et al., 2018, 2019a), the purely frictional and cohesive-frictional contact interfaces are 224 225 accounted for. As indicated in Fig. 2, frictional/cohesive-frictional behaviour is considered for 226 yellow smoothing domains in contact with the boundary while red smoothing domains have 227 potential contacts. To prevent the penetration into the boundary, the following non-penetration 228 condition is employed:

229 
$$g^{I} = g_{0}^{I} + \left(\Delta \hat{\mathbf{u}}^{I}\right)^{\mathrm{T}} \boldsymbol{n}^{I} \ge 0$$

$$p^{I} g^{I} = 0$$
(18)

- where  $\Delta \hat{\mathbf{u}}^{I}$  is the displacement increments of the node at contact *I*,  $\mathbf{n}^{I}$  is the outward normal vector of the boundary,  $p^{I}$  is the contact force from the boundary,  $g_{0}^{I}$  is the initial gap and  $g^{I}$  is the gap at the next step.
- 233



Fig. 2. The boundary condition for a deformable body. Smoothing domains are shown with dashlines. Smoothing domains with frictional/cohesive-frictional interfaces are coloured in yellow.

237

Following the approach in Meng et al. (2019b), the condition (18) can be enforced into the principleleading to:

$$\min_{\Delta \hat{\mathbf{u}}} \max_{\bar{\sigma}_{n+1}, \bar{r}_{n+1}, p, q} \Delta \hat{\mathbf{u}}^{\mathrm{T}} \mathbf{B}^{\mathrm{T}} \bar{\sigma}_{n+1} - \frac{1}{2} \Delta \bar{\sigma}_{n+1}^{\mathrm{T}} \mathbf{C} \Delta \bar{\sigma}_{n+1} - \frac{1}{2} \Delta t^{2} \tilde{r}_{n+1}^{\mathrm{T}} \mathbf{D}_{r} \tilde{r}_{n+1} + \Delta \hat{\boldsymbol{u}}_{n+1}^{\mathrm{T}} \mathbf{A}^{\mathrm{T}} \tilde{r}_{n+1} - \Delta \hat{\mathbf{u}}^{\mathrm{T}} \tilde{\mathbf{f}} - \Delta \hat{\mathbf{u}}^{\mathrm{T}} \left( \boldsymbol{np} + \hat{\boldsymbol{n}} q \right) - \sum_{l=1}^{N_{b}} g_{0}^{l} p^{l}$$
subject to
$$F(\bar{\sigma}_{n+1}) \leq 0$$

$$F_{b}(\boldsymbol{p}, \boldsymbol{q}) \leq 0$$
(19)

where  $N_b$  is the number of boundary contacts, the normal and tangential vectors of the boundaries are collected in n and  $\hat{n}$ , respectively, contact forces in the normal and tangential directions are organised into vectors p and q, respectively, and shear strength for boundary contacts is considered

244 with a constraint (i.e.,  $F_b(\boldsymbol{p}, \boldsymbol{q}) \leq 0$ ) as:

245 
$$\begin{cases} |\boldsymbol{q}| \le \mu \boldsymbol{p}, & \text{purely frictional interfaces} \\ |\boldsymbol{q}| \le \mu \boldsymbol{p} + c\boldsymbol{A}, \text{ cohesive-frictional interfaces} \end{cases}$$
(20)

where  $\mu$  is the friction coefficient, *c* is the cohesion of the shear strength and *A* is the vector collecting all cohesive interfaces' area. Principle (19) is thus the discretised optimisation problem for the nodal-integration based finite element method with contacts (N-FEM). Following the procedure in the appendix, principle (19) can be cast into a standard second-order cone program. Thus, efficient standard second-order cone programming solvers can be employed.

