

Long-term storage of riverine dissolved organic carbon

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Abstract

The effect of long-term (2 years) storage on dissolved organic carbon concentrations of filtered river water was investigated. Replicate samples were stored at 3 temperatures 20°C, 4°C or -20°C, either acidified or without acidification. The effects of sample volume were tested by storing 4 ml or 20ml samples at each temperature and again acidified or non-acidified. Only frozen, acidified samples were stable for up to 1 year and none of the treatments were suitable for storing samples for 2 years. The volume of sample stored did not influence DOC measurements.

Introduction

Approximately 1% of terrestrial primary production is exported annually by rivers to the coastal ocean as dissolved organic carbon (DOC), and this DOC flux is closely correlated with river discharge (Harrison et al., 2005; Cole et al., 2007; Worrall and Burt, 2007). Over recent decades there is growing evidence of rising DOC concentrations in central and northern Europe and North American rivers and lakes (Freeman et al., 2001; Evans et al., 2005; Skjelkvåle et al., 2005; Gedney et al., 2006; Dawson et al., 2009; Hruška et al., 2009). The reasons for these increases are subject to debate but are linked to primary productivity changes in catchments as a result of elevated atmospheric CO₂ (Freeman et al., 2004), increasing mobility of DOC within soils (Evans et al., 2006), as well as organic carbon loss from top soils (Bellamy et al., 2005). Changes in DOC may also be explained by changes in deposition chemistry and/or sensitivity to catchment acidity (Monteith et al., 2007; Dawson et al., 2009; Hruška et al., 2009). Whatever the causes of such large-scale changes, they not only have implications to long-term regional carbon balance (Cole et al., 2007), but also for the dynamics and biogeochemical cycling of the dissolved

organic matter (DOM) pool and inorganic nutrients during their transit from catchment to coast (Søndergaard and Thomas, 2003; Mattsson et al., 2009). Clearly reliable DOC measurements are crucial to this debate and the establishment of reliable long-term data sets.

The DOC pool is made up of both labile and refractory material. The refractory component can remain within aquatic systems for decades, hundreds or even thousands of years (Battin et al., 2008), whereas the labile material is degraded both photo-chemically and biologically over much shorter time periods of days or even hours (Wangersky, 1993; Battin et al., 2008). Given the potential for rapid change in labile DOC composition, even samples analyzed within hours of collection are potentially vulnerable to a loss of DOC, and reliable preservation methods are required to minimize losses. Ideally analysis of DOC samples should occur immediately after sample collection; however, during extended fieldwork campaigns or seasonal sampling programs this is not always possible. For these and other reasons storage conditions that will reliably preserve samples for longer than a few months is vital. Although there has been a considerable effort to refine the analytical methods associated with DOC analysis over the past 20

years, to our knowledge only 3 publications deal with the key issue of long-term storage of DOC samples up to 5 months studies (Sugimura and Suzuki, 1988; Peltzer and Brewer, 1994; Tupas et al., 1994) and these were within seawater samples. Here we report a brief study that addressed the issue of long-term storage (up to 2 years) of DOC samples, and investigated the merits of acid preservation *vs* non-acidified storage, under ambient laboratory temperatures, cold storage and freezing temperatures. We also compared the relative merits of different sample volumes for the suitability of long-term storage.

Materials and Procedures

Sampling was conducted on 30 January 2008 in the River Conwy, North Wales. Samples were collected by immersing a 2L acid cleaned high density polyethylene bottle approximately 20 cm below the water surface, after rinsing the bottle 3 times with sample water. The samples were returned to the laboratory within 0.5 h.

All samples were filtered directly into either 4 mL borosilicate vials or 20 mL ampoules (both pre-combusted at 500°C, 3 hours) through disposable syringe filters (Whatman[®] GD/X, pore size 0.45 µm) using an acid-cleaned plastic syringe. It is realised that with a pore size of 0.45 µm these filters do not exclude all bacteria, but routinely DOC concentrations in the literature are reported from filtrations using pre-combusted glass fibre filters (typically Whatman[®] GF/F) of effective pore sizes greater than 0.6 µm (Tipping et al., 1988; Curtis and Adams, 1995; Hessen et al., 1997; Schindler et al., 1997; Gergel et al., 1999; Pace and Cole, 2002; Pastor et al., 2003; Xenopoulos et al., 2003; Spencer et al., 2007) and so it was realistic to work with such samples treated in this manner. Additionally the contribution of sub-micrometre (< 1 µm) particles to DOC has been estimated to be only approximately 10% (Isao et al., 1990, and references therein).

