

Up, down, and round again: the circulating flow dynamics of flux-driven fractures

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Fluid-filled fracture propagation is a complex problem that is ubiquitous in Geosciences, from controlling magma propagation beneath volcanoes to water transport in glaciers. Using scaled analogue experiments, we characterized the internal flow inside a propagating flux-driven fracture and determined the relationship between flow and fracture evolution. Different flow conditions were created by varying the viscosity and flux (Q) of a Newtonian fluid injected into an elastic solid. Using particle image velocimetry we measured the fluid velocity inside the propagating fracture and mapped the flow across the crack plane. We characterized the internal flow behavior with the Reynolds number (Re) and explored Re values spanning five orders of magnitude, representing very different internal force balances. The overall fracture tip propagation velocity is a simple linear function of Q , whereas the internal velocity, and Re , may be vastly different for a given Q . We identified four flow regimes – viscous, inertial, transitional, and turbulent – and produced viscous and inertial regimes experimentally. Both flow regimes exhibit a characteristic flow pattern of a high-velocity central jet that develops into two circulating vortices on either side. However, they exhibit the opposite behavior in response to changing Q : the jet length increases with Q in the inertial regime, yet decreases in the viscous regime. Spatially variable, circulating flow is vastly different from the common assumption of unidirectional fracture flow, and has strong implications for the mixing efficiency and heat transfer processes in volcanic and glacial applications.

1 I. INTRODUCTION

2 Fluid-filled fracture propagation is a fundamental process in many geoscience applications, in-
3 cluding magma transport¹⁻³, glacier dynamics and stability⁴⁻⁶, and geothermal energy systems⁷⁻⁹.
4 Magma-filled fractures (dykes and fissures) feed volcanic eruptions, whilst glacial fractures
5 (crevasses) control the drainage of glacial lakes and the transport of melt water. The fluid dy-
6 namics within propagating fractures has a significant effect on the overall fracture behavior.
7 Propagation is driven by internal fluid pressure (due to fluid injection, buoyancy, or a combination
8 of the two), which is distributed and dissipated by the internal flow¹⁰. In dykes and fissures, the
9 flow of magma influences the style of eruption at the surface^{11,12}. Flow in glacial crevasses can
10 have a significant impact on glacier stability and melting rates¹³⁻¹⁵. Understanding and predicting
11 fracture behaviour, including the expected pathway, propagation rate, and internal fluid dynamics,
12 is essential for managing the risks associated with volcanic and climate change processes. Despite
13 its importance, the fluid flow within propagating fractures is not well understood, and is typically
14 assumed to be unidirectional — a key assumption of many theoretical and numerical models of
15 fracture propagation^{2,16-18}.

16 A major challenge in modelling fracture propagation and internal fluid flow is having a unified
17 understanding of the full range of potential behaviour. Some theoretical and numerical models
18 models neglect fluid flow and assume that buoyancy dominates, and have been able to recreate
19 fracture pathways in experiments and in nature^{19,20}. However, buoyancy-driven fractures only
20 represent a subset of natural cases, and fluid flow must be included to obtain accurate predic-
21 tions of propagation velocities^{16,17}. Flux-driven fractures are driven by the pressure created by
22 fluid injection, where buoyancy may not play any role^{2,21,22}. In theoretical and numerical mod-
23 els of flux-driven fractures, flow, as characterised by the dimensionless Reynolds number Re , is
24 typically assumed to be in one of two limiting regimes: viscosity-dominated^{2,23-25} ($Re \ll 1$) or
25 turbulent²⁶⁻²⁸ ($Re > 1000$). In reality, fracture flow spans a wide range of regimes due to the
26 vast natural parameter space, notably the fluid viscosity. Fluid in glacial fractures and geothermal
27 systems has a viscosity of the order 10^{-3} Pa s, yet for magma this varies between 10^{-2} Pa s (for ul-
28 tramafic low-silica magmas) and 10^9 Pa s (for evolved, silica-rich rhyolite)²⁹. In nature, Reynolds
29 numbers range from the order of 10^{-10} for viscous, creeping dykes³⁰, to 10^6 and beyond for turbu-
30 lent crevasses during rapid drainage events⁵. There is therefore a strong motivation to understand
31 fracture propagation for the full range of potential flow regimes and Re values, particularly in the

32 transition from viscosity-dominated flows to full turbulence^{28,31}.