# **3. Implementation of the nodal-integration based PFEM (N-PFEM)**

252 Similar to the idea of the PFEM for geomechanics problems in Zhang et al. (2013), the key feature 253 of the modelling procedure of the N-PFEM is that nodes are viewed as free "particles" that can 254 move and even separate from the domain to which they originally belong. On the basis of the 255 particle distribution, the computational domain is re-identified and discretised in space at each time 256 step and then the governing equations are resolved. Nevertheless, compared to the classical PFEM, 257 the proposed N-PFEM does not require the operation of variable mapping when it is used for solving 258 geomechanical problems. This is owing to the fact that has been indicated in Section 2.2 that 259 information of all field variables is stored on mesh nodes and the nodal integration is performed. Briefly, the computational cycle of the N-PFEM for a typical time interval is as follows (see also Fig. 260 261 3):

- 262 1. The computational domain is represented by a cloud of particles at time  $t_n$ .
- 263 2. The alpha-shape technique is adopted to identify the boundary of the computational domain and
- then triangle meshes are generated based on the particles and identified boundary;
- 265 3. Cells (e.g. smoothing domains) are constructed according to the topology of the triangle meshes,
- 266 on which the N-FEM is resolved to obtain the state of field variables at  $t_{n+1}$ ;
- 267 4. Update the position of the particles to form the new cloud of particles;
- 268 5. Loop the above process over all time steps.
- 269



# **4. Validation on a sliding block**

For the purpose of verification, the developed approach is adopted to simulate a block (2 m long and 1 m high) that slides on a slope as shown in Fig. 4 (a). Model parameters for the simulation include the unit weight of the block of 20 kN/m<sup>3</sup> and the inclined slope angle of 45° and 60°. The time step

used in the simulation is 0.01 s, which is sufficiently small to obtain convergence. Both the Young's modulus and the strength of the block are set to be sufficiently large so that the block behaves nearly as a rigid body. Both purely frictional and cohesive-frictional interfaces between the block and the inclined slope were considered. The frictional angle for the contact interface between block and plane was 20° with three different values of cohesion considered: 0 kPa, 1 kPa and 10 kPa. The analytical solution of the block sliding distance *S* can be derived from first principles:

283 
$$S = 0.5(g\sin(\beta) - g\cos(\beta)\mu - cA/m)t^2$$
(21)

where  $\beta$  is an angle of slope inclination,  $\mu$  is the interface frictional coefficient (tan20°), g is the gravitational acceleration (10 m/s<sup>2</sup>), t is the sliding time, c is the interface cohesion, A is the interface area (2 m<sup>2</sup>) and m is the block mass (4×10<sup>3</sup> kg).

287

288 The simulation results and the analytical solutions of the displacement versus time are shown in Fig.

289 4(b): excellent agreement of the simulation results with the analytical solution can be observed.





Fig. 4. The sliding of a block: (a) geometric model and (b) comparison between numerical and
analytical solutions.

# 294 5. Uniform weathering: N-PFEM Versus Limit Analysis

A vertical uniform slope subjected to the uniform weathering that has been studied using the limit analysis method in Utili and Crosta (2011a) is here considered. The vertical slope is 40 m high with a unit weight of  $20 \text{ kN/m}^3$ . Three different meshes, namely a coarse one (4876 elements), a fine one

298 (10950 elements) and a very fine one (30276 elements) are used to discretise the vertical slope as

illustrated in Fig. 5.

300





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301

Weathering causes a decrease of the strength of the slope materials over time (Gupta and Rao, 2000, Gullà et al., 2006, Tran et al., 2019). Kimmance (1988) reported experimental data of weathering of granites as shown in Fig. 6(a). It is evident that weathering causes a reduction mainly of cohesion and to a much lesser extent of the friction angle. The same phenomenon is observed for hard soils such as cemented sands (Wang and Leung, 2008) as shown in Fig. 6(b). Here, we assumed a weathering induced cohesion only decrease (constant friction) for sake of comparison with the main cliff retreat scenario presented in Utili and Crosta (2011a) in section 3.3. Following Utili and Crosta

- 311 (2011a) the initial cohesion of the geo-material is set to 500 kPa and a constant frictional angle of
- 312 34°, is adopted (Utili and Crosta, 2011a).
- 313



Fig. 6. Mechanical degradation of geo-materials due to weathering: (a) granites (after Kimmance,
1988) and (b) cemented sand (after Wang and Leung, 2008). Different weathered granite samples
have been adopted for testing and indicated with SB1, SB2 and SB3 in Fig 6(a). SB1, SB2 and SB3
samples represent weak brownish grey granites, moderately to highly altered granites and highly
altered granites, respectively.