Replicate samples were prepared for DOC analysis (completed within 4 h of collection): 5 replicates were acidified with ultrapure HP₃O₄, and 5 with of 37% HCl. The DOC concentrations were not significantly different at $422 \pm 3 \mu\text{mol L}^{-1}$ and $419 \pm 3 \mu\text{mol L}^{-1}$ respectively. Therefore in subsequent acidification of samples within this study HP₃O₄ was used.

Two set of samples were prepared: The first was acidified to a pH of 2.0 and the second was left non-acidified. The vials were then sealed with Teflon-lined screw caps, and the ampoules flame sealed.

Replicate acidified and non-acidified vials and ampoules were then stored in the dark at the following 3 storage temperatures:

- 1) Ambient laboratory ($20 \pm 5 \text{ }^\circ\text{C}$)
- 2) Cold storage ($4 \pm 2 \text{ }^\circ\text{C}$),
- 3) Frozen ($-20 \pm 2 \text{ }^\circ\text{C}$).

In the case of the ambient samples only acidified samples were stored. N.B. all sub-sampling was completed within 3 hours of collection

Two subsequent sets of analyses were completed in January 2009 and January 2010. The samples stored in vials were directly used (after shaking) in the auto-sampler of the DOC analyser. The ampoules were thoroughly mixed before opening, and a sub-sample removed from each one and transferred to pre-combusted 4 mL vials. All non-acidified samples were acidified with HP₃O₄ prior to DOC analyses.

DOC was measured using high temperature combustion on an MQ 1001 TOC Analyser (Qian and Mopper, 1996). Certified reference material (CRM) of deep Florida Strait water and low carbon water ($5\text{--}10 \mu\text{mol L}^{-1}$) from the Hansell Laboratory, University of Miami, Rosenstiel School of Marine and Atmospheric Science are run on the machine daily during routine operation. Three batches of CRM were used during this experimental period with certified DOC concentrations of 47-48 µmol, 41-44 µmol and 41-43 µmol L⁻¹. The respective DOC concentrations measured from

these standards batches were $48 \pm 3 \mu\text{mol L}^{-1}$ (n=92), $44 \pm 6 \mu\text{mol L}^{-1}$, (n=64) and $44 \pm 4 \mu\text{mol L}^{-1}$ (n=52). The detection limit of the instrument was $10 \mu\text{mol L}^{-1}$.

Discussion

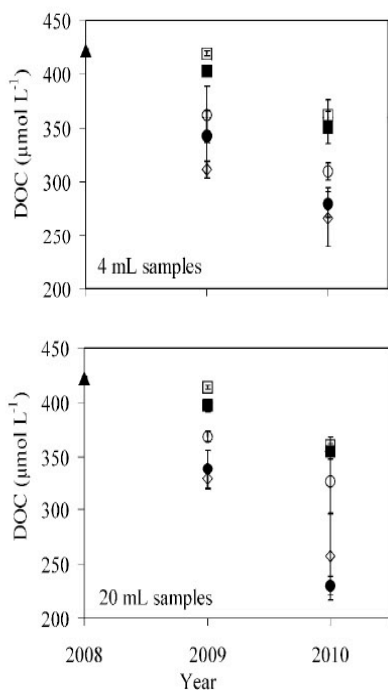


Figure 1 Changes in DOC concentration during a 2 year storage experiment using samples stored in 4 mL vials or 20 mL ampoules, and acidified with HP_3O_4 or left non-acidified. Samples were stored at $20 \pm 5^\circ\text{C}$ (\diamond), $4 \pm 2^\circ\text{C}$, acidified (\circ), non-acidified (\bullet), or $-20 \pm 2^\circ\text{C}$, acidified (\square), non-acidified (\blacksquare). $t = 0$ (\blacktriangle) was the initial measured DOC concentration of $422 \pm 3 \mu\text{mol L}^{-1}$ (n=5).

After 1 year of storage the only samples that did not show a significant decrease in DOC concentration compared to $t=0$ samples were the frozen acidified samples at -20°C (Fig. 1). These losses were 1 and 2% in both the acidified vials and ampoules, respectively, and were not significantly different (tested with a using 2 sample T-Test) from the $t=0$ DOC concentration. There were decreases in DOC concentrations of 5 and 6% in the frozen non-acidified vials and ampoules, and although small these were statistically significantly

different ($p = < 0.001$ in both cases) from the initial measured concentrations (Fig. 1). All other samples stored at ambient or 4°C exhibited too great a loss of DOC to be considered further.

After two years of storage there were significant decreases in DOC concentration in all frozen samples, regardless of whether acid had been added or not. Although these losses were not as great as those observed in the ambient and cold stored samples, the average losses were 14 and 15% in the acidified vials and ampoules, respectively, and 17 and 16% in the non-acidified vials and ampoules, respectively (Fig 1). The fact that this loss happened in both acidified and non-acidified samples would suggest that these losses are not due to bacteria activity since at pH 2 it seems highly unlikely that acidophilic bacteria would be active considering the sample water and low temperatures.