33 Scaled, analogue experiments of fluid-filled fracture propagation give crucial insight into the
34 fundamental processes of fracture dynamics (see Rivalta et al. (2015)¹⁷ and Kavanagh et al.
35 (2018)³² for a review). Laboratory experiments involving the injection of fluid into solid, elas-
36 tic, gelatine allow for direct observations of fracture and fluid dynamics during propagation^{33,34}.
37 Buoyancy-driven fracturing occurs if the injected fluid is sufficiently less dense than the solid
38 host^{35–38}. Otherwise, flux-driven fractures are created by the constant injection of fluid^{22,32,39}.
39 Recent studies have used Particle Image Velocity (PIV) to measure internal flow velocity pro-
40 files in flux-driven fractures^{32,40,41}. Whilst flow in buoyancy-dominated fractures is confirmed to
41 have a simple unidirectional profile⁴¹, Newtonian flux-driven fractures exhibit a more complex
42 flow pattern, consisting of a central, localized jet, with circulating downflow along the fracture
43 margins^{32,40,41}, which is not captured with any existing numerical model. Only a small number
44 of published experiments (all consisting of water injections with Reynolds numbers in the nar-
45 row range $1 \leq Re \leq 30$) have captured this interesting flow pattern. Experimental data across a
46 wider Re range is required to fully understand flux-driven fracture propagation and the influence
47 of internal fluid flow.

48 In this study, we provide the first experimental investigation of the dependence of fracture
49 dynamics on the Reynolds number. We restrict our attention to flux-driven fractures (that are not
50 buoyant) and conduct a series of experiments where a Newtonian fluid is injected into gelatine at
51 a constant rate. We systematically vary Re by changing the viscosity of the injected fluid and its
52 injection rate, achieving flows in the range $O(10^{-3}) \leq Re \leq O(10^2)$. Propagation velocities and
53 internal fluid flow profiles are measured across a two-dimensional plane of the growing fracture
54 for the full duration of the experiment. Our results showcase the complex fluid dynamics inside
55 flux-driven fractures and the relationship with propagation velocities. A jet and recirculation is a
56 universal feature of Newtonian flux-driven flows, yet there are key differences between viscous and
57 inertial flow regimes. We discuss the physics behind this observed behavior and the implications
58 for natural flux-driven fractures in glacial and magmatic settings.

59 II. THEORETICAL FRAMEWORK

60 In this section we present the relevant theoretical framework behind the experimental flux-
61 driven fractures, particularly related to the internal fluid flow (the main focus of this study). We

62 consider the following simplifications: a single, vertical fracture is driven by a constant continuous
 63 flux; the host is an isotropic, non-porous, elastic solid; the injected fluid is Newtonian and non-
 64 buoyant; fractures are tensile, opening in the direction of the least compressive stress σ_3 . We adopt
 65 Linear Elastic Fracture Mechanics (LEFM)^{3,42}, and assume that the solid resistance to fracture is
 66 characterized by the fracture toughness K_C (Pa/m^{1/2}). Fracture propagation occurs if the stress
 67 intensity K at the tip (a function of fluid pressure gradients) equals a critical value K_C . We assume
 68 that K_C is constant, although experiments suggest that K_C may change with fracture length³⁷.
 69 These simplifications are common assumptions in mathematical and numerical models of flux-
 70 driven fractures^{2,18,43,44}.

