
1 **Soil-expanded seismic metamaterial with ultralow and wide bandgap**

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14 **Abstract:** The low-frequency wide-bandgap characteristics of the seismic metamaterial
15 can suppress the propagation of vibrations and reduce the risk of extreme loadings such
16 as earthquakes. The stringent requirement of lattice size extensively increasing with the
17 cost of forming seismic metamaterial using general engineering materials. We design
18 soil-expanded seismic metamaterial to reduce the scale restriction on artificial materials.
19 Two types of soil-expanded lattice are created, and the bandgap characteristics for the
20 lattice are obtained through the transfer matrix method. The propagation process for
21 finite periodic lattice is simulated by the finite difference method in the time domain. It
22 is found that the acceleration amplitudes in the wave propagation region are suppressed
23 by 90% for the seismic metamaterial with rubber components. The response spectra

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24 further indicate that seismic metamaterials can reduce seismic risk in targeted areas.

25 **Keywords:** Vibration attenuation, Seismic metamaterial, Bandgap, Periodic structures,
26 Soil-structure interaction, Seismic prevention.

27

28 **1. Introduction**

29 Earthquake, sudden and devastating natural disaster, has been well known for bringing
30 huge casualties and economic losses and is still a huge challenge for all mankind [1-3].
31 Various theories and techniques have been developed to promote structural seismic
32 resistance, but these measures can only protect the structure in which it is located, and
33 earthquake-indenergy-absorbing elements can cause considerable damage under large
34 earthquakes. These damages are often difficult to repair or even so difficult to replace
35 that the structure can no longer be used [4,5]. The presence of these problems makes
36 earthquakes still a huge threat to cities. Recently, the advent of phononic crystal
37 acoustic metamaterials has allowed us to directly control the propagation of elastic
38 waves. Seismic waves are elastic; they can also be controlled by phononic crystals,
39 meaning that region-scale phononic crystals have the potential to enable earthquake
40 protection for entire urban areas. Phononic crystal is an artificial acoustic metamaterial
41 composed of periodically distributed media [6] when region-scale metamaterials are
42 particularly designed for seismic this kind of metamaterial as seismic crystals [7] or
43 seismic metamaterials(SMs) [8].

44 Seismic metamaterials consist of a series of identical lattices that prevent elastic
45 waves in certain frequency ranges from propagating through the metamaterial. These
46 frequency ranges are often referred to as bandgaps and are the essence of seismic
47 metamaterials that can control the propagation of seismic waves. The realization of
48 bandgaps is based on periodically repeating lattices. The geometry, density, and

49 Young's modulus of the lattice control the location and width of the band gaps.
50 Specially designed low-frequency bandgap seismic metamaterials can prevent seismic
51 waves from propagating into urban areas or critical structures such as power plants and
52 hospitals.

53 Over the last decade, considerable literature has grown up around the topic of surface
54 wave attenuation by seismic metamaterials. A number of authors have considered the
55 formation of bandgaps by periodically burying artificial structures below the ground to
56 attenuate seismic surface waves. Pu [9, 10] analyzed surface wave attenuation by
57 periodically cylindrical concrete piles, calculated the effect of soil stratification on
58 bandgaps, and proposed a new method for identifying surface wave bandgaps. Zhang
59 [11], Miniaci [12] and Amanat [13] presented seismic metamaterials with different
60 cross-sections, they calculated the bandgaps and transmission properties of seismic
61 metamaterials by commercial finite element software. Considering the anisotropy of
62 the soil, Guo et al [14] combined theoretical derivation and numerical modelling to
63 calculate the bandgaps and transmission curve of rubber-steel piles. The authors
64 proposed a low-frequency bandgap seismic metamaterial in anisotropic soil. Using
65 laboratory experiments, Zeng [15] and Chen [16] verified the bandgaps of two different
66 seismic metamaterials through transmission curve. Different with mentioned above,
67 Brûlé [17] and Kacin [18] constructed periodic holes with different distribution shapes
68 in natural soil and verified the ability of seismic metamaterials to attenuate surface
69 waves by amplitude distribution. In addition, some studies have shown that the periodic
70 attachment of specially designed structures to the soil surface is also capable of forming
71 bandgaps. By commercial finite element software, the bandgaps and transmission curve
72 of built-up structural steel sections [19], H-fractal seismic metamaterial [20],
73 Minkowski-like fractal seismic metamaterial [21], T-shaped seismic metamaterial [22]

74 and Matryoshka-like seismic metamaterial [23] are well analyzed. Recently, Zeng [24]
75 pointed out that negative Poisson's ratio materials are more conducive to the formation
76 of low-frequency metamaterials. However, such studies remain narrow in focus dealing
77 only with surface waves attention.

78 Seismic waves are divided into surface waves and body waves. Surface waves
79 propagate only in the horizontal direction, while body waves can propagate in both the
80 horizontal and vertical directions. There is little published data on the attention of body
81 waves through seismic metamaterials [7,25]. As noted by Geng, one-dimensional
82 layered seismic metamaterials can effectively attenuate body waves propagating in any
83 direction [25]. They calculated the bandgap of a two-component layered seismic
84 metamaterial by a theoretical method and discussed the effect of material properties on
85 the width and location of the first-order bandgap. Geng argues that it is hard for two-
86 component seismic metamaterials to form low-frequency bandgaps when the artificial
87 material thickness is less than the seismic wavelength. However, they only calculated
88 the bandgaps for shear waves. For longitudinal waves in soil, which have longer
89 wavelengths, it is even more difficult to form bandgaps with the same lattice thickness.
90 In addition, little is known about the transmission properties of seismic metamaterials
91 with only finite dimensions in the aperiodic direction.

92 This paper discusses the case of attention of body waves by one-dimensional seismic
93 metamaterial. Two kind of one-dimensional seismic metamaterial to reduce the
94 thickness of artificial material are presented, both of them expanded lattice by soil.
95 There are two primary aims of this study: 1. To investigate the influence of soil
96 thickness to bandgaps. 2. To ascertain transmission properties of this kind of seismic
97 metamaterial. Firstly, the effect of soil thickness and artificial material thickness on the
98 bandgap of both longitudinal wave and shear wave by transfer matrix method is

99 analysed. Subsequently, the finite difference method in time domain (FDTD) is used to
100 simulate the transmission properties of one-dimensional seismic metamaterials. Finally,
101 the transmission properties of layered seismic metamaterials with finite size in
102 aperiodic direction are simulated by 2D FDTD considering the practical situation.

103 **2. Model and Bandgaps of soil-expanded lattice**

104 **2.1 The model of soil-expanded lattices**

105 Many researchers have utilized soil as a matrix to form two- or three-dimensional
106 seismic metamaterials. However, one-dimensional layered seismic metamaterials have
107 no difference between matrix and scatterer, and the lattice contains only artificial
108 material. As a result, such a lattice cannot have dimensions comparable to seismic
109 wavelengths. Although it can somewhat control transverse waves, it isn't easy to control
110 longitudinal waves. To make the lattice size of the one-dimensional layered seismic
111 metamaterial comparable to the seismic wavelength, we propose two soil-expanded
112 lattices, as shown in Fig. 1.

113 Fig. 1(a) shows a schematic diagram of the protection of a target building by one-
114 dimensional seismic metamaterial. As shown in Fig. 1(a), the seismic metamaterial is
115 placed in front of the target building. As shown in Fig. 1(b) and Fig. 1(c), two types of
116 soil-expanded lattices are illustrated. The red dashed box marked the lattice. The lattice
117 shown in Fig. 1(b) consists of concrete and soil (CS lattice), while the lattice shown in
118 Figure 1(c) consists of a concrete layer wrapped in rubber in front and behind and soil
119 (CRS lattice). Compared to Fig. 1(b), the CRS lattice in Fig. 1(c) contains rubber, which
120 is more conducive to forming band gaps through the local resonance mechanism.

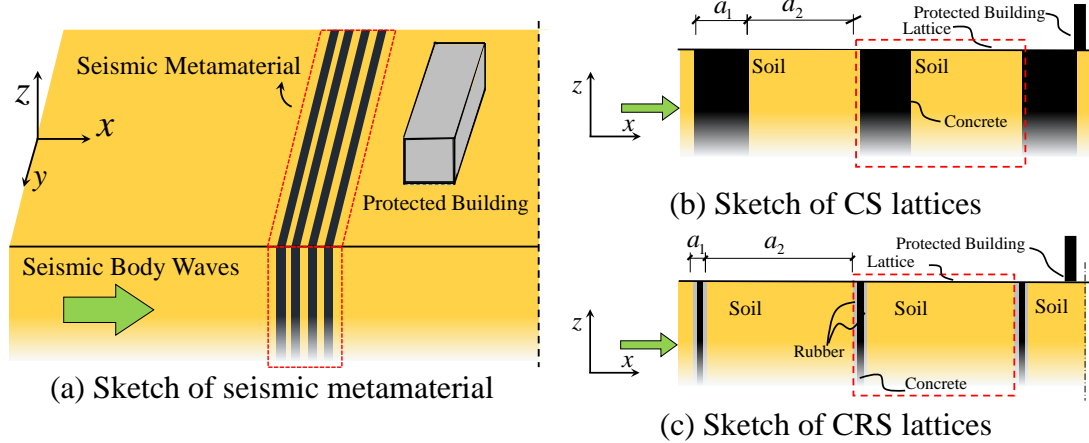


Fig. 1 Soil-expanded lattices

121

122

123 It is worth noting that both soil-expanded lattices make the soil an integral part of the
 124 lattice. So in both soil-expanded lattices, increasing the thickness of the soil increases
 125 the lattice size. As the seismic metamaterial is buried in the soil, controlling the soil
 126 thickness is easy to realize, that is, maintaining the burial distance between each
 127 concrete layer.

