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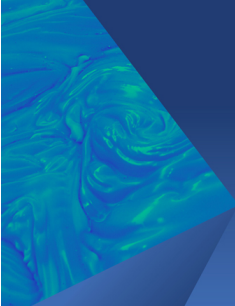


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## AFFILIATIONS

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## ABSTRACT

Fluid-filled fracture propagation is a complex problem that is ubiquitous in geosciences, from controlling magma propagation beneath volcanoes to water transport in glaciers. Using scaled analog experiments, we characterized the internal flow inside a propagating flux-driven fracture and determined the relationship between flow and fracture evolution. Different flow conditions were created by varying the viscosity and flux ( $Q$ ) of a Newtonian fluid injected into an elastic solid. Using particle image velocimetry, we measured the fluid velocity inside the propagating fracture and mapped the flow across the crack plane. We characterized the internal flow behavior with the Reynolds number ( $Re$ ) and explored  $Re$  values spanning five orders of magnitude, representing very different internal force balances. The overall fracture tip propagation velocity is a simple linear function of  $Q$ , whereas the internal velocity, and  $Re$ , may be vastly different for a given  $Q$ . We identified four flow regimes—viscous, inertial, transitional, and turbulent—and produced viscous and inertial regimes experimentally. Both flow regimes exhibit a characteristic flow pattern of a high-velocity central jet that develops into two circulating vortices on either side. However, they exhibit the opposite behavior in response to changing  $Q$ : the jet length increases with  $Q$  in the inertial regime, yet decreases in the viscous regime. Spatially variable, circulating flow is vastly different from the common assumption of unidirectional fracture flow and has strong implications for the mixing efficiency and heat transfer processes in volcanic and glacial applications.

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## I. INTRODUCTION

Fluid-filled fracture propagation is a fundamental process in many geoscience applications, including magma transport,<sup>1–3</sup> glacier dynamics and stability,<sup>4–6</sup> and geothermal energy systems.<sup>7–9</sup> Magma-filled fractures (dykes and fissures) feed volcanic eruptions, while glacial fractures (crevasses) control the drainage of glacial lakes and the transport of melt water. The fluid dynamics within propagating fractures has a significant effect on the overall fracture behavior. Propagation is driven by internal fluid pressure (due to fluid injection, buoyancy, or a combination of the two), which is distributed and dissipated by the internal flow.<sup>10</sup> In dykes and fissures, the flow of magma influences the style of eruption at the surface.<sup>11,12</sup> Flow in glacial crevasses can have a significant impact on glacier stability and melting rates.<sup>13–15</sup> Understanding and predicting fracture behavior, including the expected pathway, propagation rate, and internal fluid dynamics, is essential for managing the risks associated with volcanic and climate change processes. Despite its importance, the fluid flow within propagating fractures is not well understood and is typically assumed to be

unidirectional—a key assumption of many theoretical and numerical models of fracture propagation.<sup>2,16–18</sup>

A major challenge in modeling fracture propagation and internal fluid flow is having a unified understanding of the full range of potential behavior. Some theoretical and numerical models neglect fluid flow and assume that buoyancy dominates, and have been able to recreate fracture pathways in experiments and in nature.<sup>19,20</sup> However, buoyancy-driven fractures only represent a subset of natural cases, and fluid flow must be included to obtain accurate predictions of propagation velocities.<sup>16,17</sup> Flux-driven fractures are driven by the pressure created by fluid injection, where buoyancy may not play any role.<sup>2,21,22</sup> In theoretical and numerical models of flux-driven fractures, flow, as characterized by the dimensionless Reynolds number  $Re$ , is typically assumed to be in one of two limiting regimes: viscosity-dominated<sup>2,23–25</sup> ( $Re \ll 1$ ) or turbulent<sup>26–28</sup> ( $Re > 1000$ ). In reality, fracture flow spans a wide range of regimes due to the vast natural parameter space, notably the fluid viscosity. Fluid in glacial fractures and geothermal systems has a viscosity of the order  $10^{-3}$  Pa s, yet for magma this varies

between  $10^{-2}$  Pa s (for ultramafic low-silica magmas) and  $10^9$  Pa s (for evolved, silica-rich rhyolite).<sup>29</sup> In nature, Reynolds numbers range from the order of  $10^{-10}$  for viscous, creeping dykes,<sup>30</sup> to  $10^6$  and beyond for turbulent crevasses during rapid drainage events.<sup>5</sup> There is, therefore, a strong motivation to understand fracture propagation for the full range of potential flow regimes and  $Re$  values, particularly in the transition from viscosity-dominated flows to full turbulence.<sup>28,31</sup>

Scaled, analog experiments of fluid-filled fracture propagation give crucial insight into the fundamental processes of fracture dynamics (see Rivalta *et al.*<sup>17</sup> and Kavanagh *et al.*<sup>32</sup> for a review). Laboratory experiments involving the injection of fluid into solid, elastic, gelatin allow for direct observations of fracture and fluid dynamics during propagation.<sup>33,34</sup> Buoyancy-driven fracturing occurs if the injected fluid is sufficiently less dense than the solid host.<sup>35–38</sup> Otherwise, flux-driven fractures are created by the constant injection of fluid.<sup>22,32,39</sup> Recent studies have used particle image velocity (PIV) to measure internal flow velocity profiles in flux-driven fractures.<sup>32,40,41</sup> While flow in buoyancy-dominated fractures is confirmed to have a simple unidirectional profile,<sup>41</sup> Newtonian flux-driven fractures exhibit a more complex flow pattern, consisting of a central, localized jet, with circulating downward flow along the fracture margins,<sup>32,40,41</sup> which is not captured with any existing numerical model. Only a small number of published experiments (all consisting of water injections with Reynolds numbers in the narrow range  $1 \lesssim Re \lesssim 30$ ) have captured this interesting flow pattern. Experimental data across a wider  $Re$  range are required to fully understand flux-driven fracture propagation and the influence of internal fluid flow.

In this study, we provide the first experimental investigation of the dependence of fracture dynamics on the Reynolds number. We restrict our attention to flux-driven fractures (that are not buoyant) and conduct a series of experiments where a Newtonian fluid is injected into gelatin at a constant rate. We systematically vary  $Re$  by changing the viscosity of the injected fluid and its injection rate, achieving flows in the range  $O(10^{-3}) \leq Re \leq O(10^2)$ . Propagation velocities and internal fluid flow profiles are measured across a two-dimensional plane of the growing fracture for the full duration of the experiment. Our results showcase the complex fluid dynamics inside flux-driven fractures and the relationship with propagation velocities. A jet and recirculation is a universal feature of Newtonian flux-driven flows, yet there are key differences between viscous and inertial flow regimes. We discuss the physics behind this observed behavior and the implications for natural flux-driven fractures in glacial and magmatic settings.

## II. THEORETICAL FRAMEWORK

In this section, we present the relevant theoretical framework behind the experimental flux-driven fractures, particularly related to the internal fluid flow (the main focus of this study). We consider the following simplifications: a single, vertical fracture is driven by a constant continuous flux; the host is an isotropic, non-porous, elastic solid; the injected fluid is Newtonian and non-buoyant; fractures are tensile, opening in the direction of the least compressive stress. We adopt linear elastic fracture mechanics (LEFM)<sup>3,42</sup> and assume that the solid resistance to fracture is characterized by the fracture toughness  $K_C$  ( $\text{Pa}\cdot\text{m}^{1/2}$ ). Fracture propagation occurs if the stress intensity  $K$  at the tip (a function of fluid pressure gradients) equals a critical value  $K_C$ . We assume that  $K_C$  is constant, although experiments suggest that  $K_C$  may change with fracture length.<sup>37</sup> These simplifications are common

assumptions in mathematical and numerical models of flux-driven fractures.<sup>2,18,43,44</sup>

### A. Equations of fluid motion

Flux-driven fracture propagation requires the continuous injection of fluid as this provides a driving pressure gradient at the source, which is distributed to the fracture tips via fluid flow. The evolution of pressure and fluid flow are governed by the Navier–Stokes equations, consisting of the conservation of mass and momentum,

$$\nabla \cdot \mathbf{u} = 0, \quad (1)$$

$$\rho_f \frac{\partial \mathbf{u}}{\partial t} + \rho_f \mathbf{u} \cdot \nabla \mathbf{u} = -\nabla p + \mu \nabla^2 \mathbf{u}. \quad (2)$$

Here,  $\mathbf{u} = (u_x, u_z)$  is the fluid velocity,  $\rho_f$  is the fluid density,  $p$  is the dynamic fluid pressure (i.e., excess of hydrostatic),  $t$  is the time, and  $\mu$  is the dynamic viscosity. Along with suitable boundary conditions, Eqs. (1) and (2) describe incompressible, Newtonian flow inside a flux-driven fracture.

#### 1. Boundary conditions

Injection of fluid can be expressed as a flux boundary condition,

$$\mathbf{u} \cdot \mathbf{n} = \frac{Q}{A}, \quad \text{at the inlet}, \quad (3)$$

where  $\mathbf{n}$  is the unit normal direction to the inlet flow,  $Q$  is the volumetric flux ( $\text{m}^3/\text{s}$ ) and  $A$  ( $\text{m}^2$ ) is the surface area of the inlet.  $Q/A$  is the fluid injection velocity, also written as  $u_{in}$ . Note that Eq. (3) is equivalent to imposing a pressure gradient at the inlet.

Fluid flow satisfies the no-slip condition:

$$\mathbf{u} = \mathbf{u}_{tip}, \quad \text{at the solid–fluid interface}, \quad (4)$$

where  $\mathbf{u}_{tip}$  is the velocity of the fracture tip.

#### 2. Fluid forces

In the momentum equation (2), dynamic pressure gradients (units of force per unit volume,  $\text{N}/\text{m}^3$ ) are balanced with two forces, viscous [ $\mathbf{F}_V = (F_{Vx}, F_{Vz})$ ] and inertial [ $\mathbf{F}_I = (F_{Ix}, F_{Iz})$ ],

$$\nabla p = \mathbf{F}_I - \mathbf{F}_V, \quad (5)$$

$$\mathbf{F}_I = \rho_f \frac{\partial \mathbf{u}}{\partial t} + \rho_f \mathbf{u} \cdot \nabla \mathbf{u}, \quad (6)$$

$$\mathbf{F}_V = \mu \nabla^2 \mathbf{u}. \quad (7)$$

In addition to driving fracture growth, the imposed pressure gradient due to fluid injection is also dissipated by viscous and inertial forces.

### B. Opposing pressure scales

The forces that oppose fracture propagation can be represented by simple pressure scales.<sup>2,37,41,43,44</sup> Solid resistance to fracture is represented by the fracture pressure scale  $P_f$

$$P_f \sim \frac{K_C}{\min(L, W)^{1/2}}, \quad (8)$$

where  $\min(L, W)$  is the minimum of the fracture length  $L$  and width  $W$ .

The viscous pressure scale  $\Delta P_V$  represents the drop in pressure along the fracture due to viscous resistance,

$$\Delta P_V \sim \frac{3\mu u L}{H^2}, \quad (9)$$

where  $\mu$  is the viscosity,  $H$  is the fracture thickness, and  $u$  is the average internal velocity. This scaling (9) follows from the Navier–Stokes equations (1) and (2) under lubrication theory assumptions: laminar, unidirectional flow with negligible inertia.<sup>1,2,18,25</sup>

In high  $Re$  flows, inertial effects are important and pressure is dissipated via fluid kinetic energy.<sup>45</sup> The inertial pressure scale  $\Delta P_I$  is derived through neglecting viscosity from the full equations of fluid motion.  $\Delta P_I$  represents the loss of fluid pressure to inertial forces,

$$\Delta P_I \sim \frac{f_D \rho_f u^2 L}{2}, \quad (10)$$

where  $f_D$  is a complex empirical function of the friction factor and fracture geometry.<sup>5,46</sup> Closures for (10) have been proposed for turbulent fracture flow,<sup>5,26</sup> yet there is very little focus on the transition from viscosity-dominated laminar flow and full turbulence.<sup>46</sup>

### C. Dimensionless numbers

Flux-driven fractures propagate in different regimes according to the dominant resistive processes. Fractures in gelatin are expected to propagate in the toughness regime,<sup>32,37</sup> where the dominant opposing pressure scale is the fracture pressure  $P_F$ . In the toughness regime, the solid fracture process uses more energy than fluid forces, yet the relative balance of internal inertial and viscous forces still influences the overall fracture dynamics.<sup>17,47</sup> Denoting the characteristic flow velocity with  $\mathcal{U}$  and the characteristic length scale with  $\mathcal{L}$ , the magnitude of the fluid force terms can be estimated as follows:<sup>48</sup>

$$|F_I| = |\rho_f \mathbf{u} \cdot \nabla \mathbf{u}| \sim \frac{\rho_f \mathcal{U}^2}{\mathcal{L}}, \quad (11)$$

$$|F_V| = |\mu \nabla^2 \mathbf{u}| \sim \frac{\mu \mathcal{U}}{\mathcal{L}^2}. \quad (12)$$

The Reynolds number  $Re$  represents the ratio of inertial to viscous forces,

$$Re = \frac{|\rho_f \mathbf{u} \cdot \nabla \mathbf{u}|}{|\mu \nabla^2 \mathbf{u}|} = \frac{\rho_f \mathcal{U} \mathcal{L}}{\mu}. \quad (13)$$

For fracture flows,  $\mathcal{L}$  is the fracture thickness  $H$ , and  $\mathcal{U}$  can be approximated as  $\mathcal{U} \approx Q/WH$ .<sup>41,49</sup> This reduces  $Re$  to

$$Re_0 = \frac{\rho_f Q}{\mu W}. \quad (14)$$

In the toughness regime, two dimensionless numbers describe the relative effects of viscous and inertial forces to the fracture resistance. These are known as the dimensionless viscosity  $\mu_k$  and dimensionless inertia  $R_k$ ,

$$\mu_k = \frac{12\mu Q'}{E'} \left( \frac{E'}{K'} \right)^4, \quad (15)$$

$$R_k = \frac{\rho E'^{5/3} Q'^{5/3}}{K'^{8/3} t^{1/3}}, \quad (16)$$

where  $Q' = Q/W$  is the volumetric flux per unit width ( $\text{m}^2/\text{s}$ ) and

$$E' = \frac{E}{1 - \nu^2}, \quad K' = 4\sqrt{2/\pi} K_C, \quad (17)$$

where  $E$  is the Young's modulus and  $\nu$  is Poisson's ratio. Note that  $R_k$  is a decreasing function of time, while  $\mu_k$  is constant.  $R_k$  and  $\mu_k$  were derived from an idealized, plane-strain, flux-driven fracture model that quantifies the coupled effects of fracture resistance, viscosity and inertia.<sup>27</sup> We expect  $\mu_k$  and  $R_k$  to be small ( $<1$ ) in the toughness regime as fracture resistance dominates.