320 The condition of strong erosion (also called weathering-limited erosion) is here assumed. This 321 means that after each slope failure the landslide debris is removed from the model. The cohesion of the geo-materials is decreased by small steps of by 0.5 kPa. Slope failure occurs once a shear band 322 323 has developed from the slope toe all the way to the slope upper horizontal level as shown in Fig.7. In this study several successive failure events take place, so it is important to identify slope failures in a 324 325 consistent way for all the failures taking place. Common criteria in the literature are in terms of loss 326 of static equilibrated solution for the failing mass or the achievement of a threshold of strain or 327 displacement at some reference points in the slope (Lane & Griffiths, 1999). Here for reasons of 328 ease of computation, the horizontal displacement of the slope crest  $d_{\text{crest}}$ , normalised by the slope 329 height, was considered:  $I = d_{crest}/H$ . The threshold value for failure assumed in our simulations was  $I_{\text{failure}} = 3.5 \times 10^{-3}$ . After a slope failure is identified, the failed slope mass is deleted to replicate strong 330 331 erosion (the marked nodes in Fig. 7). Mesh convergence of the evolution processes produced by the 332 proposed method is examined and the simulation results are also compared to those from limit analysis in Utili and Crosta (2011a). Fig. 8 shows the relationship between the normalised cohesion 333 334  $c/(\gamma H)$  and the crest retreat normalised by slope height CR/H, where c is the cohesion,  $\gamma$  is the unit weight, H is the slope height and CR is the crest retreat distance. The steps shown in the figure result 335 336 from a sequence of discrete landslide events. Each landslide event leads to a finite retreat of cliff, 337 corresponding to horizontal lines. Between landslide events the slope is stable with the ground strength parameters progressively being degraded due to weathering captured by the vertical lines in 338 339 Fig. 8.

A satisfactory agreement between the results from the proposed method and the limit analysis solutions from Utili and Crosta (2011a) can be observed for the first four failure mechanisms which are enough to track how the cliff profile evolves due to the weathering action. The little discrepancy between two methods is to be attributed to the following factors: 1) in the limit analysis model of 344 (Utili and Crosta, 2011a) a logarithmic spiral failure surface is assumed, whereas in our approach 345 the failure surface is obtained as result of the simulations without any predefined assumption on its 346 shape; and 2) some approximation in identifying the failure surface in our approach due to the FEM space discretization as shown in Fig. 7. 347





Fig. 7. Shear band at the onset of the first failure. Colours are proportional to equivalent plastic strain increment. The mesh marked with nodes is considered as failure mass. 350



351

352 Fig. 8. Uniform weathering results (dimensionless normalised cohesion and the crest retreat).

With regard to the dependence of the results from mesh size, the predicted cliff retreat from the fine mesh and the very fine mesh are identical for all practical purposes. However the value of cohesion at failure is not the only feature to check since the geometry of the failure mechanism is also important. The slope profiles after the fourth failure predicted by the proposed method using different meshes are illustrated in Fig. 9. It is evident that the shapes of the slope profile left are very similar. Thus the fine mesh will be used to simulate the case of the non-uniform degradation in the following sections for the sake of computational efficiency.





Fig. 9. The slope shape after the fourth failure subject to uniform weathering with (a) a coarse mesh,(b) a fine mesh and (c) a very fine mesh.