128 2.2 bandgaps of soil-expanded lattice

129 Eqn. 1 shows the equation for the propagation of an elastic wave in a one-
 130 dimensional metamaterial.

$$\frac{\partial^2 U(x,t)}{\partial t^2} = \frac{1}{C} \cdot \frac{\partial^2 U(x,t)}{\partial x^2} \quad (1)$$

131 where $U(x,t)$ is the displacement field in the computational domain, and C is the
 132 elastic wave velocity. In infinitely repeating one-dimensional layered metamaterials,
 133 the displacement fields at the ends of the lattice satisfy the Bloch period condition
 134 shown in Eqn. 2.

$$U(x+a) = e^{-ika} \cdot U(x) \quad (2)$$

135 where $U(x)$ is the displacement at the left end of the lattice, and $U(x+a)$ is the

136 displacement at the right end of the lattice. i is an imaginary number and k is the
137 wave number, a is the length of lattice. The eigenfrequency analysis of each wave
138 number k of the lattice shown in Fig. 1(b) and Fig. 1(c) using the transfer matrix method
139 enables the dispersion equation for the field expansion lattice to be obtained, as shown
140 in Eqn. 3.

$$2 \cos[k(a_1 + a_2)] = \cos\left(\frac{\omega a_1}{C_1}\right) \cos\left(\frac{\omega a_2}{C_2}\right) - \frac{C_1 \mu_2}{\mu_1 C_2} \sin\left(\frac{\omega a_1}{C_1}\right) \sin\left(\frac{\omega a_2}{C_2}\right) - \frac{C_2}{\omega \cdot \mu_2} \left(\frac{\mu_1 \cdot \omega \cdot \sin\left(\frac{\omega a_1}{C_1}\right) \sin\left(\frac{\omega a_2}{C_2}\right)}{C_1} - \frac{\omega \cdot \mu_2 \cos\left(\frac{\omega a_1}{C_1}\right) \cos\left(\frac{\omega a_2}{C_2}\right)}{C_2} \right) \quad (3-a)$$

$$2 \cos[k(a_1 + a_2)] = \cos\left(\frac{\omega a_1}{C_1}\right) \cos\left(\frac{\omega a_2}{C_2}\right) - \frac{(\lambda_2 + 2\mu_2) C_1}{(\lambda_1 + 2\mu_1) C_2} \sin\left(\frac{\omega a_1}{C_1}\right) \sin\left(\frac{\omega a_2}{C_2}\right) - \frac{C_2}{\omega \cdot (\lambda_2 + 2\mu_2)} \left(\frac{(\lambda_1 + 2\mu_1) \cdot \omega \cdot \sin\left(\frac{\omega a_1}{C_1}\right) \sin\left(\frac{\omega a_2}{C_2}\right)}{C_1} - \frac{\omega \cdot (\lambda_2 + 2\mu_2) \cos\left(\frac{\omega a_1}{C_1}\right) \cos\left(\frac{\omega a_2}{C_2}\right)}{C_2} \right) \quad (3-b)$$

$$2 \cos[k(a_1 + a_2)] = \frac{1}{C_2 \cdot C_3^2 \cdot C_4 \cdot \mu_2 \cdot \mu_3^2 \cdot \mu_4} \cdot \{ \cos^2\left(\frac{\omega a_1}{C_3}\right) \cdot [-\sin\left(\frac{\omega a_1}{C_2}\right) \sin\left(\frac{a_2 \omega}{C_4}\right) \cdot (C_3^2 \mu_4^2 + C_4^2 \mu_3^2) \cdot (C_2^2 \mu_3^2 + C_3^2 \mu_2^2) + 4 \cos\left(\frac{\omega a_1}{C_2}\right) \cos\left(a_2 \frac{\omega a_1}{C_4}\right) C_2 C_3^2 C_4 \cdot \mu_2 \cdot \mu_3^2 \cdot \mu_4] - 2 C_3 \mu_3 \sin\left(\frac{\omega a_1}{C_3}\right) \cos\left(\frac{\omega a_1}{C_3}\right) \cdot [C_2 \mu_2 \cos\left(\frac{\omega a_1}{C_2}\right) \sin\left(a_2 \frac{\omega a_1}{C_4}\right) (C_3^2 \mu_4^2 + C_4^2 \mu_3^2) + C_4 \mu_4 \sin\left(\frac{\omega a_1}{C_2}\right) \cos\left(a_2 \frac{\omega a_1}{C_4}\right) (C_2^2 \mu_3^2 + C_3^2 \mu_2^2)] + \sin\left(\frac{\omega a_1}{C_2}\right) \sin\left(\frac{\omega a_2}{C_4}\right) (C_2^2 C_4 \mu_3^4 + C_3^4 \mu_2^2 \mu_4^2) - 2 \cos\left(\frac{\omega a_1}{C_2}\right) \cos\left(\frac{\omega a_2}{C_4}\right) C_2 \cdot C_3^2 \cdot C_4 \cdot \mu_2 \cdot \mu_3^2 \cdot \mu_4 \} \quad (3-c)$$

$$2 \cos[k(a_1 + a_2)] = \frac{1}{C_2 \cdot C_3^2 \cdot C_4 \cdot (\lambda_2 + 2\mu_2) \cdot (\lambda_3 + 2\mu_3)^2 \cdot (\lambda_4 + 2\mu_4)} \cdot \{ \cos^2\left(\frac{\omega a_1}{C_3}\right) \cdot [-\sin\left(\frac{\omega a_1}{C_2}\right) \sin\left(\frac{a_2 \omega}{C_4}\right) \cdot (C_3^2 (\lambda_4 + 2\mu_4)^2 + C_4^2 (\lambda_3 + 2\mu_3)^2) \cdot (C_2^2 (\lambda_3 + 2\mu_3)^2 + C_3^2 (\lambda_2 + 2\mu_2)^2) + 4 \cos\left(\frac{\omega a_1}{C_2}\right) \cos\left(a_2 \frac{\omega a_1}{C_4}\right) C_2 C_3^2 C_4 \cdot (\lambda_2 + 2\mu_2) \cdot (\lambda_3 + 2\mu_3)^2 \cdot (\lambda_4 + 2\mu_4)] - 2 C_3 (\lambda_3 + 2\mu_3) \sin\left(\frac{\omega a_1}{C_3}\right) \cos\left(\frac{\omega a_1}{C_3}\right) \cdot [C_2 (\lambda_2 + 2\mu_2) \cos\left(\frac{\omega a_1}{C_2}\right) \sin\left(a_2 \frac{\omega a_1}{C_4}\right) (C_3^2 (\lambda_4 + 2\mu_4)^2 + C_4^2 (\lambda_3 + 2\mu_3)^2) + C_4 (\lambda_4 + 2\mu_4)^2 \sin\left(\frac{\omega a_1}{C_2}\right) \cos\left(a_2 \frac{\omega a_1}{C_4}\right) (C_2^2 (\lambda_3 + 2\mu_3)^2 + C_3^2 (\lambda_2 + 2\mu_2)^2)] + \sin\left(\frac{\omega a_1}{C_2}\right) \sin\left(\frac{\omega a_2}{C_4}\right) (C_2^2 C_4 (\lambda_3 + 2\mu_3)^4 + C_3^4 (\lambda_2 + 2\mu_2)^2 (\lambda_3 + 2\mu_3)^2) - 2 \cos\left(\frac{\omega a_1}{C_2}\right) \cos\left(\frac{\omega a_2}{C_4}\right) C_2 \cdot C_3^2 \cdot C_4 \cdot (\lambda_2 + 2\mu_2) \cdot (\lambda_3 + 2\mu_3)^2 \cdot (\lambda_4 + 2\mu_4) \} \quad (3-d)$$

141 where a_1 is the thickness of the concrete layer of the CS lattice in Eqn.(3-a) and
142 Eqn. (3-b). a_1 is the thickness of the concrete layer of the CRS lattice including the
143 outer rubber layer in Eqn. (3-c) and Eqn. (3-d). And a_2 is the thickness of the soil in
144 Eqn. (3-a), Eqn. (3-b), Eqn. (3-c) and Eqn. (3-d). a_1, a_2 are geometrical parameters,
145 and they are all illustrated in Fig. 1. In Eqn. 3, μ_n and λ_n are the Lamé constants,
146 and C_n is wave velocity of the Nth layer. k is wave number and ω is angular
147 frequency. Eqn. (3-a) and Eqn. (3-b) describes the dispersion equation for the CS lattice,

148 Eqn. (3-c) and Eqn. (3-d) describes the dispersion equation for the CRS lattice. The
149 detailed analysis of Bloch's boundary conditions and the derivation of Eqn. 3 can be
150 found in Supplementary Material S1 and literature [26-28].