### III. METHODOLOGY

Here, we describe the experimental process in detail. We first provide an overview of the experimental setup and the materials used, followed by a description of the PIV method for measuring internal fracture velocities during propagation. We then provide details on post-processing the experimental data, including the calculation of representative fracture velocities, forces, and Reynolds numbers.

#### A. Overview

A series of experiments were conducted to establish the effect of  $Re$  on both the internal fluid dynamics and the overall propagation of flux-driven fractures. Each experiment consisted of a Newtonian fluid being injected into a  $40 \times 40 \times 25 \text{ cm}^3$  volume of transparent, solid, elastic gelatin held in a clear Perspex tank. An initial, vertical, precut of 3 cm length and 1 cm width was created in the center of the base of the gelatin using a thin blade. The fluid was injected into the precut using a needle with its tapered edge orientated parallel to the widest part of the precut. The needle had an inlet diameter  $d$  of either 1 or 2 mm and an elliptical opening surface area  $A$  of either  $\pi \times 1 \times 3.5 \text{ mm}^2$  or  $\pi \times 2 \times 4 \text{ mm}^2$ . The needle was connected to a fluid reservoir via 5 mm diameter tubing. A valve on the pipe and a small amount of petroleum jelly added to the end of the needle ensured all air was removed from the injection system prior to starting an experiment. A peristaltic pump was then used to push fluid through the tube and into the gelatin at a known, constant rate, creating a flux-driven penny-shaped fracture that propagated vertically and erupted at the surface through a thin fissure. The precut controlled the fracture orientation, ensuring it grew vertically in the  $z$  direction and radially in the  $x$ - $z$  plane, while pushing open the solid as a tensile fracture in the  $y$ - $z$  plane (see Fig. 1). Two-dimensional (2D) internal velocity profiles were measured in the  $x$ - $z$  plane using a laser-based PIV system (see Sec. III C), controlled via LaVision's DaVis ten specialized laser imaging software.<sup>50,51</sup> The laser-imaging system and experimental tank were all supported by a robust, connected metal frame that ensures experiment repeatability.

#### B. Materials

Different flow regimes were achieved by injecting Newtonian fluids with different viscosity but similar density: a high viscosity fluid (silicone oil,  $\mu = 0.45 \text{ Pa s}$ ,  $\rho_f = 998 \text{ kg/m}^3$ ) and a low viscosity fluid (water,  $\mu = 0.001 \text{ Pa s}$ ,  $\rho_f = 998 \text{ kg/m}^3$ ). The viscosity of silicone oil was determined with a series of rheometer tests at different temperatures, and the density was obtained using a 100 ml pycnometer.

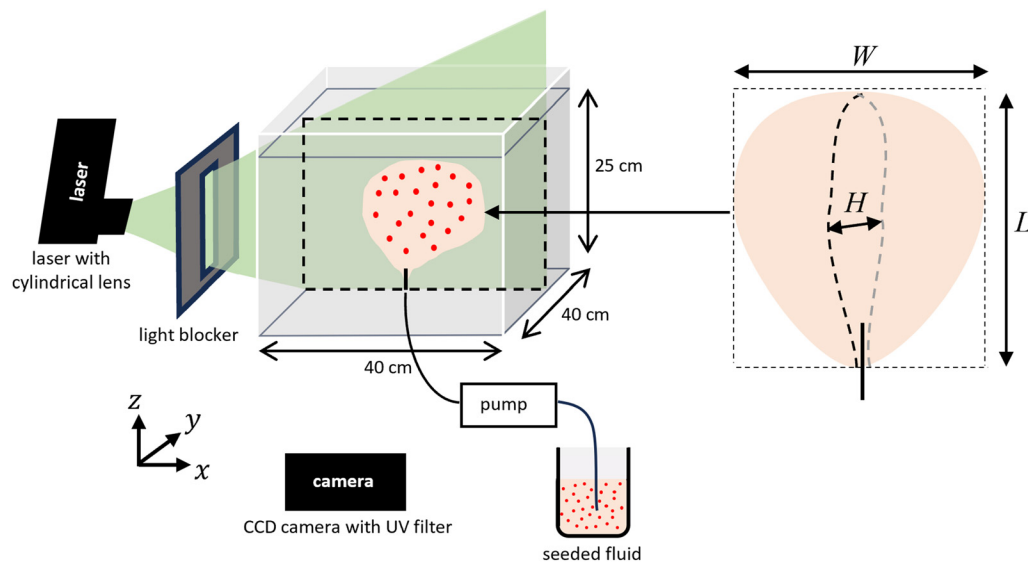


FIG. 1. Schematic of the experimental setup depicting a growing flux-driven penny-shaped fracture being illuminated with a laser sheet.

The solid, elastic gelatin had a concentration of 2.5 wt. % and was prepared following the guidelines of Kavanagh *et al.*,<sup>32,52</sup> resulting in  $1001.5 \text{ kg/m}^3$  solid density.<sup>32</sup> Gelatin preparation involves mixing 1 kg of gelatin powder (260 Bloom, ten mesh, pig-skin gelatin supplied by Gelita UK) with 39 kg of de-ionized water, resulting in a total liquid mass of 40 kg. Approximately half of the total amount of water was added hot ( $\approx 80^\circ \text{C}$ ) to initially dissolve the gelatin, and the rest was added cold ( $\approx 7 - 10^\circ \text{C}$ ) directly in the tank. The liquid mixture was then covered in a thin layer of vegetable oil and covered with plastic wrapping. It was left to cool and solidify in a refrigerator for approximately 41–50 hours to obtain a Young's modulus  $E$  in the range 3000–5000 Pa (note that the addition of cold water allowed for shorter solidification times than if using hot water only).  $E$  was measured immediately before running an experiment using the method of Kavanagh *et al.*,<sup>52</sup> which involves removing the surface oil and then applying different loads to the center of the gel surface and measuring their deflections. Gelatin's fracture toughness  $K_C$  can be approximated as  $K_C = 1.4\sqrt{E}$ .<sup>52</sup>

### C. The PIV system

Planar PIV was used to measure horizontal and vertical fluid velocities inside the fracture in the  $x$ - $z$  plane. The injected fluid was pre-seeded with Rhodamine B-coated tracer particles with diameters of  $20\text{--}50 \mu\text{m}$  and a particle density of  $1190 \text{ kg/m}^3$ . Calculations of Stokes settling velocity  $U_g$ , relaxation time  $\tau_r$ , and Stokes number  $St$  suggest that these particles suitably trace the fluid streamlines (see Sec. I of the supplementary material). Successive images of these passive tracer particles were used to track fluid motion and compute velocities (see the videos in the supplementary material). The PIV method<sup>53</sup> divides each image into subdomains of a defined size (here either  $32 \times 32$  or  $24 \times 24$  pixels) and applies a statistical correlation technique to produce a single velocity vector per subdomain (with a calculation overlap of 75% between subdomains). The time interval

between recorded images is chosen so that there is optimal particle displacement between successive images (of approximately 5 pixels<sup>51</sup>). PIV has been used to measure velocities in laboratory flows with a wide variety of geophysical applications.<sup>32,54–58</sup>

The tracer particles were fluoresced with a sheet of light emitted from a Class 4 532 nm DoublePulse Nd:YAG Litron laser (maximum energy  $2 \times 325 \text{ mJ}$ ), which illuminated the expected plane of fracture growth. Laser output was synchronized to an Imager SX 6M CCD camera facing the  $x$ - $z$  plane, positioned perpendicular to the light sheet. The camera had a resolution of  $2752 \times 2200$  pixels and was used with a Zeiss 50 mm  $f/1.4$  lens with an aperture of  $f/5.6$ . The lens was fitted with a UV filter that blocks out short wavelengths and prevents reflections from the gelatin. Note that when measuring 2D flow, the light sheet should be as thin as possible to reduce out-of-plane motion effects. However, the fracture needs to remain within the sheet to ensure that the tracer particles are recorded (it is not guaranteed to be perfectly vertical). We used a 5 mm-thick light sheet as a compromise between reducing potential three-dimensional (3D) effects and ensuring flow visualization. The laser was output through a  $\sim 25 \text{ mm}$  convex cylindrical lens (from Edmund Optics Ltd). Magnetic blockers (with an adjustable gap in the middle) positioned between the laser output and tank wall transform the laser light into a thin sheet, orientated along the center of the tank (see Fig. 1).

Images were captured in either single-frame or double-frame mode,<sup>53</sup> depending on the expected magnitude of particle displacements. In single-frame mode, each image consists of a single frame that records the emission of two laser pulses with an exposure time of  $42 \mu\text{s}$ . The shortest time interval that can be achieved between subsequent frames in single-frame mode is restricted by the maximum camera frame rate of 15 frames per second. However, double-frame recordings allow for shorter time intervals. In double-frame mode, each image is composed of two frames separated by an interval  $\Delta t$ , which is achieved via control of the camera shutter. This also defines the separation between the two laser pulses, so that the first frame

**TABLE I.** Solid and fluid material parameters for experiments S1–S5 and SH (injecting silicone oil), experiments W1–W4 and WH (injecting water). Young's modulus  $E$  (Pa), volumetric flux  $Q$  ( $\text{m}^3/\text{s}$ ), fluid viscosity  $\mu$  (Pa s), fluid density  $\rho$  ( $\text{kg}/\text{m}^3$ ), inlet diameter  $d$  (mm), average thickness  $H$  (mm) (including range and associated experiment), and inlet Reynolds number  $Re_{in}$ .

	$E$	$Q$	$\mu$	$\rho$	$d$	$H$	$Re_{in}$
S1	4337	$4.37 \times 10^{-8}$	0.450	998	1	11 (7–15, SH)	0.009
S2	4098	$7.20 \times 10^{-8}$	0.450	998	1	11 (7–15, SH)	0.015
S3	4170	$1.34 \times 10^{-7}$	0.450	998	1	11 (7–15, SH)	0.027
S4	4506	$2.60 \times 10^{-7}$	0.450	998	2	11 (7–15, SH)	0.046
S5	4214	$4.07 \times 10^{-7}$	0.450	998	1	11 (7–15, SH)	0.092
W1	4309	$4.68 \times 10^{-7}$	0.001	998	1	4 (3–5, WH)	36.650
W2	2593	$3.15 \times 10^{-6}$	0.001	998	2	4 (3–5, WH)	250.088
W3	3591	$6.97 \times 10^{-6}$	0.001	998	1	4 (3–5, WH)	633.005
W4	2278	$6.97 \times 10^{-6}$	0.001	998	1	4 (3–5, WH)	633.005
SH	3946	$4.07 \times 10^{-7}$	0.450	998	1	11 (7–15)	0.092
WH	3591	$6.97 \times 10^{-6}$	0.001	998	1	4 (3–5)	633.005

captures the first emission, and the second frame captures the second pulse. The exposure of the first frame is defined by  $\Delta t$ , whereas the exposure of the second frame cannot be controlled and may capture more ambient light. The recording settings for each experiment are provided in Table I in the supplementary material.