# 365 **6. Non-uniform degradation**

In this section, a more realistic weathering scenario, namely the non-uniform weathering over time and space, is considered. Additionally, the transport-limited condition is here investigated. The condition implies that the debris accumulates at the toe of the slope after each landslide event (Utili and Crosta, 2011b). Therefore both the onset of failure and the movement of the debris need to be correctly simulated for the model to capture the evolution of the cliff investigated. 371 In light of field hardness tests taken at different depth from the exposed surface of a cliff subject to 372 weathering (Yokota and Iwamatsu, 2000), a weathering front parallel to the exposed surfaces 373 moving inwards at a constant rate is here assumed as shown in Fig. 10. The weathering depth is 374  $W_{\text{front}} = v_{\text{front}} \times t_{real}$  where  $v_{\text{front}}$  is the speed of the weathering front and  $t_{real}$  is the physical time. The damage index D is used to quantify the degradation of the cliff subjected to weathering. Specifically, 375 for intact rocks, the damage index is D=0, while we have D=1 if the rocks are fully weathered. 376 377 Currently, data related to quantitative descriptions on strength degradation within natural slopes 378 induced by weathering are scanty. Thus, cohesion is reduced to zero for fully weathered geo-379 materials to validate the approach against the predictions of cliff evolution from Discrete Element 380 simulations reported in Utili and Crosta (2011b). The decrease of cohesion over time is given by:

$$^{381}$$
  $c(t)=c \times (1-D)$  (21)

W<sub>D</sub> as shown in Fig. 10 is the width of partially weathered zone, in which the damage index D
changes linearly from 1 to 0. W<sub>D</sub> is given by:

$$W_{\rm D} = v_{\rm front} / v_{\rm D} \tag{22}$$

where  $v_D$  is the damage rate, i.e. the increment of D over time. In the simulations, the damage index inside the slope is prescribed as shown in Fig. 10. The shear strength inside the slope can be determined from Eq. (21).

Since the spatial distribution of weathering changes over time, it is necessary to explicitly assume an initial condition to simulate the progression of degradation with time unlike the case of uniform degradation. The initial condition assumed at  $t_0$ , was that of unweathered material: D=0 throughout the whole slope. This assumption supposes the existence of a time when the cliff was characterised by a uniform strength which could be thought as the time of formation of the cliff: for instance, the formation of a scarp or a hillslope because of a deep-seated landslide, a series of rapid displacements

<sup>394</sup> along a specific plane (e.g. fault) or erosion and deposition of river terraces. This is also the standard
 <sup>395</sup> assumption in the Fisher-Lehmann and Bakker-Le Heux models.



396



Fig. 10. The weathering process within the cliff.

Weathering is usually very slow. Therefore the weathering rate needs to be speeded up to run the simulations in a feasible computation time as indicated in Utili and Crosta (2011b). This means that a fictitious simulation time needs to be employed. The ratio between the fictitious time in the simulation  $t_{sim}$  and the real time  $t_{real}$  is:

$$402 t_{real} = t_{sim} \times \kappa (23)$$

403 where  $\kappa$  is the speedup ratio defined as the ratio of the physical time to the simulation time.

405 The cliff model considered in Utili and Nova (2008) and Utili and Crosta (2011a) is used in this study which is also shown in Fig. 12 (a). The velocity of the weathering front is  $v_{\text{front}} = 0.01 \text{ m/year}$ 406 and W<sub>D</sub> is 2 m. The material parameters are the same as these for the case of uniform weathering 407 408 except that the non-associated plastic flow with a null dilation angle is assumed for the geo-material in this section. To calibrate the speedup ratio,  $\kappa$  is set to  $0.50 \times 10^{10}$ ,  $1.00 \times 10^{10}$  and  $1.58 \times 10^{10}$ 409 410 corresponding to a speedup velocity of the weathering front of 1.59, 3.17 and 5.00 m/s, respectively. 411 Note that as soon as dynamic motion in any point in the slope is identified, weathering is stopped 412 and the speed up ratio is set to unity, i.e. the real time coincides with the simulation time, for the 413 whole duration of the dynamic motion until the landslide body stops its motion. Then, once quasi-414 static conditions are resumed, weathering scaling is also resumed with the speed up ratio set to its original value until the next failure occurs. One simulation, consisting of a sequence of discrete 415 416 landslide events and slow weathering process, required around 62.6 hours to complete.