151 When the wave number k in Eqn. 3 is specified, a number of eigenfrequency f that
152 satisfy the equation can be calculated. Since the lattice is periodically repeated, the
153 dispersion curve is obtained by calculating only the wave number k within the first
154 Brillouin zone($k \in \left(-\frac{\pi}{2a}, \frac{\pi}{2a}\right)$), where a is the length of the lattice. When there is
155 no corresponding wave number k in a section of frequency, these frequency sections
156 are called bandgaps. As we can see, there is more than one bandgap. The upper and
157 lower boundaries of bandgaps determine the location and width of the bandgaps. If an
158 elastic wave has only frequencies within the bandgaps, such an elastic wave cannot
159 propagate through seismic metamaterials. Seismic waves are ultra-low frequency
160 elastic waves, the main frequency component of which is 0-15 Hz. Therefore, seismic
161 metamaterials should have a bandgap that is sufficiently low frequency and as wide as
162 possible. The bandgap with the smallest lower bound mark is bandgap 1st, and as the
163 frequency rises, the other bandgaps are 2nd, 3rd, and so on. By taking $k = \pi/(2a)$ into
164 Eqn. 3, we can obtain the upper and lower bound frequencies for the odd bandgaps.
165 Bringing $k = 0$ into Eqn.3 gives the upper and lower bounds of the bandgap for order
166 bandgaps.

167 Since the lower bound of bandgap number one has the lowest frequency, we take
168 $k = \pi/2 \cdot a$ into account Eqn. 3 and calculate odd bandgaps. Note that the dispersion
169 equations for the CS and CRS lattices have the artificial material layer thickness
170 parameter and the soil thickness parameter, which can influence the upper and lower
171 bound of the band gap. Therefore, the frequency f is calculated using the artificial

172 material layer thickness parameter a_1 and the soil thickness parameter a_2 as
 173 variables, as shown in Fig. 2. Fig. 2(a) shows the distribution of the upper and lower
 174 boundaries of the odd-numbered bandgap for the CS lattice in both modes. The vertical
 175 coordinate is the frequency, the axis a_1 represents the concrete thickness, and the axis
 176 a_2 represents the soil layer thickness. Fig. 2(b) shows the distribution of the upper and
 177 lower boundaries of the odd-numbered bandgap of the CRS lattice for both modes. The
 178 vertical coordinate is the frequency f , the axis a_1 represents the total thickness of the
 179 rubber and concrete layers, and the axis a_2 represents the soil thickness. The material
 180 data are shown in Tab. 1, where V_p and V_s is the wave velocity of the longitudinal
 181 wave and shear wave. ρ is the density of the material, μ and λ are the Lamé
 182 coefficient of the material.

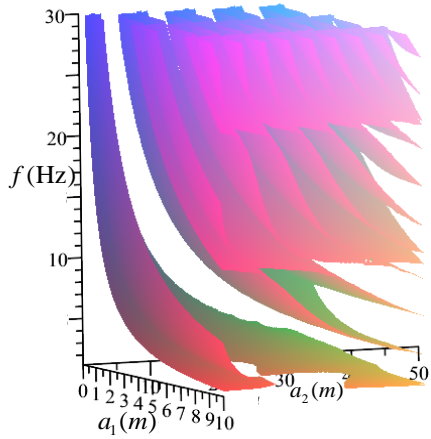
183 Table 1 Material property in CS and CRS lattice

| Material | V_p (m/s) | V_s (m/s) | ρ (g/cm ³) | μ | λ |
|--------------|----------------|----------------|--------------------------------|--------------------|-----------------------|
| C30 Concrete | 3800 | 2820 | 2.36 | 9.4×10^9 | 1.82×10^{10} |
| Soil | 1500 | 210 | 1.93 | 8.49×10^7 | 2.12×10^9 |
| Rubber | 571 | 45 | 1.3 | 2.67×10^6 | 4.18×10^8 |

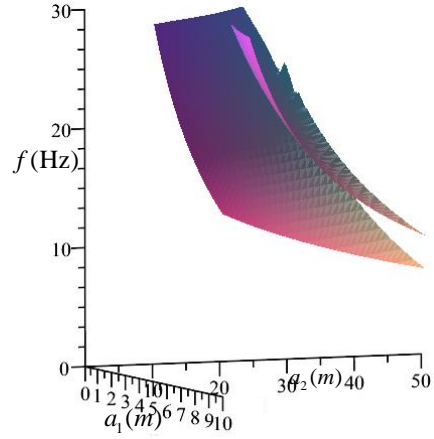
184
 185 Fig. 2 shows that the upper and lower boundaries of the odd-numbered bandgaps
 186 corresponding to different layer thicknesses. Both lattices can reduce the lower
 187 boundary of the bandgap by adding concrete and soil thickness. This may be due to the
 188 fact that increasing the concrete layer thickness is equivalent to increasing the resonator
 189 mass of the local resonance mechanism and reducing the band gap position. The CRS
 190 lattice is covered with softer rubber than soil, and the same thickness of concrete makes

191 it easier for the local resonance mechanism to form, requiring less concrete thickness.
192 However, the increased thickness of the artificial material layer increases the cost of the
193 seismic metamaterial. Fig. 2(b)(d) shows the results of the analysis of the longitudinal
194 wave mode, where we consider the value of 30m for the CS lattice and 10m for the
195 CRS lattice.

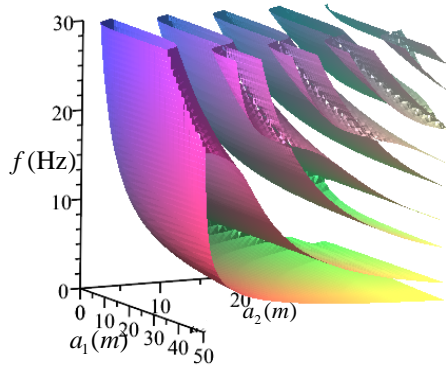
196 Another trend in the bandgap in Fig. 2 is that the upper and lower boundary spacing
197 of bandgap number one becomes smaller as the soil thickness increases, which means
198 that the bandgap width decreases. There are two reasons for this phenomenon: on the
199 one hand, increasing soil thickness allows for a significant increase in lattice size, which
200 facilitates the acquisition of low-frequency band gaps. On the other hand, when the
201 thickness of the soil increases infinitely, the seismic metamaterial will converge
202 infinitely to the natural soil. However, natural soils do not have any bandgap, which
203 makes the distance between the upper and lower boundaries of the bandgap decrease
204 with increasing soil thickness.



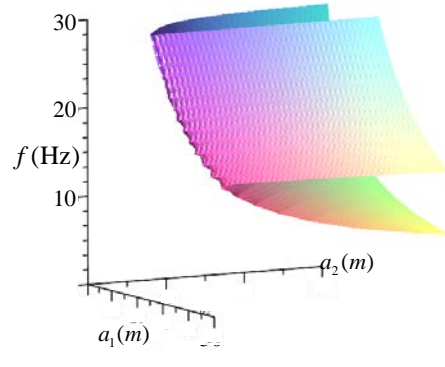
(a) CS lattice bandgaps in shear wave



(b) CS lattice bandgaps in longitudinal wave



(c) CRS lattice bandgaps in shear wave



(d) CRS lattice bandgaps in longitudinal wave

Fig. 2 3D view of the bandgaps

205

206

207 To find a reasonable soil thickness, we take the given a_1 and $k = \pi/2 \cdot a$ into Eqn. 3,

208 get the upper and lower bounds for the odd-numbered and even-numbered bandgaps,

209 respectively, as shown in Fig. 3.