Sugimura and Suzuki (1988) reported a 15 to 20% decrease in DOC concentration in frozen acidified and non-acidified seawater samples over 7 days. Much of this loss occurred during the first hour of storage, after which the DOC concentrations remained relatively constant. In contrast, Tupas et al. (1994) analysed frozen, non-acidified samples up to 5 months after collection and reported that the DOC concentrations measured were not significantly different from those measured in the initial samples. Wangersky (1993) suggests that the loss of DOC observed by Sugimura and Suzuki (1988) may have occurred during sample preparation and freezing, indicating that rapid preservation is required to minimize this effect. The small (5 to 6%) but significant decrease in DOC concentration measured after one year in non-acidified, frozen samples in this study indicates that, as demonstrated by Tupas et al. (1994), that samples may remain stable for a few months, but thereafter reliability of results may become questionable.

There were no significant differences in DOC concentration between samples stored in

either 4 ml vials or 20 ml ampoules, suggesting that the storage sample volume should be chosen primarily to minimize the risk of contamination. These results are in contrast to those of Tupas et al. (1994) who reported consistently elevated DOC concentrations in samples (frozen) stored in 2 mL ampoules compared to those stored in 10 mL ampoules. They concluded that despite both sets of ampoules being flame sealed, the sealing may have contaminated the 2 mL samples.

This study has shown that long-term storage of riverine DOC samples (up to 1 year) is possible, but we recommend that samples are filtered and acidified prior to storage at -20°C, and that these procedures are carried out as soon as possible after sample collection to minimize the loss of labile DOC.

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References

- Battin, T.J., Kaplan, L.A., Findlay, S., Hopkinson, C.S., Marti, E., Packman, A.I., Newbold, J.D., Sabater, F. 2008. Biophysical controls on organic carbon fluxes in fluvial networks. *Nat. Geosci.* 1:95-100
- Bellamy, P.H., Loveland, P.J., Bradley, R.I., Lark, R.M., Kirk, G.J.D. 2005. Carbon losses from all soils across England and Wales 1978–2003. *Nature* 437:245-248.
- Cole, J.J., Prairie, Y.T., Caraco, N.F., McDowell, W.H., Tranvik, L.J., Striegl, R.G., Duarte, C.M., Kortelainen, P., Downing, J.A., Middelburg, J.J., Melack, J. 2007. Plumbing the Global Carbon Cycle: Integrating Inland Waters into the Terrestrial Carbon Budget. *Ecosystems* 10:171-184.
- Curtis, P.J., Adams, H.E. 1995. Dissolved organic matter quantity and quality from freshwater and saltwater lakes in east-central Alberta. *Biogeochemistry* 30:59-76.
- Dawson, J.J.C., Malcolm, I.A., Middlemas, S.J., Tetzlaff, D., Soulsby, C. 2009. Is the composition of dissolved organic carbon changing in upland acidic streams? *Environ. Sci. Technol.* 43:7748–7753.
- Evans, C.D., Monteith, D.T., Cooper, D.M., 2005. Long-term increases in surface water dissolved organic carbon: Observations, possible causes and environmental impacts. *Environ. Pollut.* 137:55-71.
- Evans, C.D., Chapman, P.J., Clark, J.M., Monteith, D.T., Cresser, M.S. 2006. Alternative explanations for rising dissolved organic carbon export from organic soils. *Global Change Biol.* 12:2044-2053.
- Freeman, C., Evans, C.T., Monteith, D.T., Reynolds, B., Fenner, N. 2001. Export of organic carbon from peat soils. Warmer conditions may be to blame for the exodus of peatland carbon to the oceans. *Nature* 412:785.
- Freeman, C., Fenner, N., Ostle, N.J., Kang, H., Dowrick, D.J., Reynolds, B., Lock, M.A., Sleep, D., Hughes, S., Hudson, J. 2004. Export of dissolved organic carbon from peatlands under elevated carbon dioxide levels. *Nature* 430:195-198.
- Gedney, N., Cox, P. M., Betts, R. A., Boucher, O., Huntingford, C., Stott, P. A. 2006. Detection of a direct carbon dioxide effect in continental river runoff records. *Nature* 439:835-838.
- Gergel, S.E., Turner, M.G., Kratz, T.K. 1999. Dissolved organic carbon as an indicator of the scale of watershed influence on lakes and rivers. *Ecol. Appl.* 9:1377-1390.
- Harrison, J.A., Caraco, N., Seitzinger, S.P. 2005. Global patterns and sources of dissolved organic matter export to the coastal zone: Results from a spatially explicit, global model. *Glob. Biogeochem. Cycles* 19:1-15.
- Hessen, D.O., Gjessing, E.T., Knulst, J., Fjeld, E. 1997. TOC fluctuations in a humic lake as related to catchment acidification, season and climate. *Biogeochemistry* 36:139-151.
- Hruška, J., Krám, P., McDowell, W.H., Oulehle, F. 2009. Increased dissolved organic carbon (DOC) in Central European streams is driven by reductions in ionic strength rather than climate change or decreasing acidity. *Environ. Sci. Technol.* 43:2320–2326.
- Isao K, Hara S, Terauchi K, Kogure K. 1990. Role of sub-micrometre particles in the ocean. *Nature* 345:242-243.