The results of the simulation in terms of normalised crest retreat versus propagation distance of the Weathering Front (WF) for WF/H<0.4 are shown in Fig. 11. The crest retreats obtained for  $\kappa =$ 0.50×10<sup>10</sup> and  $\kappa = 1.00 \times 10^{10}$  are in a good agreement, indicating that the speedup ratio  $\kappa = 1.00 \times 10^{10}$ is sufficiently small not to unduly affect the results. This ratio was therefore adopted in all subsequent simulations.



423 Fig. 11. Dimensionless crest retreat versus weathering front under different speedup ratios. WF is
424 the propagation distance of the weathering front.

The predicted cliff profiles over time are shown in Figs. 12 (b)-(i) with the associated weathering 425 426 fronts plotted in Fig.13. As it can be seen in Fig.13 (i), at the end of the cliff geometric evolution, 427 most of the slope is fully weathered and the final angle of repose of the scree made of debris is about 30° over the horizontal. In Fig. 12, the debris materials are depicted in blue and the so-called non-428 429 displaced undisturbed zone is in red. It emerges that the free cliff front (e.g. the surface from the 430 crest of the slope to the rear of the accumulated debris) remains parallel to the original cliff surface 431 which is represented by a dashed line. This phenomenon produced by the N-PFEM simulation 432 confirms the assumption made by the Fisher-Lehmann model (Fisher, 1866, Lehmann, 1933), which is a classic geomorphologic model to predict the shape of the undisturbed zone of slopes subjected 433 to weathering, that in a given time weathering produces an equal retreat of all parts of the exposed 434 435 free face by the falling away of fine debris. Readers interested in the assumption of this 436 geomorphologic model and the equation of the shape of the undisturbed zone are referred to





Fig. 12. The evolution of the slope subjected to weathering propagation at simulation time of (a) 0.0
s, (b) 1.5 s (c) 1.8 s, (d) 2.4 s (e) 3.3 s (f) 4.1 s (g) 5.4 s (h) 8.2 s and (i) 16.0 s. Colours are
proportional to the accumulated horizontal displacement (m). The blue indicates the sliding mass
while the undisturbed zone in the slope is indicated in red.

443



Fig. 13. The propagation of weathering front at simulation time of (a) 0.0 s, (b) 1.5 s (c) 1.8 s, (d)
2.4 s (e) 3.3 s (f) 4.1 s (g) 5.4 s (h) 8.2 s and (i) 16.0 s. Colours are proportional to the damage index,
D. The green indicates the intact mass while the fully weathered materials are in yellow.

In Fig.14 the final deposit predicted by the N-PFEM method (Fig. 12 (i)) is compared to the DEM simulations in Utili and Crosta (2011b), yellow curves, and that from the Fisher-Lehmann geomorphologic model (Fisher, 1866, Lehmann, 1933). The results of the three methods broadly agree well with each other. It can be seen that the proposed N-PFEM and the DEM produce very similar results in terms of both the shape of the undisturbed zone and the slope profile. An advantage of the N-PFEM model over the DEM is that, as a continuum approach, it does not require the calibration process needed for the particle bond parameters of the DEM modelling.



458 Fig.14. Final deposit of the slope derived from the N-PFEM, the DEM (Utili and Crosta, 2011b) and
459 the Fisher-Lehmann model.