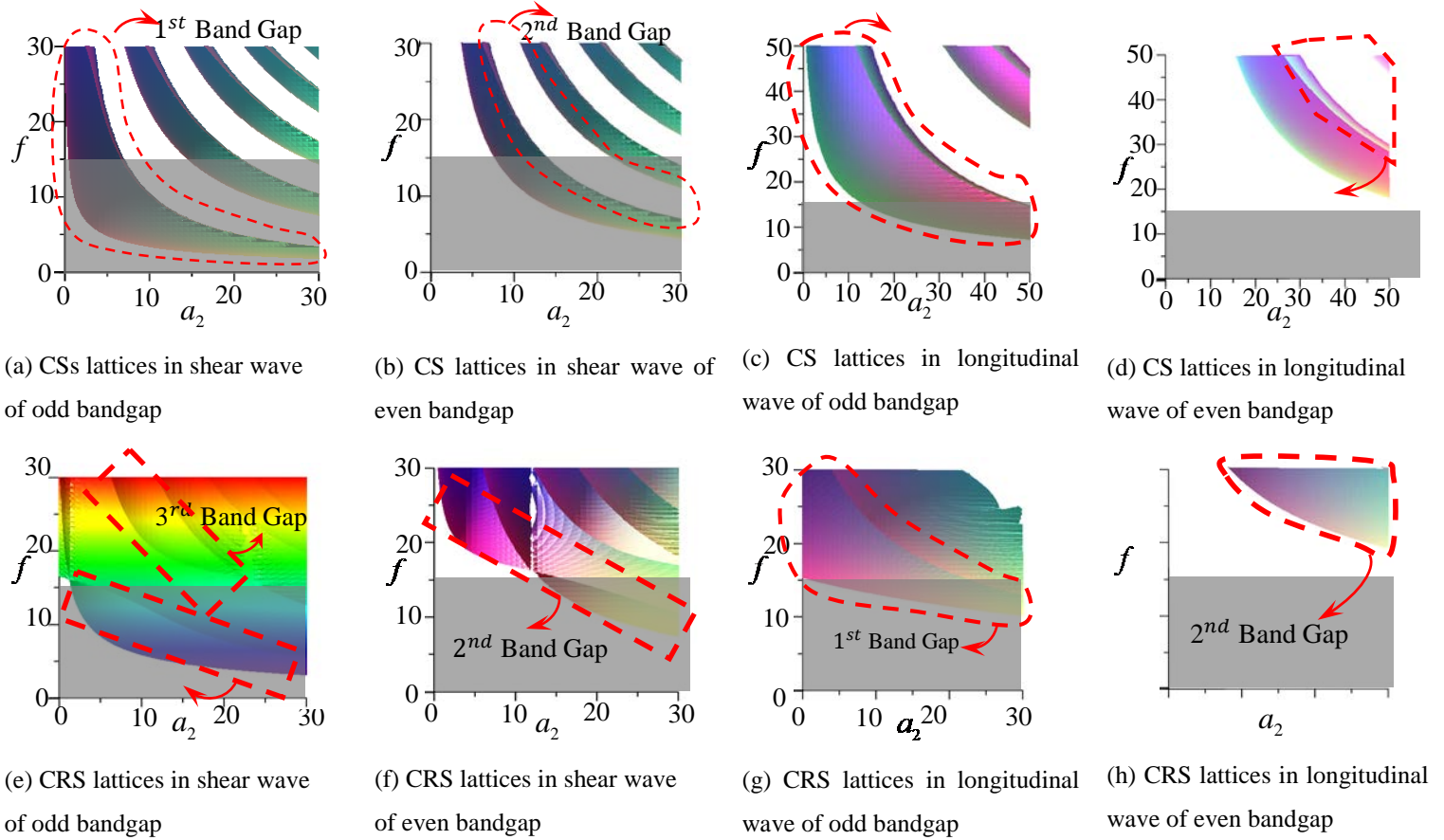


Fig. 3 Variation of bandgap with the thickness of soil

210
 211 The upper and lower boundaries of the bandgap are circled with a red dash line in
 212 Fig. 3, while the grey rectangle represents the dominant frequency of the seismic wave
 213 between 0 and 15 Hz. When the red dashed box intersects the grey part of the diagram,
 214 the soil thickness is such that a valid bandgap can be formed. For example, Fig. 3(a)(b)
 215 the complete bandgap of CS lattice in shear wave model. The 1st bandgap are circled
 216 with red dash line, and has the intersect with grey part when a_2 is less than 1m.
 217 However, Fig. 3(c)(d) shows that the 1st bandgap intersect with grey part when a_2
 218 more than 10m. For more valid bandgaps in longitudinal wave, the value of a_2 could
 219 be 50m. As the value of a_2 is more than 50m, the intersection of red wireframe and
 220 gray part gets narrower. So the value of a_2 should be 50m. Likely, we can get the soil
 221 thickness of CRS lattice. Tab. 2 shows the size of two lattices.

222

Table2 Size of two lattices

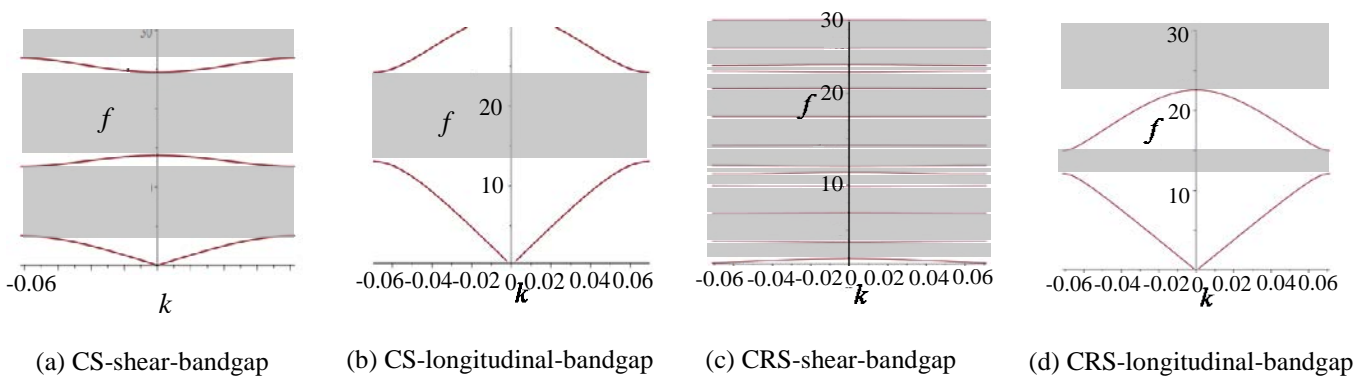
| Lattice | a_1 (m) | a_2 (m) | a (m) |
|---------|-----------|-----------|---------|
| CS | 30 | 50 | 80 |
| CRS | 10 | 30 | 40 |

223

224 Taking the thickness shown in Tab. 2 to Eqn.3, the Fig. 4 shows the dispersion curves

225 for the CS and CRS lattice. The grey part in Fig. 4 is the bandgap. The CS and CRS

226 lattices can form bandgaps within 0-15Hz in both modes.



227

Fig. 4 Bandgaps for various seismic metamaterial

228 3. Numerical simulation of soil-expanded lattice

229 The time-domain finite-difference method of acceleration loading is used to simulate

230 the propagation of seismic waves through seismic metamaterials. This numerical

231 method improves the standard time-domain finite-difference method and allows for

232 convenient statistics on the computational domain's maximum acceleration and

233 maximum displacement. The numerical calculation methods are described in

234 Supplementary Material S2 and can also be found in the literature [29-31].

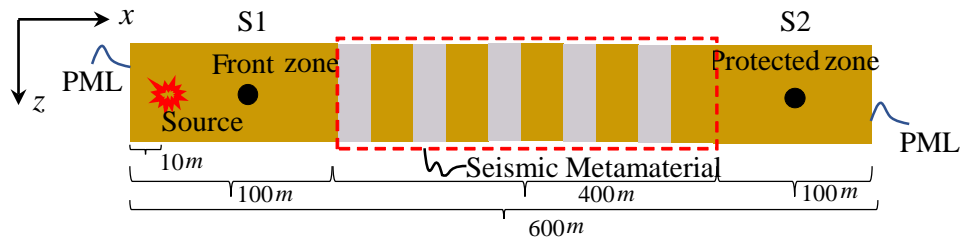
235 When the seismic metamaterial in Fig. 1(a) has infinite dimensions in the y-direction

236 and seismic waves cannot pass the seismic metamaterial, the computational domain can

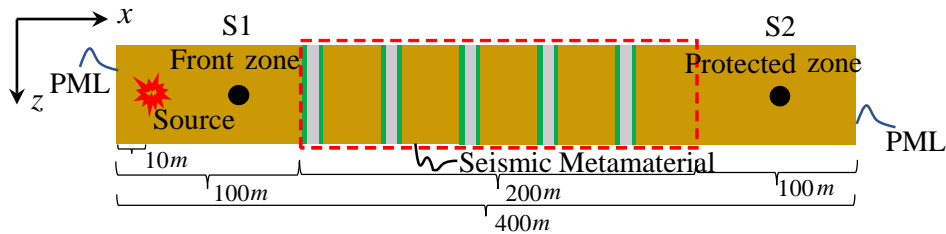
237 be reduced to one dimension. Fig. 5 shows the one-dimensional computational domain.