71 **A. Equations of fluid motion**

72 Flux-driven fracture propagation requires the continuous injection of fluid as this provides a
 73 driving pressure gradient at the source, which is distributed to the fracture tips via fluid flow. The
 74 evolution of pressure and fluid flow are governed by the Navier-Stokes equations, consisting of the
 75 conservation of mass and momentum:

$$76 \quad \nabla \cdot \mathbf{u} = 0, \quad (1)$$

$$77 \quad \rho_f \frac{\partial \mathbf{u}}{\partial t} + \rho_f \mathbf{u} \cdot \nabla \mathbf{u} = -\nabla p + \mu \nabla^2 \mathbf{u}. \quad (2)$$

77 Here $\mathbf{u} = (u_x, u_z)$ is the fluid velocity, ρ_f is the fluid density, p is the dynamic fluid pressure
 78 (i.e. excess of hydrostatic), t is time, and μ is dynamic viscosity. Along with suitable boundary
 79 conditions, Equations (1) and (2) describe incompressible, Newtonian flow inside a flux-driven
 80 fracture.

81 **1. Boundary conditions**

82 Injection of fluid can be expressed as a flux boundary condition:

$$83 \quad \mathbf{u} \cdot \mathbf{n} = \frac{Q}{A}, \text{ at the inlet}, \quad (3)$$

83 where \mathbf{n} is the unit normal direction to the inlet flow, Q is the volumetric flux (m³/s) and A (m²)
 84 is the area of the inlet surface (\mathcal{S}). $\frac{Q}{A}$ is the fluid injection velocity, also written as u_{in} . Note that
 85 Equation 3 is equivalent to imposing a pressure gradient at the inlet.

86 Fluid flow satisfies the no-slip condition:

$$\mathbf{u} = \mathbf{u}_{tip}, \text{ at the solid - fluid interface,} \quad (4)$$

87 where \mathbf{u}_{tip} is the velocity of the fracture tip.

88 **2. Fluid forces**

89 In the momentum equation (2), dynamic pressure gradients (units of force per unit volume,
90 N/m^3) are balanced with two forces, viscous ($\mathbf{F}_V = (F_{Vx}, F_{Vz})$) and inertial ($\mathbf{F}_I = (F_{Ix}, F_{Iz})$):

$$\nabla p = \mathbf{F}_I - \mathbf{F}_V \quad (5)$$

$$\mathbf{F}_I = \rho_f \frac{\partial \mathbf{u}}{\partial t} + \rho_f \mathbf{u} \cdot \nabla \mathbf{u}, \quad (6)$$

$$\mathbf{F}_V = \mu \nabla^2 \mathbf{u}. \quad (7)$$

91 In addition to driving fracture growth, the imposed pressure gradient due to fluid injection is also
92 dissipated by viscous and inertial forces.

93 **B. Opposing pressure scales**

94 The forces that oppose fracture propagation can be represented by simple pressure scales^{2,37,41,43,44}.
95 Solid resistance to fracture is represented by the fracture pressure scale P_f :

$$P_f \sim \frac{K_C}{\min(L, W)^{1/2}}, \quad (8)$$

96 where $\min(L, W)$ is the minimum of the fracture length L and width W .

97 The viscous pressure scale ΔP_V represents the drop in pressure along the fracture due to viscous
98 resistance:

$$\Delta P_V \sim \frac{3\mu u L}{H^2}, \quad (9)$$

99 where μ is viscosity, H is the fracture thickness and u is the average internal velocity. This scal-
100 ing (9) follows from the Navier-Stokes equations (1) (2) under lubrication theory assumptions:
101 laminar, unidirectional flow with negligible inertia^{1,2,18,25}.