Prior to running an experiment, the camera was first focused on fluoresced particles in a tank full of seeded water, with the laser and camera in position. A calibration procedure was then performed (within the DaVis software) where images were taken of a calibration board positioned in the imaging plane. Pixels are automatically converted to material coordinates in subsequent DaVis operations. The calibration procedure was conducted prior to running an experiment and required that the camera and laser positions were kept at a fixed position relative to the tank and imaging plane—this was ensured via the supportive frame.

## D. Data processing

### 1. Post-processing PIV data

Erroneous PIV velocity vectors were removed in “Vector Post-Processing” in DaVis<sup>50</sup> according to a threshold set by the correlation value  $r_c$  (the degree of confidence in the statistical correlation procedure). A threshold of  $r_c = 0.2$  eliminated vectors lying outside of the seeded fracture flow.

Velocity data were exported from DaVis as a series of csv files (one for each time step), containing the velocity components  $u_x$  and  $u_z$  and the corresponding spatial coordinates  $x$  and  $z$ . All further analysis was performed in Matlab.<sup>59</sup> Data were imported using the `readtable` function, and each variable was converted to a 2D grid and processed with median filtering and Gaussian smoothing functions (`medfilt2` and `smoothdata`). Velocity data collected in single-frame mode were time averaged over an interval representing 5% of the experimental duration, resulting in an averaged velocity profile and standard deviation. Double-frame velocity data were not time averaged due to the time separation between two successive images being greater

than 5%. All processing scripts are available in an accompanying data publication.<sup>60</sup>

### 2. Tracking fracture geometry

The fracture outline was extracted from the raw images by cropping around the illuminated particles. Images were first converted to binary using `im2bw` (from the Image Processing Toolbox<sup>61</sup>) and an appropriate pixel intensity threshold was selected to distinguish black from white. The binary image was reduced in size and outliers removed, before applying the `rangesearch` function (in the Statistics and Machine Learning Toolbox<sup>62</sup>) to remove individual pixels with fewer than a specified number of neighbors. The boundary function was applied to detect the bounding shape of the reduced set of pixels, and the coordinates were converted from pixels to mm. An ellipse was fitted to the boundary points (with function `fitellipse`<sup>63</sup>). Fracture length  $L$  and width  $W$  were calculated from the length and width of the fitted ellipse at its center point.

### 3. Representative velocities

Different representative velocities are used to characterize the fracture at a given time, including tip velocities and internal flow velocities. Tip velocities consist of the fracture propagation rate in the vertical and horizontal directions— $u_{tip}$  and  $u_W$ , respectively. Representative flow velocities include the spatially averaged mean velocity  $u_{mean}$ , a representative jet velocity  $u_{jet}$ , a representative downwards velocity  $u_{down}$ , and a circulation velocity<sup>41</sup>  $u_{circ}$ . The latter represents the degree of internal flow circulation:  $u_{circ} = (u_{jet} - u_{down})/u_{jet}$ . A value of  $u_{circ} = 1$  means that there is zero downwards flow and no circulation, while  $u_{circ} = 2$  corresponds to the downwards velocity being of equal magnitude to the upwards velocity, indicating strong circulation.

To obtain values for  $u_{jet}$  vectors within the jet region were systematically cropped in each experiment and averaged in this area (within the 65–90 percentile range); a comparison with full velocity contours confirmed that this method gives a velocity value that is representative of the jet. A similar method was applied to get  $u_{down}$  where the data were instead cropped near the lateral fracture margins, and filtered according to  $u_z < 0$ . Tip velocities  $u_{tip} = dL/dt$  and  $u_W = dW/dt$  were calculated by first fitting third-order polynomials to the temporal evolution of  $L$  and  $W$ , before integration over  $t$ . Error terms for  $u_{jet}$ ,  $u_{mean}$ , and  $u_{down}$  were calculated using the corresponding velocity standard deviation terms (note that these errors could not be obtained for the non-time-averaged double-frame experiments W2–W4). Errors for  $u_{tip}$  and  $u_W$  were approximated by considering the difference between the temporal geometry data ( $L$ ,  $W$ ) and their fitted curves.

### 4. Calculating fluid forces and pressure scales

For each experiment, the viscous and inertial forces (7) can be calculated using the measured velocity data  $\mathbf{u}_{i,j}$ . Here,  $\mathbf{u}_{i,j}$  denotes a single PIV grid measurement where indices  $i$  and  $j$  represent the spatial location in the  $x$  and  $z$  directions, respectively. Adjacent grid points are separated by a constant distance  $\Delta x$  in both directions. Two adjacent points are denoted by  $\mathbf{u}_{i,j}$  and  $\mathbf{u}_{i+1,j}$  in the  $x$  axis, and  $\mathbf{u}_{i,j}$  and  $\mathbf{u}_{i,j+1}$  in the  $z$  axis. First- and second-order velocity derivatives are calculated

using a finite difference method<sup>64</sup> and substituted into (7) to obtain approximations of inertial and viscous forces at a given time. These forces have a horizontal and vertical component (in 2D), such that  $\mathbf{F}_v = (F_{vx}, F_{vz})$  and  $\mathbf{F}_I = (F_{Ix}, F_{Iz})$ . The numerical (finite difference) approximations of the force terms (denoted with a  $\hat{\cdot}$  notation) are defined as follows:

$$\hat{F}_{vxi,j} \approx \frac{\mu}{\Delta x^2} (u_{xi-1,j} + u_{xi+1,j} + u_{xi,j-1} + u_{xi,j+1} - 4u_{xi,j}), \quad (18)$$

$$\hat{F}_{vzi,j} \approx \frac{\mu}{\Delta x^2} (u_{zi-1,j} + u_{zi+1,j} + u_{zi,j-1} + u_{zi,j+1} - 4u_{zi,j}), \quad (19)$$

$$\hat{F}_{Ixi,j} \approx \rho_f \frac{\partial u_{xi,j}}{\partial t} + \frac{\rho_f}{\Delta x} (u_{xi,j}(u_{xi+1,j} - u_{xi,j}) + u_{zi,j}(u_{xi,j+1} - u_{xi,j})), \quad (20)$$

$$\hat{F}_{Izi,j} \approx \rho_f \frac{\partial u_{zi,j}}{\partial t} + \frac{\rho_f}{\Delta x} (u_{xi,j}(u_{zi+1,j} - u_{zi,j}) + u_{zi,j}(u_{zi,j+1} - u_{zi,j})). \quad (21)$$

These terms are spatially averaged across the full fracture profile to approximate the average (absolute) forces (viscous  $\overline{F}_v$  and inertial  $\overline{F}_I$ ) at a given time,

$$|\overline{F}_v| = \frac{1}{N-1} \sum_{i=2}^{N-1} |\hat{F}_{vi,j}|, \quad (22)$$

$$|\overline{F}_I| = \frac{1}{N-1} \sum_{i=2}^{N-1} |\hat{F}_{Ii,j}|, \quad (23)$$

where  $N$  is the total number of grid points and  $|\hat{F}_{vi,j}|$  and  $|\hat{F}_{Ii,j}|$  denote the modulus of the numerical viscous and inertial force terms at a single point.

Estimates of the viscous and inertial pressure scales are then obtained by multiplying the average force by the fracture length,

$$\Delta \hat{P}_v \approx |\overline{F}_v|L, \quad (24)$$

$$\Delta \hat{P}_I \approx |\overline{F}_I|L. \quad (25)$$

### E. Reynolds number calculations

The Reynolds number summarizes the bulk flow behavior as a single parameter, yet in reality flow can be spatially and temporally variable, with a range of characteristic velocity and length scales. In addition to  $Re_0 = \rho_f Q / \mu W$  (14), we explore several alternative Reynolds numbers using different characteristic velocities and length scales, which may potentially better represent the force balance during fracture flow. We define four alternative Reynolds numbers: (1) the inlet Reynolds number,  $Re_{in}$ , (2) the tip Reynolds number,  $Re_{tip}$ , (3) the jet Reynolds number  $Re_{jet}$ , and (4) the mean Reynolds number  $Re_{mean}$ . The flow at the source of fluid injection is characterized by  $Re_{in}$  and is known prior to running an experiment. Flow Reynolds numbers  $Re_{tip}$ ,  $Re_{jet}$  and  $Re_{mean}$  represent the internal flow during fracture propagation and require measured velocity values (see Sec. III D 3).

#### 1. Inlet Reynolds number, $Re_{in}$

$Re_{in}$  represents the fluid force balance at the inlet and does not require any information about the fracture flow or geometry. This is defined as

$$Re_{in} = \frac{\rho_f u_{in} d}{\mu}. \quad (26)$$

#### 2. Tip Reynolds number, $Re_{tip}$

$Re_{tip}$  uses the vertical fracture tip velocity  $u_{tip}$  as the characteristic velocity scale, and  $H$  as the length scale,

$$Re_{tip} = \frac{\rho_f u_{tip} H}{\mu}. \quad (27)$$

#### 3. Jet Reynolds number, $Re_{jet}$

$Re_{jet}$  represents the internal flow behavior, with the characteristic velocity defined as the jet velocity  $u_{jet}$ ,

$$Re_{jet} = \frac{\rho_f u_{jet} H}{\mu}. \quad (28)$$

#### 4. Mean Reynolds number, $Re_{mean}$

$Re_{mean}$  also represents the internal flow, but instead uses the mean internal velocity  $u_{mean}$  as the characteristic value,

$$Re_{mean} = \frac{\rho_f u_{mean} H}{\mu}. \quad (29)$$

#### 5. Fracture thickness measurements $H$

The fracture thickness  $H$  is required to calculate the representative flow Reynolds numbers  $Re_{mean}$ ,  $Re_{jet}$ , and  $Re_{tip}$ . We conducted experiments to approximate  $H$  for low-viscosity, water fractures, and high-viscosity silicone oil fractures, respectively. The fracture evolution was instead recorded in the  $y$ - $z$  plane, using either a seeded fluid or a seeded gelatin (with no fluid seeding, see Sec. III in the supplementary material). The representative thickness was approximated via manual image inspection within the DaVis software.

### IV. RESULTS

In total, eleven experiments were completed with  $Re_{in}$  ranging from 0.009 to 633 (Table I). All experiments produced a broadly penny-shaped fluid-filled crack that grew and eventually erupted at the surface. Nine experiments measured fluid velocities in the  $x$ - $z$  plane [five silicone oil (S) and four water (W) injections] and two experiments measured a representative  $H$  for the different fluid injections [one for silicone oil (SH), and one for water (WH)] (see Table II). All silicone oil injections have  $Re_{in} < 1$ , while the water injections all have  $Re_{in} > 1$ . The fracture thickness  $H$  varies with height, and measurements of  $H$  may be affected by optical distortions related to out-of-plane fracture growth and mismatching refractive indices between the fluid and gelatin (see Sec. II in the supplementary material). We, therefore, report a conservative range of  $H$  values:  $H \approx 3 - 5$  mm for water, and  $H \approx 7 - 15$  mm for silicone oil (S experiments have a high refractive index mismatch and a larger error margin, as a closer refractive index matching leads to more accurate measurements). These variations in  $H$  have a minor effect on the overall change in Reynolds number, and the average  $H$  value was used in  $Re$  calculations.

In Secs. IV A–IV D, we present the fluid flow profiles, temporal fracture evolution, and governing force balance results. Note that in

**TABLE II.** Flux-driven fracture thickness and characteristic velocity results for each experiment, at a dimensionless time of  $t^* = 0.5$ : inlet velocity  $u_{in}$  (mm/s), vertical tip velocity  $u_{tip}$  (mm/s), horizontal tip velocity  $u_W$  (mm/s), mean velocity  $u_{mean}$  (mm/s), jet velocity  $u_{jet}$  (mm/s), downward velocity  $u_{down}$  (mm/s),  $\frac{\partial u_{mean}}{\partial t}$  (mm/s<sup>2</sup>), and  $\frac{\partial u_{jet}}{\partial t}$  (mm/s<sup>2</sup>).