238 The PML absorption boundary is set at both ends of the computational domain, and the

239 source is 50 m from the left-end boundary. The time history of point S1 and point S2
 240 and the distribution of maximum acceleration and maximum displacement are recorded
 241 during the computation.

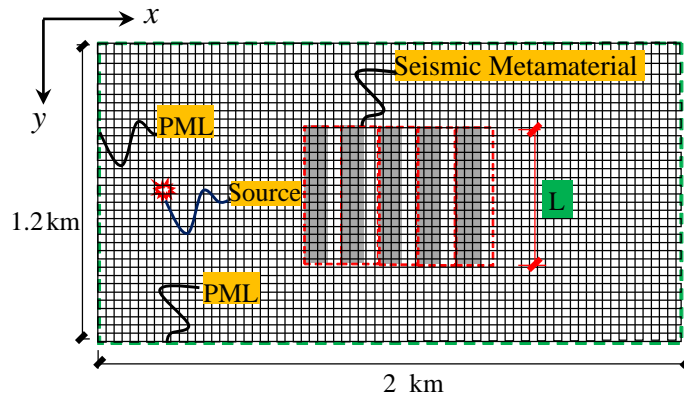


(a) seismic metamaterial by CS lattice



(b) seismic metamaterial by CRS lattice

242 Fig. 5 Seismic metamaterial with infinite length in y direction



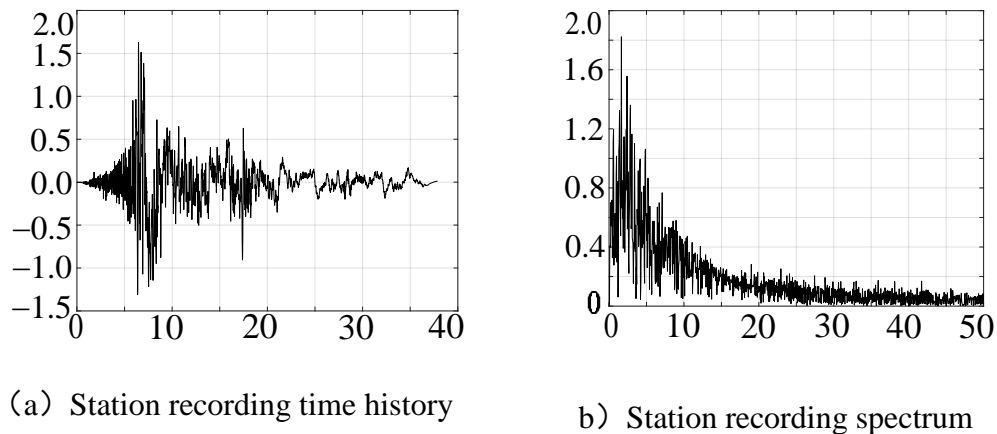
243
 244 Fig. 6 Seismic metamaterial with finite length in y direction

245 When the seismic metamaterial in Fig. 1(a) has only a finite size in the y-direction,
 246 there is a possibility that the seismic wave will go around to the back of the seismic
 247 metamaterial during propagation, as shown in Fig. 6. The computational domain is a
 248 square area of 1.2km×2.0km. The four sides of the square are set with PML absorption
 249 boundaries. The length of seismic metamaterial in the y-direction is 600m.

250 **3.1 1D numerical analysis**

251 We used seismic wave data from the PEER website, El-Centro Array #9, from the
252 earthquake of Imperial Valley-05. this seismic wave has a rich low frequency
253 component. Fig. 7 shows the time history and spectrum of the seismic wave.

254



255 Fig. 7 The source schedule of seismic wave

256 Considering shear waves, the seismic waves were analyzed using the finite difference
257 method in the time domain through the one-dimensional computational domain shown
258 in Fig. 5, and the maximum displacement and acceleration distributions within all
259 computational domains were counted, as shown in Fig. 8.

260 Fig. 8 shows the distribution of the maximum acceleration and displacement of the
261 seismic wave through the 1D computational domain. The "Uniform" and "SMs" labels
262 mark the natural soil and the seismic metamaterial, respectively, as shown by the blue
263 and red curves in Fig. 8. "F zone" and "P zone" mark the front and back of the seismic
264 metamaterial. Fig. 8(a)(b) shows the CS seismic metamaterial compared to natural soil,
265 and Fig. 8(c)(d) shows the CRS seismic metamaterial compared to natural soil.

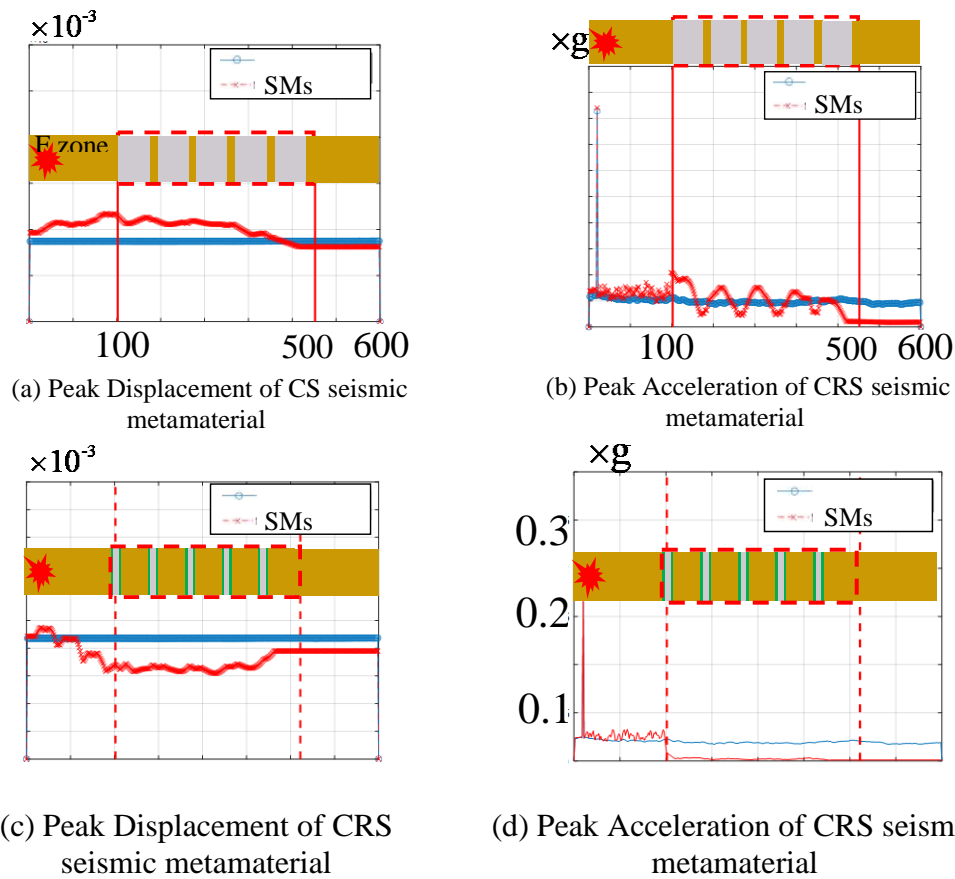


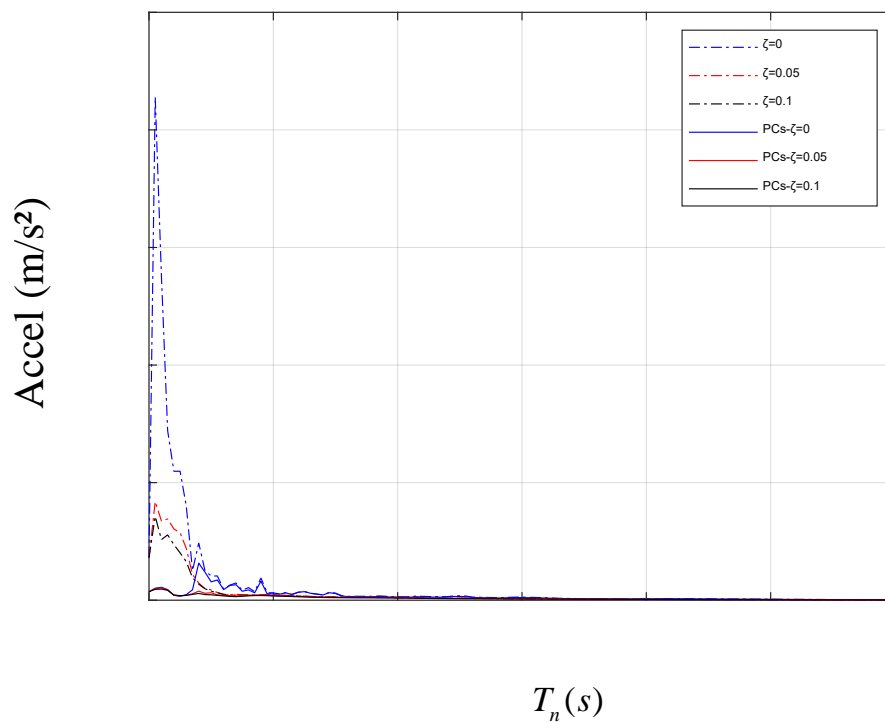
Fig. 8 Max displacement and acceleration