#### 460 **7. Retaining wall design**

457

461 In order to prevent the cliff recession due to weathering induced landslides, one of the measures 462 commonly taken in practical engineering is the construction of retaining wall structures. In this 463 study, the performance of the retaining wall against weathering induced cliff failures is shown in Fig. 464 15. The higher the retaining wall, the safer the slope but the associated construction and 465 maintenance costs will also be higher. Therefore, it is important to strike the right balance with respect to the wall construction costs. For this purpose, the performance of retaining wall with 466 467 different heights (i.e. 0 m, 5 m, 10 m, 15 m, 20 m to 25 m) is investigated using the proposed N-468 PFEM. As a matter of fact, the retaining wall will eventually be degraded due to weathering in the 469 very long term. In our simulations the rigid retaining wall is modelled as rough boundary condition. 470 The parallel weathering model described for the non-uniform weathering case in section 6 was 471 adopted together with the slope material parameters. A faster weathering rate is considered with 472  $v_{\text{front}} = 0.1 \text{ m/year}$ ;  $W_D$  is 2 m and the speedup ratio  $\kappa$  is reduced to  $1.00 \times 10^9$ .





Fig.15. The numerical model with a retaining wall structure.

The evolution of the slope subjected to weathering is shown in Fig. 16 while the simulation results for the crest retreat over time are shown in Fig. 17 in which the height of the retaining wall is normalized by the height of the slope, h/H, and so does the distance of crest retreat (i.e. CR/H).

478 Fig. 17 indicates that the first crest retreat event occurs at about 57.1 years (simulation time = 1.5 s) 479 regardless of the height of the retaining wall. There is also no clear difference on the crest retreat 480 distance between different cases until the slope with the highest retaining wall reaches its final 481 deposit profile at 199.8 years (simulation time = 6.3 s). The detached debris covers the whole 482 retaining wall for the lower retaining wall structure (i.e.  $h/H \le 0.375$ ) implying that the effect of the 483 retaining wall on the shape of the scree resting on the slope base is negligible. In addition, the crest 484 retreat distances for these four cases are comparable. The slope with second highest retaining wall 485 (h/H=0.5) reaches its final deposit profile at 255.3 years (simulation time = 8.05 s). At the final stage, 486 the slopes with  $h/H \ge 0.5$  has a much lower crest retreat while, for the cases with low retaining walls 487 (i.e.  $h/H \le 0.375$ ), the effect of the retaining wall on the crest retreat is insignificant.



Fig. 16. Evolution of the cliff with different retaining wall heights at different simulation time
instants: (a) 1.8 s, (b) 6.3 s, (c) 8.05 s and (d) 16.0 s. Colours are proportional to accumulated the
horizontal displacement (m) and black dash lines indicate crest retreat of the cliff without the
retaining wall.



Fig. 17. Dimensionless crest retreat over time.

# 496 8. Discussion and conclusions

497 A novel computational framework called the Nodal-integration based Particle Finite Element 498 Method (N-PFEM) was developed for modelling cliff recession resulting from weathering-caused 499 landslides. To this end, the nodal integration technique is first embedded into the finite element 500 formulation in second-order cone programming. The resulting nodal-integration based finite element 501 formulation (N-FEM) is then merged into the framework of the particle finite element method (PFEM) to form the N-PFEM for simulating cliff recession involving very large material 502 503 deformations. Comparing to the classical PFEM, the developed N-PFEM stores information of all 504 field variables on mesh nodes and performs nodal integration on the cell associated with each mesh node. Consequently, it requires no operation of variable mapping from old meshes to new meshes, 505 when new meshes are generated, which is essential in the classical PFEM for modelling history-506

507 dependent materials and inevitably introduces errors and considerable complexity to solution508 procedures.

509 The ability of the proposed N-PFEM for predicting cliff recession in weathering has been 510 showcased through a series of examples. The results from the N-PFEM modelling have been compared with these from the limit analysis, the DEM simulation and the Fisher-Lehmann 511 512 geomorphologic model available in literatures. Analytical methods, such as the limit analysis 513 method and the Fisher-Lehmann geomorphologic model, can provide a quick estimation of the slope 514 profile. But the limit analysis approach can only handle the weathering-limited case and the shape of 515 the failure surface has to be defined in advance. The simulation results from the N-PFEM and the 516 DEM agree well with each other in terms of both the final deposit shape and the shape of the 517 undisturbed zone of the slope. In comparison with the DEM, the appeal of the N-PFEM is that the 518 time onerous calibration of the material properties at microscopic levels is not required.