	$u_{in}$	$u_{tip}$	$u_W$	$u_{mean}$	$u_{jet}$	$u_{down}$	$\frac{\partial u_{mean}}{\partial t}$	$\frac{\partial u_{jet}}{\partial t}$
S1	3.975	0.023	0.021	0.050	0.379	-0.053	$6.6 \times 10^{-7}$	$4.4 \times 10^{-5}$
S2	6.551	0.040	0.023	0.084	0.686	-0.088	$5.7 \times 10^{-7}$	$5.3 \times 10^{-5}$
S3	12.209	0.076	0.054	0.124	1.229	-0.119	$5.8 \times 10^{-6}$	$3.3 \times 10^{-4}$
S4	10.347	0.258	0.308	0.247	1.925	-0.183	$3.3 \times 10^{-4}$	0.0032
S5	41.430	0.147	0.182	0.233	1.591	-0.106	$1.1 \times 10^{-4}$	$6.6 \times 10^{-4}$
W1	36.724	0.236	0.171	2.432	6.907	-2.573	$5.5 \times 10^{-4}$	-0.0042
W2	125.295	1.418	1.480	30.601	135.641	-58.565	0.052	-0.19
W3	634.273	2.539	2.929	108.404	211.867	-153.974	-0.60	1.36
W4	634.273	3.676	3.517	82.717	235.844	-124.827	-1.49	-0.80

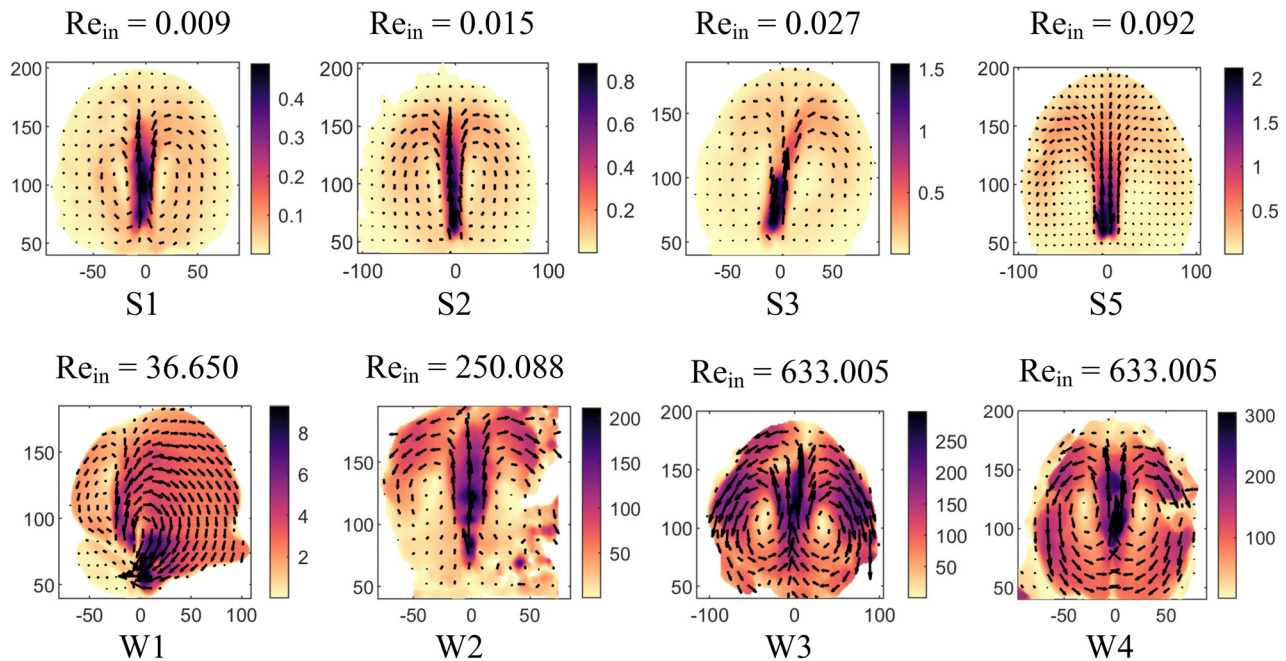
these descriptions, our main focus is the developed flow pattern (the initial flow development is presented in Fig. 3 in the supplementary material). To allow for a direct comparison between different experiments, time is normalized as  $t^* = (t - t_0)/t_{erupt}$ , where  $t_0$  is the time when  $L = 10$  cm and  $t_{erupt}$  is the time interval between  $t_0$  ( $t^* = 0$ ) and fluid eruption ( $t^* = 1$ ).

**A. Flow profiles—central jet and recirculation**

Every experiment produced a high-velocity central jet that increased in width with height, before transitioning to a recirculating

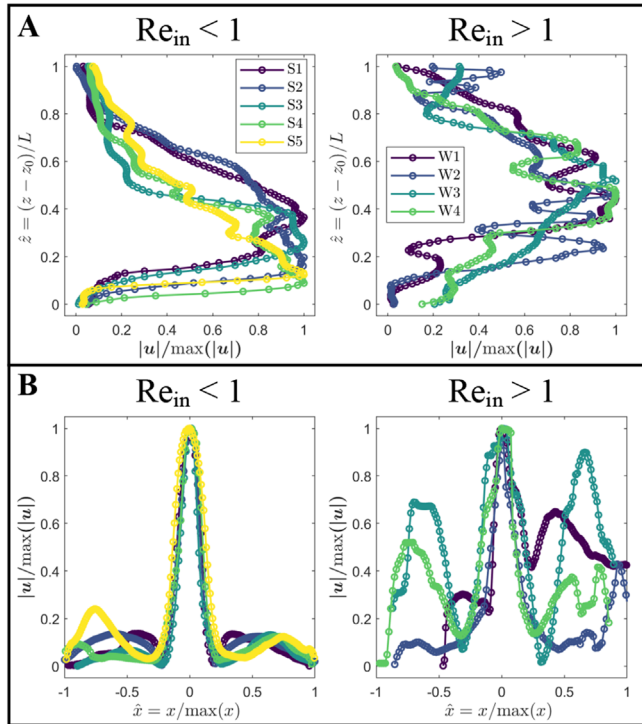
flow on either side of the jet, around two stagnant points (Fig. 2,  $t^* = 0.5$ ). There is a strong spatial velocity variation, with  $u_{jet}$  being the most dominant characteristic flow velocity and ranging from 0.379 to 236 mm/s across all experiments (Table II,  $t^* = 0.5$ ).  $u_{jet}$  is at least one order of magnitude greater than the tip velocities in every experiment (which range from 0.021 to 3.676 mm/s at  $t^* = 0.5$ ).

While all experiments have the same overall pattern of a jet and recirculating flow, there are some clear differences between the  $Re_{in} < 1$  and  $Re_{in} > 1$  experiments (Figs. 2 and 3). In the  $Re_{in} < 1$  experiments (S1–S5), flow is mostly localized in the central jet region, and the downward flow velocities are significantly lower than in the jet.



**FIG. 2.** Filled contours of velocity magnitude (mm/s) and vectors of flow direction (black arrows) for flux-driven fracture experiments, at a normalized time of  $t^* = 0.5$ . Four of the high-viscosity  $Re_{in} < 1$  silicone oil experiments (S1–S3 and S5) are shown on the top row, and the low-viscosity  $Re_{in} > 1$  experiments (W1–W4) are on the bottom row. For each row, the experiments are ordered in terms of increasing  $Q$  and  $Re_{in}$ . The vectors show the flow direction, and their size represents the velocity magnitude, scaled up by a factor of two. Only every third vector is plotted (horizontally and vertically), while the filled contours show the full resolution of the flow velocity magnitude.



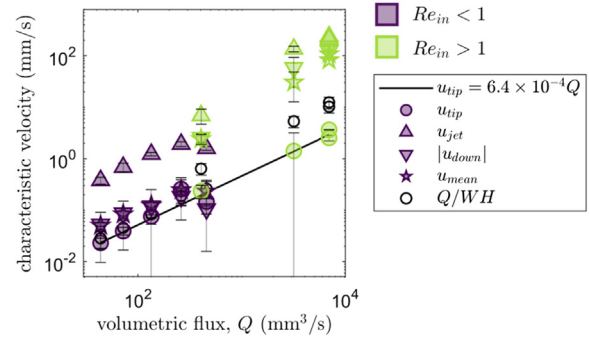


**FIG. 3.** Normalized velocity magnitude profiles at  $t^* = 0.5$ : (a) along the central vertical line above the injector and (b) along a horizontal line 60 mm above the injector, for silicon oil (S1–S5,  $Re_{in} < 1$ ) and water (W1–W4,  $Re_{in} > 1$ ) experiments. Height has been normalized so that  $\hat{z} = 0$  is just above the injector, and  $\hat{z} = 1$  is the tip location. The horizontal distance has been normalized so that the  $\hat{x} = 0$  corresponds to the location of the maximum velocity and the horizontal extent is between  $\hat{x} = \pm 1$ .

In these experiments, the jet terminates prior to reaching the vertical fracture tip, and an increase in  $Re_{in}$  correlates with a decrease in the jet height. Conversely, in all  $Re_{in} > 1$  experiments (W1–W4) the jet reaches the vertical fracture tip and recirculates along the upper boundary, distributing high velocities throughout the fracture profile. An increase in  $Re_{in}$  leads to a stronger degree of downwards flow and recirculation.

Along the central jet region, all experiments show an increase in velocity with height up to the normalized location of the velocity maximum,  $\hat{z}_{max}$ . Above  $\hat{z}_{max}$ , the velocity decreases with height as it approaches  $u_{tip}$  [Fig. 3(a),  $t^* = 0.5$ ]. All  $Re_{in} < 1$  experiments have a particularly steep velocity increase up to  $\hat{z}_{max}$ , which decreases from approximately 0.4 to 0.1 for experiments S1–S5, respectively, thus indicating the decrease in the jet height with increasing  $Re_{in}$  for  $Re_{in} < 1$ . As the  $Re_{in} > 1$  experiments have a longer jet than the  $Re_{in} < 1$  experiments, their velocity maximum is higher at  $\hat{z}_{max} \approx 0.5$  and the profiles are more or less symmetric about  $\hat{z}_{max}$  (for all  $Re_{in} > 1$  experiments).

The normalized horizontal velocity line profiles highlight the focused flow around the jet region—which has the same relative thickness for all experiments—and the local velocity minima on either side. They collapse onto one another in the central jet region at  $t^* = 0.5$  [Fig. 3(b)]. The secondary velocity peaks represent the downwards



**FIG. 4.** Characteristic fracture velocities (tip velocity  $u_{tip}$ , jet velocity  $u_{jet}$ , absolute downwards velocity  $|u_{down}|$ , mean velocity  $u_{mean}$ , mm/s) relative to the volumetric flux  $Q$  ( $\text{mm}^3/\text{s}$ ) at a dimensionless time of  $t^* = 0.5$ . Experiments with  $Re_{in} < 1$  are shown in purple, and  $Re_{in} > 1$  experiments are shown in green. The line of best fit between  $u_{tip}$  and  $Q$  is shown for all experiments ( $u_{tip} = 6.4 \times 10^{-4} Q$ ).

circulating flow, which is significantly stronger for  $Re_{in} > 1$ . While the central jet velocity dominates the flow in all experiments, W3 and W4 reach particularly high downwards velocities of around 60% of the maximum value.