266

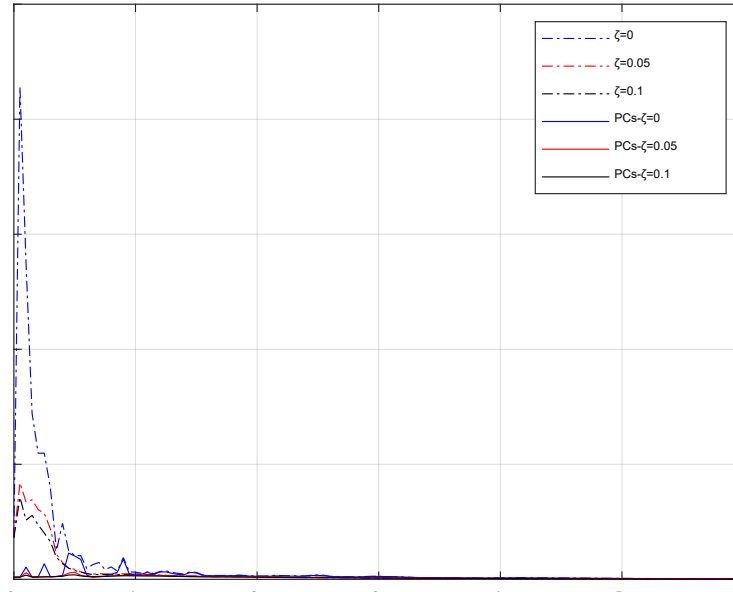
267 Firstly, both lattices effectively attenuate seismic waves, as seen from the red curve
 268 in Fig. 8, which is lower in the 'P zone' than the blue curve. It can be seen that the red
 269 curve in Fig. 8(b)(d) is significantly lower than the blue curve. Both lattices are better
 270 at attenuating acceleration than displacement. As acceleration is an important indicator
 271 of earthquake engineering, it can directly respond to the magnitude of the seismic forces
 272 in a structure. Therefore, installing seismic metamaterials in front of the protected
 273 building can effectively reduce the seismic forces on the structure.

274 Secondly, the attenuation process of seismic waves differs significantly between the
 275 two lattices. The maximum acceleration within the CS seismic metamaterial in Fig. 8(b)
 276 produces a peak at the concrete layer, and the peak decreases as the lattice increases.
 277 The peak is probably due to the high shear force where the concrete is located, and the
 278 acceleration required to balance the shear force increases accordingly, creating a peak.

279 In contrast, the maximum acceleration within the CRS seismic metamaterial decays
280 directly to near zero in the first lattice, and there is no peak in the acceleration within
281 the seismic metamaterial. The wave impedance of the concrete layer is much higher
282 than that of the rubber layer, resulting in a direct reflection of the elastic wave so that
283 the maximum acceleration in the CRS seismic metamaterial is more evenly distributed
284 and very close to 0. The red curve in the "F zone" section of Fig. 8(d) is higher than the
285 blue curve, illustrating the reflection phenomenon.



(a) CR lattice acceleration response spectrum



$$T_n (s)$$

(b) CRS lattice acceleration response spectrum

286

Fig. 9 response spectrum of seismic metamaterial

287

The dashed line in Fig. 9 is the response spectrum for the time course recorded at

288

monitoring point S1, and the solid line is the response spectrum for the time course

289

recorded at monitoring point S2. The different colored curves represent the additional

290

damping of the response spectra. It can be seen that the peak of the response spectrum

291

at point S2 is significantly lower than the peak of the response spectrum at point S1,

292

irrespective of the damping level. Buildings in the 'P zone' are at substantially lower

293

seismic risk than those in the 'F zone.' Both seismic metamaterials significantly reduce

294

the seismic hazard in the protected area.

295

3.2 Lattice Equivalence Phenomenon

296

In contrast to higher-dimensional seismic metamaterial, the scatterer and matrix

297

within the lattice of a 1D seismic metamaterial do not strictly distinguish between each

298

other, which are merely different layers. When we introduce such soil-expanded lattice,

299 the lattice just is a finite number of layers inserted periodically in the soil. The lattice
 300 positions indicated by the red boxes in Fig. 1 are only one case, and Fig. 10 shows a
 301 fully equivalent lattice schematic to the red boxes.

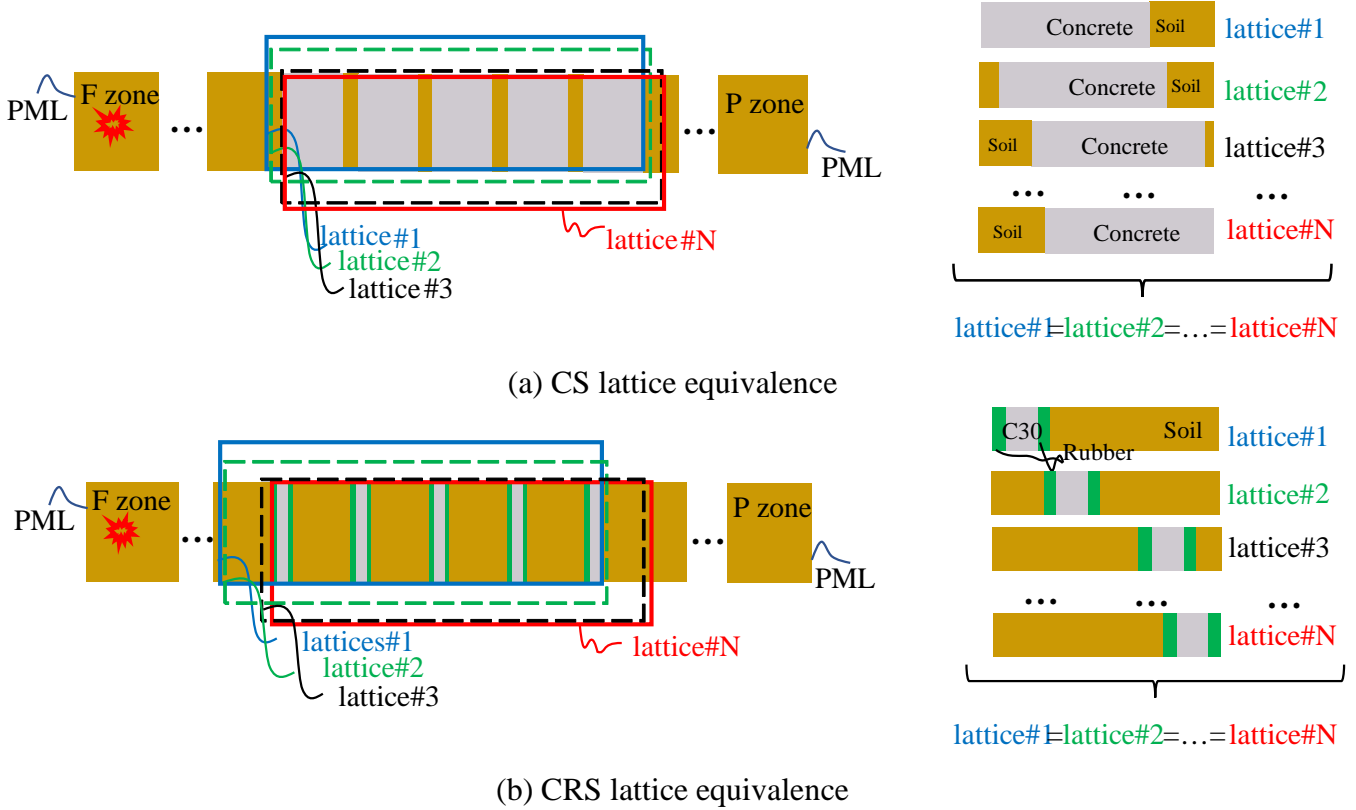


Fig. 10 Phenomenon of lattice equivalence

302 In Fig. 10(a) and Fig. 10(b), the red wireframe and other frames are equivalent
 303 lattices, and the difference between them is mainly the difference in lattice form. In the
 304 figure's red wireframe and other color wireframes, any wireframes of equal length can
 305 be completely equivalent lattices. This lattice equivalence makes it impossible for
 306 elastic waves to identify a complete CS or CRS lattice. At both ends of the seismic
 307 metamaterial, the part that acts as a soil infill will still have significant fluctuations and
 308 will not get shock absorption because it is a soil infill.