The prevention of cliff recession by retaining wall structures is also studied using the developed N-PFEM. Since the purpose of this example is to illustrate the capability of the proposed N-PFEM for studying the effect of a retaining wall on cliff erosions, we only consider a retaining wall of a simple geometry. Further studies on the influences of the geometry and the engineering properties of a retaining wall that may change during the weathering process are critical for the optimum design of retaining walls in practice and achievable using the N-PFEM.

525

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#### 529 Appendix: Second-order cone programming

530 In this work, optimisation problem (18) is transformed into a standard SOCP which is then can be 531 resolved using the advanced interior-point method. Very efficient solvers capable of dealing with 532 large-scale SOCP problems have been developed in last decades or so. Of particular notes are the 533 packages MOSEK (Mosek, 2015) and SeDuMi (Sturm, 1999).

534

535 The SOCP is a generalization of linear and quadratic programming that allows for affine 536 combinations of variables to be constrained inside a special convex set, called second-order cone 537 (Calafiore and Ghaoui, 2014). The following primal standard form of the SOCP is often used:

$$\begin{array}{ccc}
\min & \mathbf{a}^{\mathrm{T}} \mathbf{y} \\
538 & \text{subject to} & \mathbf{D} \mathbf{y} = \mathbf{e} \\
& \mathbf{y} \in \mathcal{K}
\end{array} \tag{A1}$$

539 where *y* are the full problem variables and  $\mathcal{K}$  is a Cartesian product of second-order cones i.e., 540  $\mathcal{K} = \mathcal{K}_1 \times \mathcal{K}_2 \times \cdots \times \mathcal{K}_n$ . Two most common conic cones are:

• the quadratic cone:

542 
$$\mathcal{K}_q = \left\{ \mathbf{y} \in \mathbb{R}^m \mid y_1 \ge \sqrt{y_2^2 + \dots + y_m^2} \right\}$$
 (A2)

• the rotated quadratic cone:

544 
$$\mathcal{K}_r = \left\{ \mathbf{y} \in \mathbb{R}^m \mid 2y_1 y_2 \ge y_3^2 + \dots + y_m^2, y_1 \ge 0, y_2 \ge 0 \right\}$$
 (A3)

545 The minimisation part of principle (18) with respect to  $\Delta \hat{\mathbf{u}}$  can be solved analytically resulting in a 546 maximisation problem:

$$\max_{\bar{\sigma}_{n+1}, \tilde{r}_{n+1}, p, q} -\frac{1}{2} \Delta \bar{\sigma}_{n+1}^{T} \mathbf{C} \Delta \bar{\sigma}_{n+1} - \frac{1}{2} \Delta t^{2} \tilde{r}_{n+1}^{T} \mathbf{D}_{r} \tilde{r}_{n+1} - \sum_{l=1}^{N_{b}} g_{0}^{l} p^{l}$$
subject to  $\mathbf{B}^{T} \bar{\sigma}_{n+1} + \mathbf{A}^{T} \tilde{r}_{n+1} = \tilde{\mathbf{f}} + np + \hat{n}q$ 

$$F(\bar{\sigma}_{n+1}) \leq 0$$
547
$$F_{b}(p, q) \leq 0$$
(A4)

548 Obviously, this maximum problem is equivalent to the following minimum problem:

$$\min_{\overline{\sigma}_{n+1}, \overline{r}_{n+1}, p, q} \quad \frac{1}{2} \Delta \overline{\sigma}_{n+1}^{\mathrm{T}} \mathbf{C} \Delta \overline{\sigma}_{n+1} + \frac{1}{2} \Delta t^{2} \widetilde{r}_{n+1}^{\mathrm{T}} \mathbf{D}_{r} \widetilde{r}_{n+1} + \sum_{I=1}^{N_{b}} g_{0}^{I} p^{I}$$
subject to  $\mathbf{B}^{\mathrm{T}} \overline{\sigma}_{n+1} + \mathbf{A}^{\mathrm{T}} \widetilde{r}_{n+1} = \widetilde{\mathbf{f}} + np + \hat{n}q$ 

$$F(\overline{\sigma}_{n+1}) \leq 0$$

$$F_{b}(\mathbf{p}, \mathbf{q}) \leq 0$$
(A5)