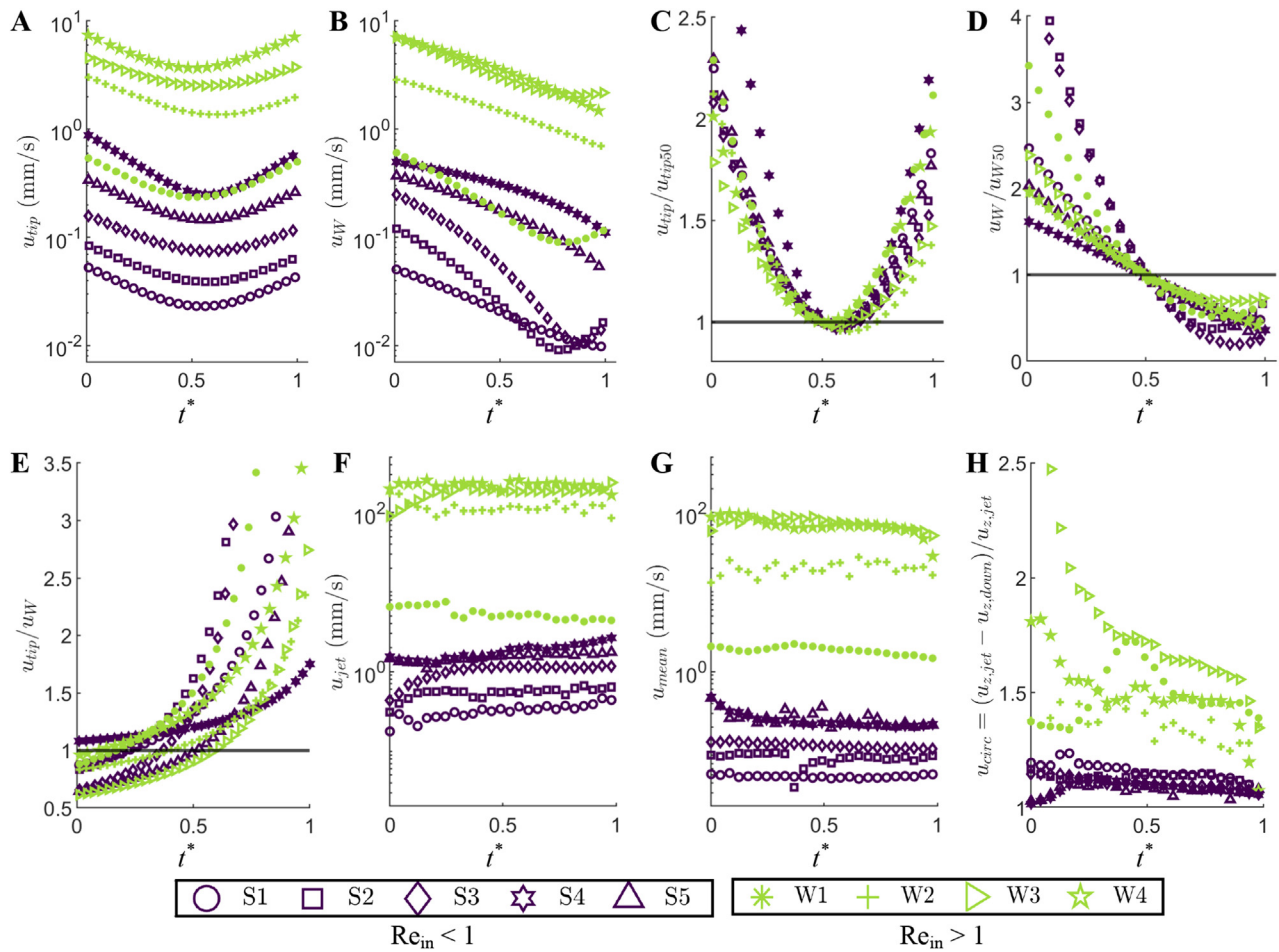
## B. Characteristic velocities

Characteristic flow velocities ( $u_{jet}$ ,  $u_{down}$ ,  $u_{mean}$ ) are nonlinear functions of  $Q$ , exhibiting a unique relationship for the two sets of experiments  $Re_{in} < 1$  and  $Re_{in} > 1$ . In both cases, internal velocities initially increase with  $Q$ , before appearing to reach a limiting value. This is in contrast to the vertical tip velocity  $u_{tip}$ , which is a linear function of  $Q$  (Fig. 4,  $t^* = 0.5$ ). The approximate inlet velocity  $u_{in}$  does not increase linearly with  $Q$  for all experiments due to differences in the size of the injection needle. However, when comparing  $u_{in}$  and  $u_{tip}$ , it is clear that  $u_{tip}$  is less than 1% of  $u_{in}$  at any given time (see Table II, and Fig. 2 in the supplementary material). Although  $u_{tip}$  is a linear function of  $Q$  overall, S4 has greater tip velocities than S5 at all times [Figs. 5(a) and 5(b)], despite having a lower  $Q$ . S4 contained a trapped air bubble at the fracture tip, which we interpret to have enhanced its overall propagation rate.

When  $Re_{in} < 1$ ,  $u_{mean}$  and  $|u_{down}|$  have similar values to  $u_{tip}$ , which all lie in the range 0.02 – 0.25 mm/s (Fig. 4 and Table II). The simple velocity approximation  $Q/WH$  is also very similar to  $u_{tip}$ ,  $u_{mean}$  and  $u_{down}$ . When  $Re_{in} > 1$ ,  $u_{mean}$  and  $u_{down}$  are significantly closer in value to  $u_{jet}$  than  $u_{tip}$ . The simple velocity approximation  $Q/WH$  is larger than  $u_{tip}$  when  $Re_{in} > 1$ , but still significantly under-predicts the mean internal flow.

## C. Temporal evolution of characteristic velocities

Overall, all experiments exhibit a similar pattern in terms of the temporal behavior of different characteristic velocities (Fig. 5). The nine experiments have a widely dispersed range of fracture tip velocities in the vertical ( $u_{tip}$ ) and horizontal ( $u_w$ ) directions [Figs. 5(a) and 5(b)], however, when normalized they show consistent behavior. In all cases,  $u_{tip}$  initially decreases, then reaches a short-lived state of steady propagation around  $t^* = 0.5$  [Fig. 5(c)] and finally increases to eruption. In contrast,  $u_w$  decreases rapidly until near-eruption when it approaches a steady value [Fig. 5(d)]. In the initial stage of



**FIG. 5.** Evolution of characteristic fracture velocities (purple = silicon oil, green = water) with normalized time  $t^*$ : (a) Fracture tip velocity in the vertical direction,  $u_{tip}$ , (b) velocity of crack breadth increase (horizontal tip velocity,  $u_W$ ), (c) normalized  $u_{tip}$  (according to  $u_{tip50}$ , the mean velocity at  $t^* = 0.5$ ), (d) normalized  $u_W$  (according to  $u_{W50}$ , the mean horizontal tip velocity at  $t^* = 0.5$ ), (e) ratio of vertical to horizontal tip propagation velocities ( $u_{tip}/u_W$ ), (f) internal jet flow velocities (from PIV)  $u_{jet}$ , (g) mean absolute velocities (from PIV)  $u_{mean}$ , and (h) flow circulation velocities (from PIV)  $u_{circ} = (u_{jet} - u_{down})/u_{jet}$ , where  $u_{circ} = 2$  indicates strong circulation.

propagation,  $u_{tip}/u_W < 1$  for all experiments except S4 [Fig. 5(e)]. The two velocities then approach one another (at a different  $t^*$  value for each experiment), after which  $u_{tip}$  significantly exceeds  $u_W$  up until eruption.

Compared to temporal variations in tip velocities, temporal variations in internal velocities are generally insignificant [Figs. 5(f) and 5(g), Table II]. Depending on the experiment,  $u_{mean}$  and  $u_{jet}$  are either approximately constant in time [e.g., S1 and S2, where  $\partial u_{jet}/\partial t = O(10^{-7})$ ] or vary slowly in time (e.g., S4 where  $\partial u_{jet}/\partial t \approx 0.0032 \text{ mm/s}^2$ ). Accelerations were approximated as the gradient of the linear curve fitted to the temporal velocity data plotted in Figs. 5(f) and 5(g) (reported in Table II).

Experiments with  $Re_{in} > 1$  (W experiments) exhibit a strong initial degree of circulation ( $u_{circ} \approx 2$ , and  $u_{circ} > 2$ ) that decreases in intensity over time [Fig. 5(h)]. In the early stages of experiment W3,  $u_{circ} > 2$  and the downwards flow is faster than the jet flow. When  $Re_{in} < 1$  (S experiments),  $u_{circ}$  is small for the entire duration of the

experiment (with a maximum value of 1.1–1.2 in the early stages that very gradually decreases over time).

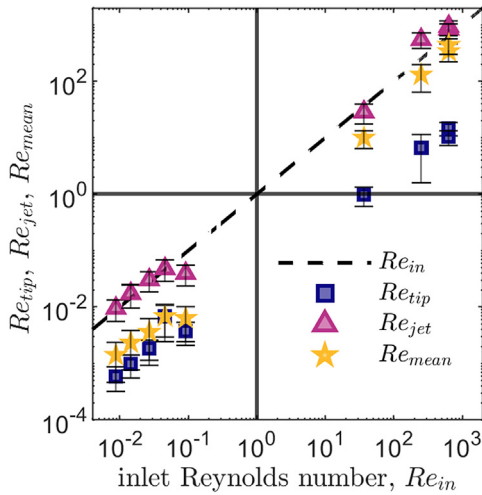
#### D. Force balance during fracture propagation

$Re_{jet}$  is consistently higher than the alternative  $Re$  definitions (Table III), as  $u_{jet}$  is the largest characteristic velocity (Table II).  $Re_{jet}$  is most similar to  $Re_{in}$ , despite  $u_{in}$  being significantly greater than  $u_{jet}$  (Table II). In all experiments,  $Re_{tip}$  is one order of magnitude smaller than  $Re_{jet}$  and  $Re_{in}$ , reflecting the small  $u_{tip}$  values compared to  $u_{jet}$  (Fig. 4). When  $Re_{in} < 1$ ,  $Re_{mean}$  is very similar to  $Re_{tip}$ . Conversely, when  $Re_{in} > 1$ ,  $Re_{mean}$  is the same order of magnitude as  $Re_{jet}$  and  $Re_{in}$ . When  $Re_{in} < 1$ ,  $Re_{tip}$ ,  $Re_{jet}$ , and  $Re_{mean}$  reach limiting values (all  $< 0.05$ ) with increasing  $Re_{in}$ . When  $Re_{in} > 1$ ,  $Re_{tip}$ ,  $Re_{jet}$ , and  $Re_{mean}$  do not (yet) reach a limiting value with increasing  $Re_{in}$ .

Variations of the mean viscous and inertial forces [ $|\overline{F}_V|$  and  $|\overline{F}_I|$ ], see Eq. (7) with respect to  $Re_{in}$  are shown in Fig. 7. As expected,

**TABLE III.** Characteristic Reynolds numbers for each experiment:  $Re_{in}$ ,  $Re_0$ ,  $Re_{tip}$ ,  $Re_{mean}$ , and  $Re_{jet}$ .

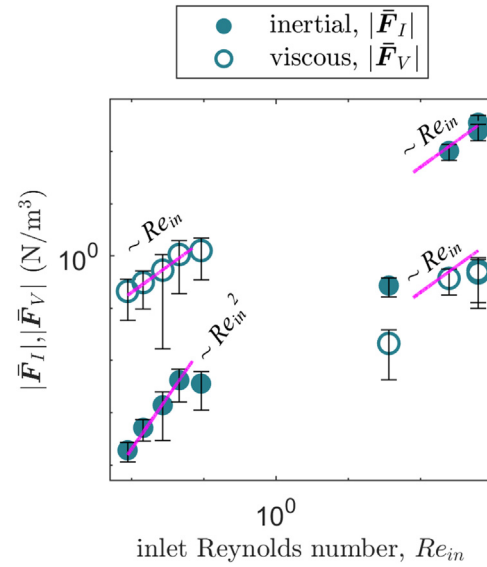
	$Re_{in}$	$Re_0$	$Re_{tip}$	$Re_{mean}$	$Re_{jet}$
S1	0.009	$6.05 \times 10^{-4}$	$5.91 \times 10^{-4}$	0.001	0.010
S2	0.015	$9.70 \times 10^{-4}$	$9.93 \times 10^{-4}$	0.002	0.017
S3	0.027	0.002	0.002	0.004	0.030
S4	0.046	0.004	0.007	0.007	0.048
S5	0.092	0.006	0.004	0.006	0.039
W1	36.65	2.401	0.973	9.964	28.680
W2	250.088	19.991	6.541	131.081	549.470
W3	633.005	38.205	10.284	439.670	845.773
W4	633.005	46.563	17.713	339.39	941.492



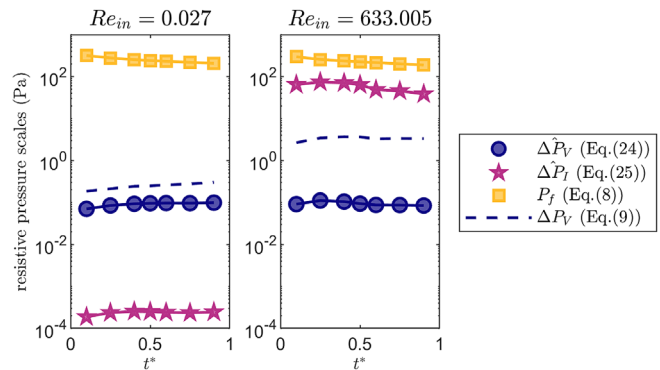
**FIG. 6.** The relationship between alternative flow Reynolds number [ $Re_{tip}$  (blue squares),  $Re_{jet}$  (pink triangles) and  $Re_{mean}$  (yellow stars)] and the inlet Reynolds number,  $Re_{in}$ . The dashed line depicts  $Re_{in}$ , and the solid lines show  $Re = 1$ . All calculations are made using velocity measurements at  $t^* = 0.5$ . Error bars are shown, which incorporate the error from velocity and  $H$  measurements.

viscous forces dominate over inertial when  $Re_{in} < 1$  ( $|\overline{F}_V| > |\overline{F}_I|$ ), and conversely, inertial forces dominate over viscous when  $Re_{in} > 1$  ( $|\overline{F}_I| > |\overline{F}_V|$ ). When  $Re_{in} \ll 1$ ,  $|\overline{F}_V|$  scales linearly with increasing  $Re_{in}$ , while  $|\overline{F}_I|$  scales with  $Re_{in}^2$ . At the highest  $Re_{in}$  value less than one (experiment S5),  $|\overline{F}_V|$  and  $|\overline{F}_I|$  deviate from this pattern. Naturally, for  $Re_{in} \approx 1$ , the two forces are expected to be of similar magnitude. Therefore, in the transitional region between  $Re_{in} < 1$  and  $Re_{in} > 1$ , curves of  $|\overline{F}_V|$  and  $|\overline{F}_I|$  will crossover— $|\overline{F}_V|$  decreases, while  $|\overline{F}_I|$  continues to increase. However, there are not enough data points in this region to determine how the forces behave during the transition. When  $Re_{in} \gg 1$ ,  $|\overline{F}_V|$  and  $|\overline{F}_I|$  scale linearly with  $Re_{in}$ . At the lowest  $Re_{in}$  value greater than one (experiment W1),  $|\overline{F}_I|$  and  $|\overline{F}_V|$  deviate from the linear scaling law.