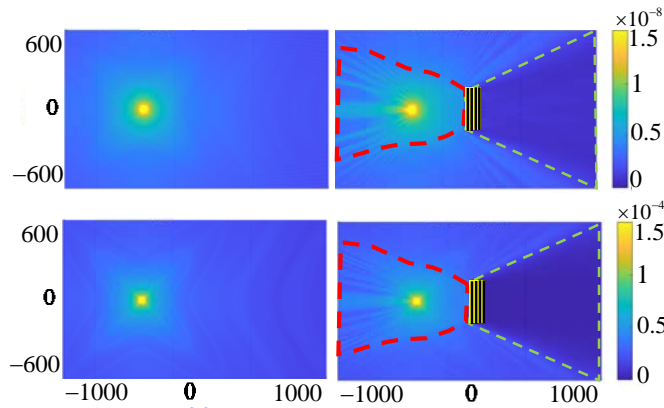
310 3.3 2D numerical analysis

311 For the CRS lattice, when the seismic metamaterial has finite dimensions in the y-

312 direction, we use a two-dimensional time domain finite difference method to simulate
 313 the propagation of elastic waves in the seismic metamaterial. We choose a harmonic
 314 wave located within the bandgaps as the seismic source, and Eqn. 5 shows the time
 315 expression.

$$a(t) = 0.01 \times [0.5 \cdot \sin(2\pi \cdot 8t) + 0.3 \cdot \sin(2\pi \cdot 18t) + 0.2 \cdot \sin(2\pi \cdot 23t)] \quad (5)$$

316 Fig. 11 gives the results of numerical calculations of seismic metamaterials with a
 317 finite length in the y-direction. On the left are the results for natural soil, labeled
 318 'uniform.' The seismic metamaterials simulation results are on the right, labeled "SMs."
 319 The red boxes mark the areas of enhanced vibration due to reflections, and the green
 320 boxes mark the areas of seismic protection provided by the seismic metamaterials.
 321 When the length of the seismic metamaterial is limited in the y-direction, the area of
 322 protection provided by the seismic metamaterial is a trapezoid.



323
 324 Fig. 11 The attenuation and amplification zone of finite length without rotation

325 4. Conclusion

326 In this paper, two kinds of soil-expanded seismic metamaterial models were proposed.
 327 Then, the transfer matrix method investigated the effect of layer thickness on the
 328 bandgaps of the CS lattice and CRS lattice. Subsequently, the propagation of elastic
 329 waves through the seismic metamaterial is simulated using the time domain finite

330 difference method. The way in which the seismic metamaterials work is discussed by
331 comparing the maximum acceleration and maximum displacement distributions of the
332 seismic metamaterials and the natural soil in the one-dimensional computational
333 domain. By analyzing the response spectrum in front of the seismic metamaterial and
334 the response spectrum behind the seismic metamaterial, it is demonstrated that the two
335 seismic metamaterials can enhance the seismic safety of the target building. The range
336 of seismic safety zones formed by the seismic metamaterials is found by comparing the
337 maximum acceleration and maximum displacement distributions of the seismic
338 metamaterials with the natural soil in the two-dimensional calculation domain.

339 (1) Both CRS lattice and CS lattice could form lower bandgaps by soil expanded
340 mechanism. While the thickness of the soil layer can lower the position of the
341 bandgap, it can also lead to a reduction in the width of the band gap. A reasonable
342 soil layer thickness should take into account an appropriate reduction of the band
343 gap position and not an excessive narrowing of the bandgap width.

344 (2) The CS lattice attenuates seismic waves slowly and more strongly with a higher
345 number of lattices, while the CRS lattice reflects them directly, independent of
346 the number of lattices.

347 (3) The distribution of the maximum displacement and the distribution of the
348 maximum acceleration indicates that the wave cannot distinguish lattice position
349 due to the lattice equivalence phenomenon.

350 (4) Compared with natural soil, the peak value of the response spectrum in the
351 protected area after the addition of lattice decreased, among which the peak value
352 of the response spectrum in the s-wave mode decreased by 90 %.

353 (5) Seismic metamaterials of finite length can form trapezoidal seismic safety zones.
354 The vibration-strengthening zone due to reflections is also trapezoidal.

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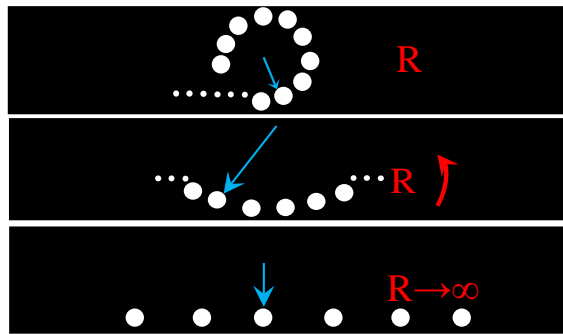
362

363 **Supplement Material S1**

364 The one-dimensional fluctuation equation is as in Eqn. S1-1.

$$\frac{\partial^2 U}{\partial t^2} = \frac{1}{C} \cdot \frac{\partial^2 U}{\partial x^2} \quad (S1-1)$$

365 where C is the wave velocity, which can be determined as longitudinal wave velocity
 366 by $c_p = \sqrt{(2\mu + \lambda)/\rho}$ or shear wave velocity $c_s = \sqrt{\mu/\rho}$, corresponding to the shear
 367 wave propagation equation and the longitudinal wave propagation equation. Assuming
 368 that the model satisfies infinite repetition in the period domain, the computational
 369 model satisfies the Bonn-Kamen boundary condition, as shown in Fig. 1.



370

Fig. S1 Bonn-Kamen Boundary

371 One-dimensional seismic metamaterials possess the Bonn-Kamen boundary
 372 condition shown in Fig. S1. It is assumed that the blue arrows represent the field
 373 functions $\Phi(x)$ at the coordinates x , and the white origin represents the position of the
 374 lattice. When the number of lattices is limited, the crystal is assumed to take the form

375 of a ring of radius R. Assuming that there are n_1 lattices in the ring, the field function
 376 after one lattice satisfies equation S1-2.

$$|\Phi(x)|^2 = |\Phi(x+a)|^2 \quad (\text{S1-2})$$

377 where a is the lattice length. Eqn. S1-2 is equivalent to Eqn. S1-3.

$$\Phi(x) = q_1 \cdot \Phi(x+a) \quad (\text{S1-3})$$

378 After n_1 lattice, the field function returns to the coordinate x and satisfies the Eqn.
 379 S1-4

$$\Phi(x) = q_1^{n_1} \cdot \Phi(x) \quad (\text{S1-4})$$

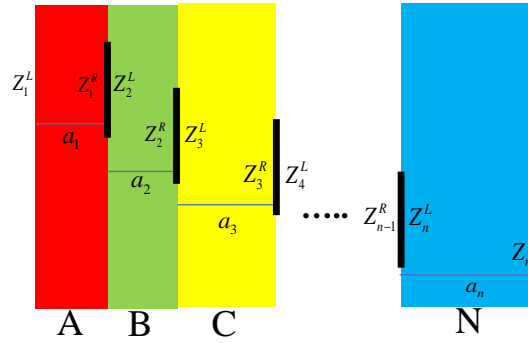
380 Obviously, $q_1^{n_1} = 1$, and $q_1 = \sqrt[n_1]{1}$, and q_1 can be replaced by Eqn. S1-5.

$$q_1 = e^{i \cdot 2\pi \cdot \frac{n}{n_1}} \quad (\text{S1-5})$$

381 As the radius R of the ring crystal increases, the number of lattices n_1 increases
 382 accordingly, as shown in Figure S1-(b). As R increases to infinity, the crystal can be
 383 considered to be in the form of a straight chain, and for each lattice passed, the form of
 384 the field function will satisfy Eqn. S1-6.

$$\Phi(x) = e^{i \cdot \frac{2\pi n}{N} \cdot \frac{(x+a)}{a}} \Phi(x+a) = e^{i \cdot k \cdot a} \Phi(x+a) \quad (\text{S1-6})$$

385 where, $k = \frac{2\pi \cdot n}{N}$, N is the total number of lattices. Eqn. S1-6 is the Bonn-Kamen
 386 boundary condition.