102 In high Re flows inertial effects are important and pressure is dissipated via fluid kinetic
103 energy⁴⁵. The inertial pressure scale ΔP_I is derived through neglecting viscosity from the full
104 equations of fluid motion. ΔP_I represents the loss of fluid pressure to inertial forces:

$$\Delta P_I \sim \frac{f_D \rho_f u^2 L}{2}, \quad (10)$$

105 where f_D is a complex empirical function of the friction factor and fracture geometry^{5,46}. Closures
 106 for (10) have been proposed for turbulent fracture flow^{5,26}, yet there is very little focus on the
 107 transition from viscosity-dominated laminar flow and full turbulence⁴⁶.

108 C. Dimensionless numbers

109 Flux-driven fractures propagate in different regimes according to the dominant resistive pro-
 110 cesses. Fractures in gelatine are expected to propagate in the toughness regime^{32,37}, where the
 111 dominant opposing pressure scale is the fracture pressure P_F . In the toughness regime, the solid
 112 fracture process uses more energy than fluid forces, yet the relative balance of internal inertial and
 113 viscous forces still influences the overall fracture dynamics^{17,47}. Denoting the characteristic flow
 114 velocity with \mathcal{U} and the characteristic length scale with \mathcal{L} , the magnitude of the fluid force terms
 115 can be estimated as⁴⁸:

$$|\mathbf{F}_I| = |\rho_f \mathbf{u} \cdot \nabla \mathbf{u}| \sim \frac{\rho_f \mathcal{U}^2}{\mathcal{L}} \quad (11)$$

$$|\mathbf{F}_V| = |\mu \nabla^2 \mathbf{u}| \sim \frac{\mu \mathcal{U}}{\mathcal{L}^2}, \quad (12)$$

116 The Reynolds number Re represents the ratio of inertial to viscous forces:

$$Re = \frac{|\rho_f \mathbf{u} \cdot \nabla \mathbf{u}|}{|\mu \nabla^2 \mathbf{u}|} = \frac{\rho_f \mathcal{U} \mathcal{L}}{\mu}. \quad (13)$$

117 For fracture flows, \mathcal{L} is the fracture thickness H , and \mathcal{U} can be approximated as $\mathcal{U} \approx$
 118 Q/WH ^{41,49}. This reduces Re to:

$$Re_0 = \frac{\rho_f Q}{\mu W}. \quad (14)$$

119 In the toughness regime, two dimensionless numbers describe the relative effects of viscous
 120 and inertial forces to the fracture resistance. These are known as the dimensionless viscosity μ_k
 121 and dimensionless inertia R_k :

$$\mu_k = \frac{12\mu Q'}{E'} \left(\frac{E'}{K'} \right)^4 \quad (15)$$

$$R_k = \frac{\rho_f E'^{5/3} Q'^{5/3}}{K'^{8/3} t^{1/3}}, \quad (16)$$

122 where $Q' = Q/W$ is the volumetric flux per unit width (m²/s) and

$$E' = \frac{E}{1 - \nu^2}, \quad K' = 4\sqrt{2/\pi} K_C, \quad (17)$$

123 where E is the Young's modulus and ν is Poisson's ratio. Note that R_k is a decreasing function
124 of time, whilst μ_k is constant. R_k and μ_k were derived from an idealized, plane-strain, flux-driven
125 fracture model that quantifies the coupled effects of fracture resistance, viscosity and inertia²⁷. We
126 expect μ_k and R_k to be small (< 1) in the toughness regime as fracture resistance dominates.

127 **III. METHODOLOGY**

128 Here we describe the experimental process in detail. We first provide an overview of the ex-
129 perimental set-up and the materials used, followed by a description of the PIV method for mea-
130 suring internal fracture velocities during propagation. We then provide details on post-processing
131 the experimental data, including the calculation of representative fracture velocities, forces, and
132 Reynolds numbers.