The fracture pressure  $P_f$  is the largest resistive pressure in all experiments, which decreases with time (Fig. 8, showing S3 and W4 as representative low and high  $Re_{in}$  experiments, respectively). The viscous pressure drop  $\Delta P_V$  [approximated numerically, Eq. (24)] has a



**FIG. 7.** Mean fluid forces [inertial  $|\overline{F}_I|$  (filled circles) and viscous  $|\overline{F}_V|$  (empty circles)] as a function of  $Re_{in}$ , at a dimensionless time of  $t^* = 0.5$ . The error bars represent one standard deviation from the mean (across the full 2D fracture profile). Lines indicate the power law scaling (linear or quadratic) of the forces with respect to  $Re_{in}$ .



**FIG. 8.** Resistive pressure scales against dimensionless time  $t^*$ , for representative experiments S3 ( $Re_{in} = 0.027$ ) and W4 ( $Re_{in} = 633.005$ ). Numerical approximations of viscous  $\Delta P_V$  and inertial  $\Delta P_I$  pressure drops are depicted by blue circles and pink stars respectively. The fracture  $P_f$  is shown by yellow squares, and the blue dashed line represents the viscous pressure scale derived from lubrication theory ( $\Delta P_V$ ).

similar magnitude for S3 and W4, despite these experiments having very different  $Re_{in}$  values. For S3 ( $Re_{in} = 0.027$ ), the lubrication theory approximation of the viscous pressure drop [ $\Delta P_V$ , Eq. (9)] is similar to the numerical profile, and both increase with  $t^*$  (and  $L$ ). For W4 ( $Re_{in} = 633.005$ ),  $\Delta P_V$  is considerably larger than  $\Delta \hat{P}_V$ , which remains approximately constant over time. The inertial pressure drop  $\Delta \hat{P}_I$  [Eq. (25)] is negligible for experiment S3, yet it is of the same order of magnitude as  $P_f$  in W4. Before  $t^* = 0.5$ ,  $\Delta \hat{P}_I$  increases slightly in W4 before decreasing after this (at a faster rate than  $P_f$ ).

V. DISCUSSION

A. Self-similar flow in flux-driven fractures: Central jet and recirculating zones

Our experiments show that a central, localized jet and recirculating flow are consistent features of Newtonian, flux-driven fractures for a wide range of inlet Reynolds numbers ( $0.009 \leq Re_{in} \leq 633$ ) and internal flow velocities ( $0.3 \leq u_{jet} \leq 235$  mm/s). Similar flux-driven fracture experiments to ours (injecting a Newtonian fluid into gelatin) have been shown to exhibit the same flow structure, with a narrow range of internal flow velocities ( $u_{jet} \approx 5 - 10$  mm/s).<sup>32,40,41</sup> This characteristic flow pattern also occurs in different jet flow problems, as first shown in the pioneering experiments of Zauner<sup>65</sup> where fluid was injected into a tank filled with the same fluid. The resultant jets increased in thickness with height due to entrainment from the outer flow (which we also observe). For low  $Re$  ( $Re \approx 10$ ), the jet terminated at a finite distance from the injector and transitioned into regions of recirculatory flow (also known as viscous toroidal eddies<sup>65</sup>) Using asymptotic analysis on jets with  $Re > 1$ , Schneider<sup>66</sup> showed that momentum flux decays with increasing distance from the injector, primarily due to convection at the interface between the jet and the outer flow (i.e., momentum within the jet is transferred to the outer flow). This analysis suggested that viscous stresses do not contribute to the momentum flux decay, and showed that jet termination and recirculatory flow is induced when the momentum flux becomes very small. Further examples of where this flow pattern occurs are inside a balloon being inflated with air,<sup>67,68</sup> and in the “stable, recirculatory flow” stage of cavity formation in a porous soil due to an increasing flow rate.<sup>69,70</sup> Here, we explain why our experiments exhibit this characteristic flow pattern.

In flux-driven fractures, fluid is injected at a higher rate than the fracture can propagate (Table IV). The resultant flow is a complex coupling of a jet flow and a solid-fluid boundary problem, where viscous effects are fundamental. Viscous forces are proportional to velocity gradients [Eq. (7)] so that viscous effects are always important in shear layers, even when viscosity is negligible in the main flow.<sup>71</sup> Shear layers comprise localized regions of rotating fluid elements, aka vorticity  $\omega = \nabla \times \mathbf{u}$ . For a 2D flow,  $\omega = (0, 0, \omega)$  has one non-zero component,

$$\omega = \frac{\partial u_z}{\partial x} - \frac{\partial u_x}{\partial z}. \tag{30}$$

Inertial forces convect vorticity toward shearing boundary layers, whereas viscous forces act to diffuse vorticity away from these boundaries. Diffusive viscous flow at boundary layers controls the dynamics in the main flow.<sup>71</sup>

In our experiments, measured velocity profiles show that there are two regions of high shear where the fluid velocity rapidly changes value: at the interface between the jet and the main flow, and the no-slip boundary (Figs. 2 and 3). Localized vorticity is created at the jet margins and convected with the flow, where the degree of convection depends on  $Re_{in}$  (Fig. 9). Combined with the integral no-slip condition at the solid fracture boundary,<sup>66</sup> viscous diffusion of vorticity from the main jet leads to a recirculating vortex on either side of it. The relative strengths of convection and diffusion of vorticity vary with  $Re_{in}$  and lead to variations of the characteristic flow pattern (see Sec. V C). In summary, the characteristic jet and recirculating flow pattern is controlled by viscous shear layers, and we propose that this is a consistent feature of Newtonian flux-driven fractures. We expect that this flow

TABLE IV. Characteristic parameters of flux-driven fractures in the current experiments and magmatic and glacial settings: Young’s modulus  $E$  (Pa), fracture toughness  $K_C$  (pa m<sup>3/2</sup>), Poisson’s ratio  $\nu$  (dimensionless), volumetric flux  $Q$  (m<sup>3</sup>/s), viscosity  $\mu$  (Pa s), fluid density  $\rho_f$  (kg/m<sup>3</sup>), fracture thickness  $H$  (m), fracture width  $W$  (m), Reynolds number  $Re_0 = \frac{Q\rho_f}{W\mu}$ , dimensionless viscosity  $\mu_k = \frac{12\mu Q}{E} (\frac{E}{K})^4$ , and dimensionless inertia  $R_k = \frac{\rho E^{6/3} Q^{6/3}}{K^{6/3} \mu^{6/3}}$ .  $R_k$  was calculated at  $t^* = 0.5$  for the experiments, and not estimated in nature. References for natural values: (a) Heap *et al.*,<sup>80</sup> (b) Vaughan,<sup>79</sup> (c) Balme *et al.*,<sup>85</sup> (d) Fischer *et al.*,<sup>86</sup> (e) Rist *et al.*,<sup>87</sup> (f) Traversa *et al.*,<sup>84</sup> (g) Das *et al.*,<sup>15</sup> (h) Fountain *et al.*,<sup>73</sup> (i) Wada *et al.*,<sup>82</sup> (j) Roman *et al.*,<sup>83</sup> (k) values for water, (l) typical basalt value,<sup>89</sup> (m) Rubin,<sup>3</sup> (n) Holdsworth *et al.*,<sup>88</sup> (o) range in experimental values at  $t^* = 0.5$ , and (p) assuming a range of sizes.

	Experiments	Magma (basalt)	Ice
$E$	2278–4337	$1 \times 10^9 - 10^{10}$ <sup>a</sup>	$10^8 - 10^{10}$ <sup>b</sup>
$K_C$	66–93	$1.4 \times 10^6 - 3.8 \times 10^6$ <sup>c</sup>	$10^5 - 20^5$ <sup>d,e</sup>
$\nu$	0.5	0.1–0.5 <sup>a</sup>	0.3 <sup>b</sup>
$Q$	$4.4 \times 10^{-8} - 7.0 \times 10^{-6}$	$1 - 10^3$ <sup>f</sup>	$10^{-5} - 10^3$ <sup>g,h</sup>
$\mu$	0.001–0.45	$10^1 - 10^{3i,j}$	$10^{-3k}$
$\rho_f$	998, 1040	2700 <sup>l</sup>	998 <sup>k</sup>
$H$	0.003–0.015	$0.1 - 10^{i,m}$	$10^{-3} - 1^n$
$W$	0.14 – 0.18 <sup>o</sup>	$10^1 - 10^{4p}$	$1 - 10^{4p}$
$Re_0$	$6 \times 10^{-4} - 5 \times 10^1$	$10^{-4} - 10^4$	$10^{-4} - 10^8$
$\mu_k$	$5 \times 10^{-7} - 3 \times 10^{-4}$	$10^{-5} - 10^{10}$	$10^{-8} - 10^{10}$
$R_k$	$4 \times 10^{-10} - 9 \times 10^{-6}$	...	...

pattern is unique to flux-driven Newtonian fractures: buoyancy-driven fractures can achieve greater tip velocities and exhibit unidirectional flow profiles,<sup>41</sup> and non-Newtonian fluids have a shear rate-dependent viscosity that would likely result in markedly different flow patterns.

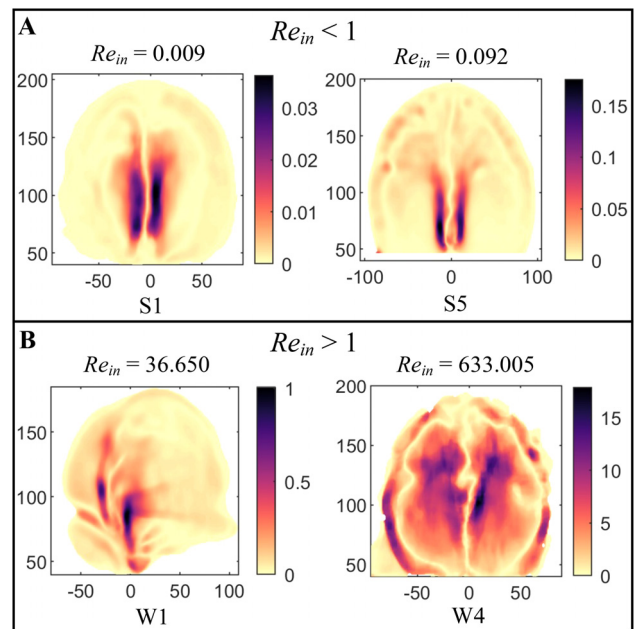
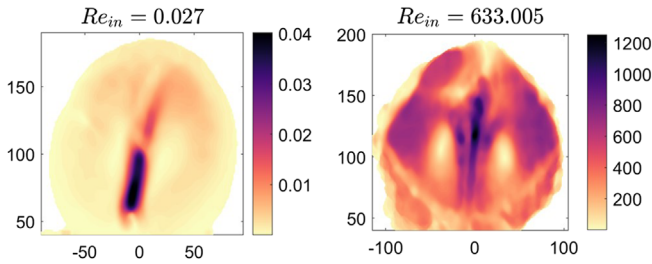


FIG. 9. Filled contours of vorticity (1/s) at a normalized time of  $t^* = 0.5$  for (a) low  $Re_{in}$  experiments S1 and S5 and (b) high  $Re_{in}$  experiments W1 and W4. Velocity gradients were calculated according to the finite difference method.<sup>54</sup>



**FIG. 10.** Filled contours of local flow Reynolds numbers for experiments S3 ( $Re_{in} = 0.027$ ) and W3 ( $Re_{in} = 633.005$ ), at a normalized time of  $t^* = 0.5$ . Local  $Re$  values were obtained at each measurement point throughout the 2D fracture profile, using the local velocity magnitude  $|\mathbf{u}|$  and constant values of  $\rho_f$ ,  $\mu$ , and  $H$ .

**B. Characteristic Reynolds numbers of fracture flow**

Fluid velocity in a flux-driven fracture is strongly spatially variable, which challenges the meaning of assigning a single Reynolds number to characterize fracture flow. The Reynolds number varies locally—this is highlighted by the range in alternative  $Re$  values for a single experiment (Fig. 6) and also by profiles of local Reynolds numbers. Figure 10 exhibits filled contours of local  $Re$  for experiments S3 ( $Re_{in} < 1$ ) and W3 ( $Re_{in} > 1$ ), showing the range of different flow regimes that can arise within a single fracture at a snapshot in time. In experiment S3, local  $Re$  reaches 0.04 in the jet, yet is of the order  $10^{-3}$  throughout the majority of the profile (outside the jet). This is reflected in the alternative characteristic  $Re$  definitions (for S3), where  $Re_{jet} = 0.03$  is significantly higher than  $Re_{mean} = 0.004$  (Table III). Conversely, W4 has high local  $Re$  values throughout the fracture profile, and  $Re_{jet} \approx 845$  is much closer in value to  $Re_{mean} \approx 440$ . Although the  $Re$  distribution is relatively uniform in W4, local  $Re$  contours also show spatial variation. Regions of both  $Re > 1000$  (in the jet) and  $Re < 1000$  potentially indicate simultaneous turbulent and laminar regimes.