- Mattsson, T., Kortelainen, P., Laubel, A., Evans, D., Pujo-Pay, M., Raike, A., Conan, P. 2009. Export of dissolved organic matter in relation to land use along a European climatic gradient. *Sci. Total Environ.* 407:1967-1976.
- Monteith, D.T., Stoddard, J.L., Evans, C.D., de Wit, H.A., Forsius, M., Hogasen, T., Wilander, A., Skjelkvale, B.L., Jeffries, D.S., Vuorenmaa, J., Keller, B., Kopacek, J., Vesely, J. 2007. Dissolved organic carbon trends resulting from changes in atmospheric deposition chemistry. *Nature* 450:537-541.
- Pace, M.L., Cole, J.J. 2002. Synchronous variation of dissolved organic carbon and color in lakes. *Limnol. Oceanogr.* 47:333-342.
- Pastor, J., Solin, J., Bridgham, S. D., Updegraff, K., Harth, C., Weishampel, P., Dewey, B. 2003. Global warming and the export of dissolved organic carbon from boreal peatlands. *Oikos* 100:380-386.
- Peltzer, E.T., Brewer, P.G. 1993. Some practical aspects of measuring DOC - sampling artifacts and analytical problems with marine samples. *Mar. Chem.* 41:243-252.
- Qian, J., Mopper, K. 1996. Automated high-performance, high-temperature combustion total carbon analyser. *Anal. Chem.* 68:3090-3097.
- Schindler, D.W., Curtis, P.J., Bayley, S.E., Parker, B.R., Beaty, K.G., Stainton, M.P. 1997. Climate-induced changes in the dissolved organic carbon budgets of boreal lakes. *Biogeochemistry* 36:9-28.
- Skjelkvale, B.L., Stoddard, J.L., Jefferies, D.S., Torseth, K., Hogasen, T., Bowman, J., Mannio, J., Monteith, D.T., Mosello, R., Rogora, M., Rzychon, D., Vesely, J., Wieting, J., Wilander, A., Worsztynowicz, A. 2005. Regional scale evidence for improvements in surface water chemistry. *Environ Pollut.* 137:165-176.
- Sondergaard, M., Thomas, D.N. 2003. *Dissolved Organic Matter (DOM) in Aquatic Ecosystems: A Study of European Catchments and Coastal Waters*. A publication by the EU project DOMAINE (EVK3-CT-2000-00034), pp 72.
- Spencer, R.G.M., Ahad, J.M.E., Baker, A., Cowie, G.L., Ganeshram, R., Upstill-Goddard, R.C., Uher, G. 2007. The estuarine mixing behaviour of peatland derived dissolved organic carbon and its relationship to chromophoric dissolved organic matter in two North Sea estuaries (U.K.). *Estuarine, Coastal Shelf Sci.* 74:131-144.
- Sugimura, Y., Suzuki, Y. 1988. A high-temperature catalytic oxidation method for the determination of nonvolatile dissolved organic carbon in seawater by direct injection of a liquid sample. *Mar. Chem.* 24:105-131.
- Tipping, E., Hilton, J., James, B. 1988. Dissolved organic matter in Cumbrian lakes and streams. *Freshwater Biol.* 19:371-378.
- Tupas, L.M., Popp, B.N., Karl, D.M. 1994. Dissolved organic carbon in oligotrophic waters: experiments on sample preservation, storage and analysis. *Mar. Chem.* 45:207-216.
- Wangersky, P.J. 1993. Dissolved organic carbon methods: a critical review. *Mar. Chem.* 41:61-74.
- Worrall, F., Burt, T.P. 2007. Flux of dissolved organic carbon from U.K. rivers. *Glob. Biogeochem. Cycles* 21:GB1013, doi:10.1029/2006GB002709.
- Xenopoulos, M.A., Lodge, D.M., Frentress, J., Kreps, T.A., Bridgham, S.D., Grossman, E., Jackson, C.J. 2003. Regional comparisons of watershed determinants of dissolved organic carbon in temperate lakes from the Upper Great Lakes region and selected regions globally. *Limnol. Oceanogr.* 48:2321-2334.