133 **A. Overview**

134 A series of experiments were conducted to establish the effect of Re on both the internal fluid
135 dynamics and the overall propagation of flux-driven fractures. Each experiment consists of a New-
136 tonian fluid being injected into a $40 \times 40 \times 25$ cm³ volume of transparent, solid, elastic gelatine
137 held in a clear Perspex tank. An initial, vertical, pre-cut of 3 cm length and 1 cm width is
138 created in the centre of the base of the gelatine using a thin blade. The fluid is injected into the
139 pre-cut using a needle with its tapered edge orientated parallel to the widest part of the pre-cut.
140 The needle has an inlet diameter d of either 1 or 2 mm and an elliptical opening surface area A of
141 either $\pi \times 1 \times 3.5$ mm² or $\pi \times 2 \times 4$ mm². The needle is connected to a fluid reservoir via 5 mm
142 diameter tubing. A valve on the pipe and a small amount of petroleum jelly added to the end of
143 the needle ensures all air is removed from the injection system prior to starting an experiment. A
144 peristaltic pump then pushes fluid through the tube and into the gelatine at a known, constant rate,
145 creating a flux-driven penny-shaped fracture that propagates vertically and erupts at the surface
146 through a thin fissure. The pre-cut controls the fracture orientation, ensuring it grows vertically in
147 the z direction and radially in the $x - z$ plane, whilst pushing open the solid as a tensile fracture
148 in the $y - z$ plane (see Fig.1). Two-dimensional (2D) internal velocity profiles are measured in the
149 $x - z$ plane using a laser-based particle image velocimetry (PIV) system (see Sec. III C), controlled
150 via LaVision's DaVis 10 specialized laser imaging software^{50,51}. The laser-imaging system and

151 experimental tank are all supported by a robust, connected metal frame that ensures experiment
152 repeatability.

153 **B. Materials**

154 Different flow regimes were achieved by injecting Newtonian fluids with different viscosity
155 but similar density: a high viscosity fluid (silicone oil, $\mu = 0.45$ Pa s, $\rho_f = 998$ kg/m³), and a
156 low viscosity fluid (water, $\mu = 0.001$ Pa s, $\rho_f = 998$ kg/m³). The viscosity of silicone oil was
157 determined with a series of rheometer tests at different temperatures, and the density was obtained
158 using a 100 ml pycnometer.

159 The solid, elastic gelatine had a concentration of 2.5 wt % and was prepared following the
160 guidelines of Kavanagh et al.^{32,52}, resulting in 1001.5 kg/m³ solid density³². Gelatine preparation
161 involves mixing 1 kg of gelatine powder (260 Bloom, 10 mesh, pig-skin gelatine supplied by
162 Gelita UK) with 39 kg of deionised water, resulting in a total liquid mass of 40 kg. Approximately
163 half of the total amount of water was added hot (≈ 80 °C) to initially dissolve the gelatine, and
164 the rest was added cold ($\approx 7 - 10$ °C) directly in the tank. The liquid mixture was then covered
165 in a thin layer of vegetable oil and covered with plastic wrapping. It was left to cool and solidify
166 in a refrigerator for approximately 41-50 hours to obtain a Young's modulus E in the range 3000-
167 5000 Pa (note that the addition of cold water allowed for shorter solidification times than if using
168 hot water only). E was measured immediately before running an experiment using the method of
169 Kavanagh et al.⁵², which involves removing the surface oil and then applying different loads to the
170 centre of the gel surface and measuring their deflections. Gelatine's fracture toughness K_C can be
171 approximated as $K_C = 1.4\sqrt{E}$ ⁵².