Although  $Re$  varies locally (Fig. 10), it remains useful to characterize internal fracture flow with a single Reynolds number estimate.  $Re_{mean} = \rho_f H u_{mean} / \mu$  characterizes the overall flow well for both  $Re_{in} < 1$  and  $Re_{in} > 1$ . When  $Re_{in} < 1$ , the high-velocity (and high  $Re$ ) jet

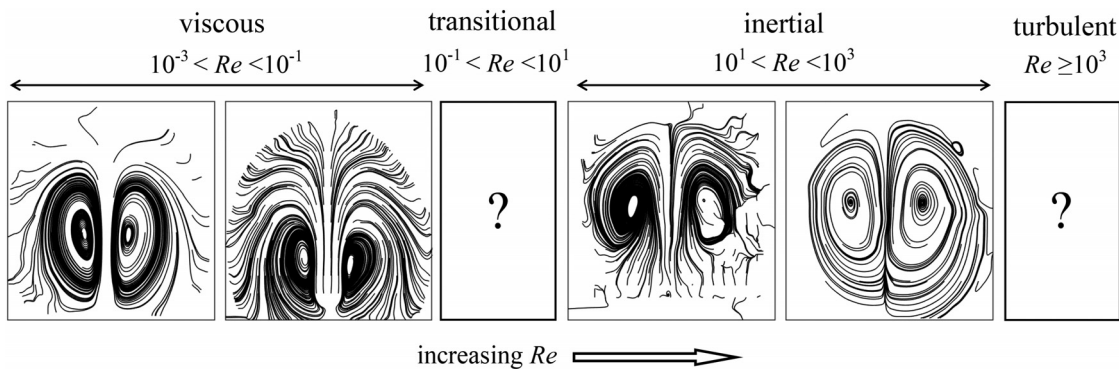
region is concentrated to a relatively small area. Momentum flux near the fluid inlet is rapidly dissipated, and  $Re_{mean}$  is significantly smaller than  $Re_{in}$  and  $Re_{jet}$ .  $Re_{mean}$  is in fact approximately equal to  $Re_{tip}$  and  $Re_0$  (Fig. 6 and Table III). When  $Re_{in} > 1$ , momentum flux is distributed throughout the fracture, and  $Re_{mean}$  is of a similar value to  $Re_{in}$  and  $Re_{jet}$ . In practical applications,  $u_{tip}$  and  $u_{in}$  can be measured<sup>15,72,73</sup> whereas internal fluid velocities  $u_{mean}$  and  $u_{jet}$  cannot. Therefore, we propose that  $Re_{tip}$  or  $Re_0$  provides an appropriate characteristic Reynolds number for slow, viscosity-dominated fractures while  $Re_{in}$  is more appropriate for fractures with important inertial effects. Note that calculating  $Re_{in}$  also requires knowledge of the area of the injection source—this is straightforward in analog experiments, but not necessarily in nature.

**C. Flow regimes in flux-driven fractures**

We propose that flux-driven fracture flow can be split into four regimes (Fig. 11) according to the inlet Reynolds number: viscous ( $Re_{in} < 10^{-1}$ ), transitional ( $10^{-1} \leq Re_{in} \leq 10^1$ ), inertial ( $10^1 < Re_{in} < 10^3$ ), and turbulent ( $Re_{in} \geq 10^3$ ). These regimes have been identified based on internal flow patterns and the behavior of average fluid forces across the  $Re_{in}$  range. While fracture flow is characterized by a localized jet and recirculation, our experiments show that this pattern can vary significantly within the range  $0.009 \leq Re_{in} \leq 633$ . However, our experiments do not cover the range  $0.1 \leq Re_{in} \leq 37$ —which is where we suggest a transitional regime between viscous and inertia dominated flow exists. The dynamics of flux-driven fracture flow in transitional and turbulent regimes have proved challenging to explore experimentally and therefore should be the subject of future research.

**1. Viscous regime**

In the viscous regime ( $Re_{in} < 0.1$ ), the jet always terminates before reaching the vertical fracture tip. Increasing  $Re_{in}$  leads to shorter jets with higher velocities near the inlet [Figs. 2 and 3(a)] and a higher magnitude of vorticity over a smaller region (Fig. 9). Viscous diffusion of vorticity also reduces, and the recirculatory zones become smaller yet more intense (i.e., contain higher velocities and vorticity) with increasing  $Re_{in}$  (Fig. 11). A decrease in jet height with increasing  $Re$  is



**FIG. 11.** Flow streamlines showing the different patterns that occur for different flow regimes (at dimensionless time  $t^* = 0.5$ ), based on the inlet Reynolds number. From left to right, these streamlines represent experiments S1, S5, W2, and W4. Each fracture shape has been normalized by its maximum length and width. Streamlines were generated using the Matlab<sup>59</sup> streamlines function.

the opposite of what occurs for unconfined jets in a fluid tank with no upper boundary,<sup>65</sup> for  $Re > 1$ . Those jets decrease in height with decreasing  $Re$ .

Viscous forces are greater than inertial forces and scale linearly with  $Re_{in}$  initially, while inertial forces increase at a faster rate and scale with  $Re_{in}^2$  (Fig. 7). This is expected from simple order of magnitude estimates of the fluid force terms [see Eqs. (11) and (12)], where  $|\overline{F}_V|$  scales with characteristic velocity  $\mathcal{U}$  and  $|\overline{F}_I|$  scales with  $\mathcal{U}^2$  (note that the Reynolds number represents a velocity scale). Higher inertial forces lead to higher velocities and vorticity in the jet, yet the simultaneous increase in viscous forces inhibits the jet from increasing in length. As  $Re_{in}$  approaches unity, the increase in both  $|\overline{F}_I|$  and  $|\overline{F}_V|$  slows down, indicating that the flow is near the onset of the transitional regime.

## 2. Inertial regime

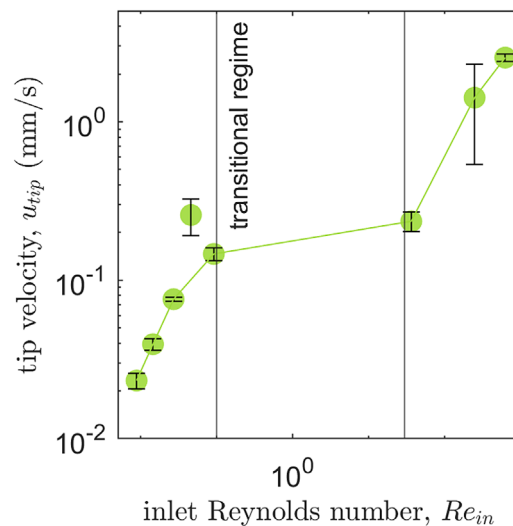
For  $10^1 < Re_{in} < 10^3$ , the jets do not terminate prior to reaching the upper solid boundary [Fig. 3(a)]. Vorticity is convected with the jet flow and along the fracture margins (Fig. 9). For the highest  $Re_{in} = 633$  (experiments W3 and W4), convection of vorticity toward the solid boundary dominates over viscous diffusion away from it, and a layer of high vorticity is confined to the entire fracture boundary. This leads to flow circulation throughout the fracture. With decreasing  $Re_{in}$ , vorticity is convected some distance along the upper fracture boundary before the flow loses momentum and vorticity is diffused from the boundary into the main flow. This results in vortices that are located closer to the upper fracture tip (as opposed to  $Re_{in} < 1$  where the vortices are located near the injector Fig. 11). Based on numerical simulations of an air-inflated balloon, it is expected that at higher  $Re$ , each vortex will split into multiple smaller scale vortices.<sup>68</sup> However, this pattern would be altered significantly by turbulent flow.

Scaling arguments [see Eqs. (11) and (12)] suggest that  $|\overline{F}_I|$  should increase with the square of the velocity (and therefore the Reynolds number). While this is observed within the viscous regime, when  $Re_{in} > 1$  inertial forces increase linearly with  $Re_{in}$  (Fig. 7). Unlike in the viscous regime where velocities vary smoothly over larger length scales (Fig. 3), in the inertial regime, flow is strongly spatially variable with finer-scale flow structures. Therefore, an average approximation of the inertial force may not be fully representative. Our results suggest that as  $Re_{in}$  decreases and approaches one,  $|\overline{F}_I|$  and  $|\overline{F}_V|$  deviate from their scaling laws. This potentially indicates that the flow is near the transitional regime, at the higher end of the  $Re_{in}$  range.

## 3. Transitional regime

Our experiments do not span the transitional range  $0.1 \leq Re_{in} \leq 36$ , which is challenging to achieve experimentally. When injecting silicone oil ( $Re_{in} < 0.1$ ), we reached the maximum possible  $Re_{in} (= 0.092)$  that could be achieved experimentally. For the water experiments, it was not possible to inject fluid at a lower rate than presented here, without potential settling of the tracer particles. Therefore, it is currently unclear how internal fracture flow behaves in the transition from viscous to inertial flow. However, the flow behavior in the viscous and inertial regimes gives some insight into what occurs during the transition.

Fracture tip velocities are very similar for experiments that lie at the transitional margins, which suggests that the transitional regime has a narrow range in tip velocities (Fig. 12). This region of



**FIG. 12.** Vertical fracture tip velocity  $u_{tip}$  (mm/s) against  $Re_{in}$  at a dimensionless time of  $t^* = 0.5$ . The line connects all experiments except S4, which is an outlier due to a trapped bubble and additional buoyancy effects.

approximately constant tip velocity coincides with a shift in the behavior of the fluid forces (Fig. 7). Unlike in the viscous and inertial regimes  $|\overline{F}_V|$  and  $|\overline{F}_I|$  do not appear to be a simple function of  $Re_{in}$ . During the transition from viscous to inertial flow,  $|\overline{F}_V|$  must decrease from the high values in the  $Re_{in} < 1$  experiments to the lowest values in the  $Re_{in} > 1$  experiments. Similarly,  $|\overline{F}_I|$  must increase across the transition, although Fig. 7 indicates that this increase is non-linear and  $|\overline{F}_I|$  potentially plateaus before increasing again. However, there are not enough data points to determine how the fluid forces evolve across the transitional regime. When  $Re \approx 1$ , inertial and viscous forces become similar in magnitude and are of equal importance. More experiments are needed to understand how fluid forces evolve across the transitional regime and determine the exact  $Re$  values at which the transition occurs. Future experiments could use tracer particles with a lower density or a Newtonian fluid with a higher density and viscosity than water (yet less viscous than silicone oil).

## 4. Turbulent regime

The onset of turbulence in fractures is commonly assumed to occur at  $Re \approx 1000$ .<sup>1,3,74</sup> Experiments W3 and W4 have regions of flow with both local  $Re > 1000$  and  $Re < 1000$  (Fig. 10), potentially indicating simultaneous laminar and turbulent regimes. However, the uniform structure of the jet (Fig. 4 in the supplementary material) and consistent, non-chaotic flow behavior suggest that fracture flow is not turbulent in our experiments.<sup>75</sup> We injected fluid at the fastest rate achievable with our injection equipment, yet could not achieve turbulence when injecting water. Note that it could be possible to achieve turbulence by injecting liquids with a lower viscosity than water, yet these liquids also needs to satisfy the condition of having a similar density to gelatin and being able to hold tracer particles in suspension. An appropriate liquid may be challenging to identify, and this could be explored in future work. Thus, the dynamics of turbulent fracture flow remain an open question. Local  $Re$  contours (Fig. 10) suggest that the

central jet would be the first region of flow to become turbulent, but it is currently unclear whether the characteristic circulating flow structure would persist at higher Reynolds numbers. Note that 2D flow profiles do not reveal how the fluid is behaving in the third out-of-plane dimension—3D imaging is required to know if the flow is chaotic across the fracture thickness.

#### D. Controls on fracture propagation

Pressure scale estimates (Fig. 8) suggest that the initial deceleration in  $u_{tip}$  is due to an increase in either the viscous or inertial resistive pressure as the fracture grows, depending on  $Re_{in}$ . The resistance to fracture (characterized by  $P_f$ ) decreases with fracture length, causing the fracture to accelerate toward the free surface.<sup>76</sup> The horizontal fracture growth,  $u_w$ , consistently decelerates, which does not coincide with any pressure scale. However, this does coincide with a decrease in the circulation velocity over time [Fig. 5(h)], suggesting that a reduction in the downwards flow velocity leads to a reduction in  $u_w$ .

The tip velocity  $u_{tip}$  is a linear function of the flux  $Q$  (Fig. 4), yet for a given tip velocity, there is a wide range of potential fluid behavior within. Across the transitional regime, injection rates and tip velocities are very similar, yet  $Re_{in}$  ranges from approximately 0.1–30 (Fig. 12). This suggests that a constant proportion of the driving pressure (due to fluid injection) contributes toward fracture propagation, regardless of the internal fluid behavior. The remainder of this applied pressure is distributed via different combinations of the inertial and viscous fluid forces (Fig. 8), producing different internal flow patterns (Figs. 2 and 11). Although the fluid injection rate controls fracture propagation, we expect that the internal fluid dynamics have a subtle but potentially significant effect on the coupled solid host deformation. Further experiments focusing on solid displacement measurements are needed to investigate this.