550 Comparing (A5) to the standard SOCP form (A1), the quadratic terms, namely 551  $\frac{1}{2}\Delta \bar{\sigma}_{n+1}^{T} \mathbf{C}\Delta \bar{\sigma}_{n+1}$  and  $\frac{1}{2}\Delta t^{2} \tilde{\mathbf{r}}_{n+1}^{T} \mathbf{D}_{r} \tilde{\mathbf{r}}_{n+1}$ , in the objective function have to be reformulated. To this end,

552 two auxiliary variables X and Y are introduced in the objective function and we have

$$\min_{\overline{\sigma}_{n+1}, \overline{r}_{n+1}, p, q} \quad X + I + \sum_{I=1}^{N_b} g_0^I p^I$$
subject to  $\mathbf{B}^{\mathrm{T}} \overline{\sigma}_{n+1} + \mathbf{A}^{\mathrm{T}} \widetilde{r}_{n+1} = \widetilde{\mathbf{f}} + np + \hat{n}q$ 

$$X \ge \frac{1}{2} \Delta \overline{\sigma}_{n+1}^{\mathrm{T}} \mathbf{C} \Delta \overline{\sigma}_{n+1}$$

$$I \ge \frac{1}{2} \Delta t^2 \widetilde{r}_{n+1}^{\mathrm{T}} \mathbf{D}_r \widetilde{r}_{n+1}$$

$$F(\overline{\sigma}_{n+1}) \le 0$$

$$F_b(p, q) \le 0$$
(A6)

553

554 The newly introduced inequality constraints can be converted to rotated quadratic cones:

$$\min_{\overline{\sigma}_{n+1}, \overline{r}_{n+1}, p, q} X + I + \sum_{l=1}^{N_b} g_0^l p^l$$
subject to  $\mathbf{B}^T \overline{\sigma}_{n+1} + \mathbf{A}^T \widetilde{r}_{n+1} = \widetilde{\mathbf{f}} + np + \hat{n}q$ 

$$\boldsymbol{\xi}_{\overline{\sigma}} = \mathbf{C}^{\frac{1}{2}} \Delta \overline{\sigma}_{n+1}, Y = 1, (X, Y, \boldsymbol{\xi}_{\overline{\sigma}}) \in \mathcal{K}_r$$

$$\mathcal{K}_r = \left\{ (X, Y, \boldsymbol{\xi}_{\overline{\sigma}}) \in \mathbb{R}^{m+2} \mid 2XY \ge \boldsymbol{\xi}_{\overline{\sigma}}^T \boldsymbol{\xi}_{\overline{\sigma}}, X \ge 0, Y \ge 0 \right\}$$

$$\boldsymbol{\xi}_{\overline{r}} = \Delta t \mathbf{D}_r^{\frac{1}{2}} \widetilde{r}_{n+1}, J = 1, (I, J, \boldsymbol{\xi}_{\overline{r}}) \in \mathcal{K}_r$$

$$\mathcal{K}_r = \left\{ (I, J, \boldsymbol{\xi}_{\overline{r}}) \in \mathbb{R}^{m+2} \mid 2IJ \ge \boldsymbol{\xi}_{\overline{r}}^T \boldsymbol{\xi}_{\overline{r}}, X \ge 0, Y \ge 0 \right\}$$

$$F(\overline{\sigma}_{n+1}) \le 0$$

$$F_b(p, q) \le 0$$
(A7)

The yield criteria  $F(\bar{\sigma}_{n+1}) \le 0$  and  $F_b(p, q) \le 0$  can be reformulated as a quadratic cone as well. Readers are referred to (Zhang et al., 2013) for more details. Problem (A7) is now the formulation of the N-FEM in a standard SOCP form which is the eventual problem to be resolved at each incremental analysis step.

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