388
389

Fig. S1 Transfer matrix calculation process

390 Fig. S2 shows a one-dimensional lattice, with A, B, C...N being the different
391 material layers. Two adjacent layers should satisfy displacement continuity and shear
392 stress continuity. The nth layer displacement and shear stress are shown in Eqn. S2-1
393 and Eqn. S2-2 respectively.

$$u_n(x_n) = A_n \sin\left(\frac{\omega}{C_{sn}} x_n\right) + B_n \cos\left(\frac{\omega}{C_{sn}} x_n\right) \quad (\text{S2-1})$$

$$\tau_n(x_n) = \frac{\mu_n \omega}{C_{sn}} \left[A_n \cos\left(\frac{\omega}{C_{sn}} x_n\right) - B_n \sin\left(\frac{\omega}{C_{sn}} x_n\right) \right] \quad (\text{S2-2})$$

394 where x_n is the local coordinate of the nth level. A_n and B_n is the coefficient
395 to be determined. C_{sn} is the shear wave velocity. Let $Z_n(x_n)$ be the state function, as
396 shown in Eqn. S2-3.

$$Z_n(x_n) = H_n(x_n) \cdot \Psi(x_n) \quad (\text{S2-3})$$

397 where $H_n(x_n)$ is the state quantity, as shown in Eqn. S2-4. $\Psi(x_n)$ is the
398 coefficient like vector to be determined, as shown in Eqn. S2-5.

$$H_n(x_n) = \begin{bmatrix} \sin\left(\frac{\omega}{C_{sn}}x_n\right) & \cos\left(\frac{\omega}{C_{sn}}x_n\right) \\ \frac{\mu_n\omega}{C_{sn}}\cos\left(\frac{\omega}{C_{sn}}x_n\right) & -\frac{\mu_n\omega}{C_{sn}}\sin\left(\frac{\omega}{C_{sn}}x_n\right) \end{bmatrix} \quad (\text{S2-4})$$

$$\Psi = [A_n \quad B_n]^T \quad (\text{S2-5})$$

399 The state function on the left side of the nth level is shown in Equation S2-6. The
400 state function on the right-hand side of the nth level is shown in Eqn. S2-7.

$$Z_n^L = H_n(0) \cdot \Psi_n \quad (\text{S2-6})$$

$$Z_n^R = H_n(a_n) \cdot \Psi_n \quad (\text{S2-7})$$

401 The relationship between Z_n^L and Z_n^R is shown in Eqn. S2-8.

$$Z_n^R = H_n(a_n)[H_n(0)]^{-1} \cdot Z_n^L \quad (\text{S2-8})$$

402 Due to the continuous displacement between two adjacent layers, there is

$$Z_{n+1}^L = Z_n^R \quad (\text{S2-8})$$

403 Thus, the state function from the leftmost to the rightmost part of the lattice satisfies
404 Eqn. S2-9.

$$Z_n^R = \left(H_n(a_n)[H_n(0)]^{-1} \cdot H_{n-1}(a_{n-1})[H_{n-1}(0)]^{-1} \dots H_1(a_1)[H_1(0)]^{-1} \right) Z_1^L \quad (\text{S2-9})$$

405 And the state function is one that satisfies the Bonn-Kamen boundary condition, as
406 shown in Eqn. S2-10.

$$Z_n^R = e^{ika} \cdot Z_1^L \quad (\text{S2-10})$$

407 The characteristic equation for lattice is shown in Eqn. S2-11.

$$0 = \left(\mathbf{H}_n(a_n)[\mathbf{H}_n(0)]^{-1} \cdot \dots \mathbf{H}_1(a_1)[\mathbf{H}_1(0)]^{-1} - e^{ika} \mathbf{I} \right) Z_1^L \quad (\text{S2-10})$$

408 By solving this characteristic equation, the corresponding dispersion curve of the
409 lattice can be obtained.

410 **Supplement Material S3**

411 The one-dimensional fluctuation equation is shown in Eqn. S3-1. The two-
412 dimensional z-mode fluctuation equation is shown in Eqn. S3-2.

$$\frac{\partial^2 U(x,t)}{\partial t^2} = \frac{1}{C} \cdot \frac{\partial^2 U(x,t)}{\partial x^2} \quad (\text{S3-1})$$

$$\rho \frac{\partial^2 U_z(x,y,t)}{\partial t^2} = \mu \frac{\partial^2 U_z(x,y,t)}{\partial x^2} + \mu \frac{\partial^2 U_z(x,y,t)}{\partial y^2} \quad (\text{S3-2})$$

413 Where, C_s and C_p are the transverse and longitudinal wave velocities of the
414 material, ρ and μ and λ are the density and Lamé coefficients of the material,
415 respectively. $U(x,t)$ and $U(x,y,t)$ are the displacement fields in the 1D elastic
416 wave equation and the 2D elastic wave equation.

417 Using the Taylor expansion to approximate the partial derivative term, the central
418 difference form partial derivative term is shown in Eqn. S3-3 and Eqn. S3-4.

$$D_{xx}U(x,t) = \frac{\partial^2 U(x,t)}{\partial x^2} = \frac{U(x+\Delta x,t) + U(x-\Delta x,t) - 2 \cdot U(x,t)}{\Delta x^2} \quad (\text{S3-3})$$

$$D_{tt}U(x,t) = \frac{\partial^2 U(x,t)}{\partial t^2} = \frac{U(x,t+\Delta t) + U(x,t-\Delta t) - 2 \cdot U(x,t)}{\Delta t^2} \quad (\text{S3-4})$$

419 When a reasonable spatial step Δx and time step Δt are chosen, the entire
420 computational domain is discretized into a spatial grid and a temporal grid. The spatial
421 grid consists of two multidimensional arrays representing the displacements in the
422 computational domain and the accelerations in the computational domain: the
423 accelerations and displacements at positions i, j at the k th instant are denoted as $U_{i,j}^k$
424 and $a_{i,j}^k$.

425 In the numerical simulation of elastic waves in the computational domain, it is
426 necessary to calculate from the moment 0 at the beginning of the time interval to the

427 end of the time interval. The central difference format of the time-domain finite-
428 difference method is to extrapolate the displacement of the computational domain at
429 the next moment from the displacement at the current moment and the historical
430 displacement point by point. At the time t_0 , the acceleration at a point in the
431 computational domain is $a_{i,j}^k$. At the time $t_0 + \Delta t$, the computational domain is excited
432 by an external force and the acceleration of the entire computational domain is denoted
433 as $a_{i,j}^{k+1}$. In the above equation, the acceleration layer $a_{i,j}^k$ at the moment of t_0 should
434 be calculated from the displacement layer $U_{i,j}^k$ as shown in Eqn. S3-5.

$$a_i^k = \frac{U_{i+1}^k + U_{i-1}^k - 2U_i^k}{\Delta t^2} \quad (\text{S3-5})$$

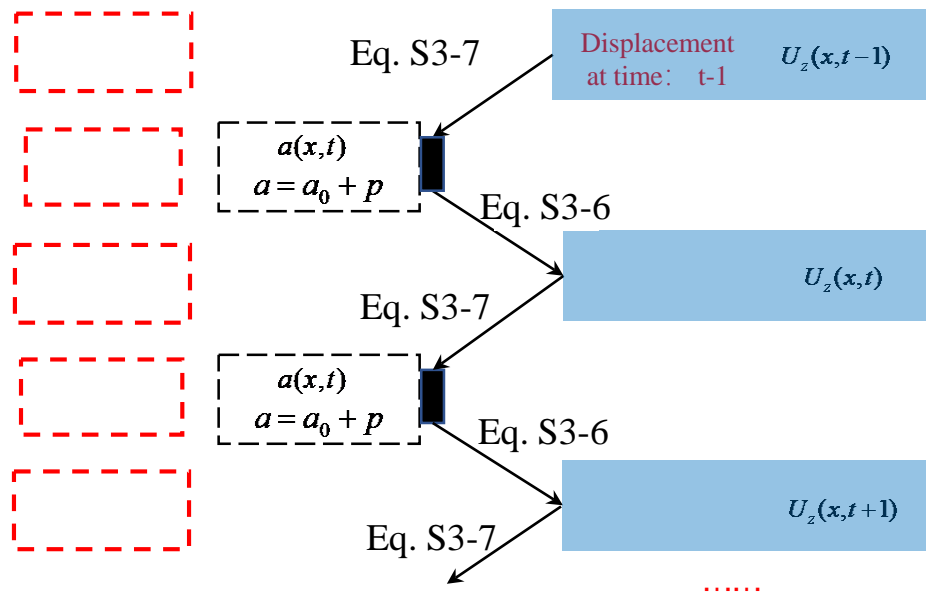
435 When the computational domain is affected by an external force at time t_0 , the
436 acceleration distribution a_i^{k+1} of the whole computational domain, which consists of
437 the acceleration of the whole domain at time t_0 and the acceleration of the affected
438 domain after the external force, is shown in Eqn. S3-6.

$$a_i^k = a_{0i}^k + a_{pi}^k \quad (\text{S3-6})$$

439 Eqn. S3-6 is the value of the acceleration layer at the moment t_0 after the
440 specified acceleration has been loaded. Based on the acceleration distribution in the
441 calculation domain of Eqn. S3-6, the displacement $U_{i,j}^{k+1}$ at the next moment can be
442 back-calculated, as shown in Eqn. S3-7.

$$U_{i,j}^{k+1} = \Delta t^2 \cdot a_{i,j}^k + 2U_{i,j}^k - U_{i,j}^{k-1} \quad (\text{S3-7})$$

443 The acceleration and displacement layers are interconverted as shown in Fig. S3.



444

445

Fig. S3 Acceleration layer loading method

446