Through a theoretical analysis of a 2D flux-driven fracture propagating in an infinite, elastic medium, Emerman *et al.*<sup>10</sup> found the tip velocity to be a linear function of the inlet velocity, with  $u_{tip} \approx 0.45u_{in}$ . This is a markedly different relationship to our experiments, where  $u_{tip}$  is consistently less than 1% of  $u_{in}$  (Fig. 2 in the supplementary material). This difference is likely due to the 2D plane-strain model assumption with an infinite fracture width, as opposed to the point source injection method in our experiments. In propagating, non-buoyant fractures, the fracture width is always expected to exceed the length scale of the fluid inlet due to radial fracture growth. However, the size of the injection area affects the inlet Reynolds number and will have a strong effect on flow dynamics. This was investigated in numerical simulations of the flow within an inflating balloon:<sup>68</sup> smaller inlets led to longer and more focused jets, yet jet formation and flow circulation always occurred. Future experimental work could investigate the effect of the fluid inlet area on flux-driven fracture dynamics.

#### E. Application to magmatic and glacial systems

Experimental, flux-driven fractures in gelatin are an idealized analog of natural, geophysical flux-driven fractures. The dimensionless viscosity  $\mu_k$  and inertia  $R_k$  are  $\ll 1$  (Table IV), confirming that the analog fractures propagate in the toughness regime. Although fracture toughness dominates overall, the ratio of viscous and inertial forces varies significantly across the experiments. We now consider the dynamic similarity between the analog experiments and natural glacial

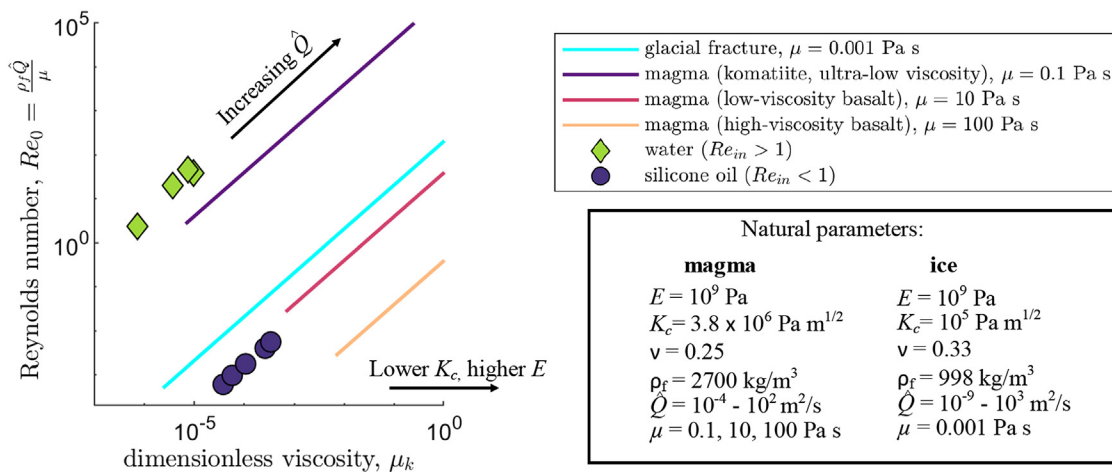
and magmatic systems by comparing the dimensionless parameter space defined by  $\mu_k$  and the Reynolds number. For the latter, we use the definition  $Re_0 = \rho_f \dot{Q} / \mu$ , where  $\dot{Q} = Q/W$  is the flux per unit width, in order to directly compare with nature. This requires appropriate estimates of magmatic and glacial parameters.

The applicability of the experiments to natural systems is limited by the model assumptions (e.g., elastic solid, Newtonian fluid). Glacial ice is not strictly linear elastic, but it is expected to behave in an elastic way under fracture.<sup>77</sup> In magmatic systems, the assumption of elasticity is only applicable to the lithosphere.<sup>78</sup> Appropriate  $E$  values range from  $10^8$  to  $10^{10}$  Pa for glaciers<sup>79</sup> and  $10^9$  to  $10^{10}$  Pa for the elastic crust.<sup>80</sup> Water in glacial crevasses is Newtonian ( $\mu \approx 10^{-3}$  Pa s), whereas the rheology of magma depends on the relative proportions of crystals, melt and bubbles.<sup>81</sup> Newtonian magmas are relatively crystal-poor with no bubbles, representing a primitive mafic (low-silica) magma:  $\mu$  ranges from  $10^1$  to  $10^2$  Pa s for basaltic magma,<sup>82,83</sup> yet can be as low as  $10^{-2}$  for ultramafic magmas such as komatiite or carbonatite.<sup>29</sup> Numerical models of basaltic dykes with a constant flux suggest that  $Q$  can range from 1 to  $1000 \text{ m}^3/\text{s}$ ,<sup>84</sup> whereas glacial fractures exhibit a wider range of  $Q$  values—from  $O(10^{-5}) \text{ m}^3/\text{s}$  in thin fracture networks,<sup>73</sup> to  $O(10^3) \text{ m}^3/\text{s}$  in rapid drainage events.<sup>15</sup> Fracture lengths, widths, and thicknesses have a wide variety of potential values, and we consider a range of fracture sizes. The natural parameter estimates are summarized in Table IV.

According to the dimensionless parameter space defined by  $\mu_k$  and  $Re_0$  (Fig. 13), our experiments represent the lowest end of the  $\mu_k$  spectrum for natural magmatic and glacial fractures. The natural parameter space is depicted as a series of linear lines, each representing a different fluid viscosity, and a wide range of  $\dot{Q}$  values. All other parameters are assumed to be constant: while  $\rho_f$  and  $\nu$  have little effect on the overall parameter space,  $E$  and  $K_C$  do act to shift the dimensionless viscosity range significantly. Here, we have selected values within the valid natural range that best represent the experiments in this parameter space (the upper and lower ends of the  $E$  and  $K_C$  ranges respectively, see Fig. 13 inset box). Both sets of experiments represent natural injections with a low flux (per unit width). Recall that the rate of injection is limited by the fluid viscosity (silicone oil could not be injected at a higher rate than achieved here). The silicone oil experiments ( $Re_{in} < 1$ ) are fairly well representative of glacial fractures and basaltic dykes, whereas the water experiments ( $Re_{in} > 1$ ) are more representative of a very low viscosity magma, such as a primitive komatiite. Future work could explore fractures with higher  $\mu_k$  and fill in the gaps in our understanding of flux-driven fluid dynamics across the natural parameter space. A higher  $\mu_k$  could be achieved by using other fluids with different viscosities and injecting them at a range of rates.

#### VI. CONCLUSIONS

Analog experiments of flux-driven fractures have shown that internal fracture flow has a self-similar pattern of a high-velocity central jet with a zone of fluid recirculation on either side, consistent across a range of regimes. We have utilized PIV velocity data to identify four potential regimes: viscous, inertial, transitional, and turbulent. Viscous and inertial regimes were produced experimentally (with some experiments perhaps bordering the transitional regime) for inlet Reynolds numbers spanning  $O(10^{-3}) \leq Re_{in} \leq O(10^2)$ . In the viscous regime, the jet and adjacent vortices shrink with increasing  $Re$  yet become more intensely localized near the jet. To our knowledge, this is the first experimental insight into the behavior of jets at  $Re < 1$ . In the



**FIG. 13.** Dimensionless parameter space for experiments and natural geophysical examples, defined by the dimensionless viscosity  $\mu_k$  and the Reynolds number  $Re_0 = \dot{Q}\rho_f/\mu$ . The water and silicone oil experiments are depicted by green diamonds and purple circles respectively. Each line represents the potential range of  $\mu_k$  and  $Re_0$  for a range of  $\dot{Q} = Q/W$  values, with all other parameters constant. The light blue line depicts ice fractures, while the pink and purple lines represent dykes with three different magma viscosities.

inertial regime, the jet length always exceeds the fracture length, and an increase in  $Re_{in}$  leads to a greater degree of flow circulation. Although data are lacking for the transitional regime ( $Re \approx 1$ ) due to experimental limitations, we propose that the average fluid forces have a complex relationship with  $Re_{in}$ , yet fractures propagate at similar tip velocities within this regime. Despite the complexity of the internal flow, the propagation velocity is a linear function of the flux  $Q$ . These results have important implications for interpreting natural data on propagating fractures and developing better numerical models to predict them. A key advantage of our experimental model is that the solid transparency allows for measurements of fracture and flow dynamics in real-time. Furthermore, the model scales appropriately with natural flux-driven fractures, as shown by the dimensionless parameter space defined by  $Re$  and the dimensionless viscosity  $\mu_k$ . However, there remains a knowledge gap regarding transitional and turbulent flow in fractures. Model simplifications also restrict our analysis to fractures in elastic solids injected by Newtonian fluids with a constant viscosity and density. These assumptions are most restrictive in the application to volcanology, where hot rocks can deform inelastically, and crystal and bubble content can lead to variations in magma viscosity and density. Future experiments (using different fluid and solid properties) are required to understand the complete range of flow regimes in flux-driven fractures, across the full natural parameter space. Experimental measurements in 3D would bring further advancement to our understanding of fracture dynamics.

**SUPPLEMENTARY MATERIAL**

See the supplementary material for details on the experimental methodology and additional results visualizations.

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**AUTHOR DECLARATIONS**

**Conflict of Interest**

The authors have no conflicts to disclose.

**Author Contributions**

**Caitlin M. Chalk:** Data curation (equal); Formal analysis (lead); Methodology (equal); Visualization (lead); Writing – original draft (lead); Writing – review & editing (supporting). **Janine L. Kavanagh:** Conceptualization (lead); Data curation (equal); Formal analysis (supporting); Funding acquisition (lead); Methodology (equal); Supervision (lead); Writing – review & editing (lead).

**DATA AVAILABILITY**

The data are available in the NGDC repository at <https://doi.org/10.5285/92383dba-8b6e-43da-9d78-e8784b97124c>, Ref. 60.



## NOMENCLATURE

$A$	Fluid inlet area
$c_g$	Gelatin concentration
$d$	Fluid inlet diameter
$E$	Young's modulus
$F_I$	Inertial force, $F_I = (F_{Ix}, F_{Iz})$
$\hat{F}_I$	Inertial force (numerical approximation)
$\overline{F}_I$	Average inertial force (numerical approximation)
$F_V$	Viscous force, $F_V = (F_{Vx}, F_{Vz})$
$\hat{F}_V$	Viscous force (numerical approximation)
$\overline{F}_V$	Average viscous force (numerical approximation)
$H$	Fracture thickness
$K_C$	Fracture toughness
$L$	Fracture length
$\mathbf{n}$	Unit normal
$P_F$	Fracture pressure scale
$p$	Dynamic pressure
$Q$	Volumetric flux
$\hat{Q}$	Flux per unit width
$Re_{in}$	Inlet Reynolds number
$Re_{jet}$	Jet Reynolds number
$Re_{mean}$	Mean Reynolds number
$Re_{tip}$	Tip Reynolds number
$Re_0$	Flux Reynolds number
$R_k$	Dimensionless inertia
$r_c$	Correlation value
$St$	Stokes number
$t$	Time, $t_0 \leq t \leq t_{max}$
$t_{erupt}$	Time between $t_0$ and eruption
$t_0$	Time when $L = 10$ cm
$t^*$	Dimensionless time, $t^* = (t - t_0)/t_{erupt}$
$U_g$	Stokes particle velocity
$\mathbf{u}$	Velocity vector, $\mathbf{u} = (u_x, u_z)$
$u_{circ}$	Circulation velocity
$u_{down}$	Representative downwards velocity
$u_{in}$	Inlet velocity
$u_{jet}$	Jet velocity
$u_{mean}$	Mean absolute velocity
$u_{tip}$	Vertical tip velocity
$u_{tip50}$	Vertical tip velocity when $t^* = 0.5$
$u_W$	Horizontal tip velocity
$u_{W50}$	Horizontal tip velocity when $t^* = 0.5$
$W$	Fracture width
$x$	Horizontal axis
$\hat{x}$	Normalized horizontal coordinate
$y$	Out of plane axis
$z$	Vertical axis
$\hat{z}$	Normalized vertical coordinate
$\hat{z}_{max}$	Normalized height of velocity maximum
$\Delta P_I$	Inertial pressure scale
$\hat{\Delta P}_I$	Inertial pressure scale (numerical approximation)
$\Delta P_V$	Viscous pressure scale
$\hat{\Delta P}_V$	Viscous pressure scale (numerical approximation)
$\Delta t$	Time increment
$\Delta x$	Grid spacing
$\partial/\partial t$	Partial time derivative
$\mu$	Dynamic viscosity

$\mu_k$	Dimensionless viscosity
$\nu$	Poisson's ratio
$\rho_f$	Fluid density
$\rho_s$	Solid density
$\tau_r$	Particle relaxation time
$\omega$	Vorticity magnitude
$\boldsymbol{\omega}$	Vorticity, $\boldsymbol{\omega} = \nabla \times \mathbf{u}$
$\nabla^2$	Laplace operator

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