172 **C. The PIV system**

173 Planar PIV is used to measure horizontal and vertical fluid velocities inside the fracture in
174 the $x - z$ plane. The injected fluid is pre-seeded with Rhodamine B-coated tracer particles with
175 diameters of $d_p = 20 - 50$ μ m and a particle density of $d_p = 1190$ kg/m³. Calculations of Stokes
176 settling velocity U_g , relaxation time τ_r , and Stokes number St suggest that these particles suitably
177 trace the fluid streamlines (see Sec. I in the Supplementary Material). Successive images of these
178 passive tracer particles are used to track fluid motion and compute velocities (see the videos in the

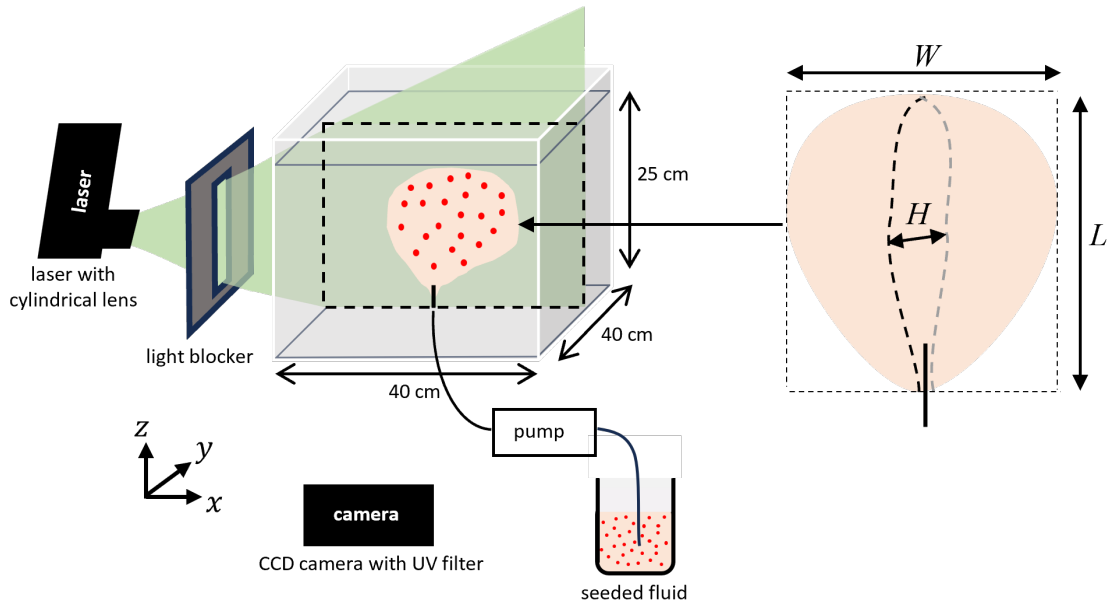


FIG. 1. Schematic of the experimental set-up depicting a growing flux-driven penny-shaped fracture being illuminated with a laser sheet.

179 Supplementary Material). The PIV method⁵³ divides each image into subdomains of a defined size
 180 (here either 32×32 or 24×24 pixels), and applies a statistical correlation technique to produce
 181 a single velocity vector per subdomain (with a calculation overlap of 75% between subdomains).
 182 The time interval between recorded images is chosen so that there is optimal particle displacement
 183 between successive images (of approximately 5 pixels⁵¹). PIV has been used to measure velocities
 184 in laboratory flows with a wide variety of geophysical applications^{32,54–58}.

185 The tracer particles are fluoresced with a sheet of light emitted from a Class 4 532nm Dou-
 186 blePulse Nd:YAG Litron laser (maximum energy 2x325 mJ), which illuminates the expected plane
 187 of fracture growth. Laser output is synchronized to an Imager SX 6M CCD camera facing the $x - z$
 188 plane, positioned perpendicular to the light sheet. The camera has a resolution of 2752×2200 pix-
 189 els and is used with a Zeiss 50mm f/1.4 lens with an aperture of f/5.6. The lens is fitted with a UV
 190 filter that blocks out short wavelengths and prevents reflections from the gelatine. Note that when
 191 measuring 2D flow, the light sheet should be as thin as possible to reduce out-of-plane motion
 192 effects. However, the fracture needs to remain within the sheet to ensure that the tracer particles
 193 are recorded (it is not guaranteed to be perfectly vertical). We used a 5mm-thick light sheet as a
 194 compromise between reducing potential three-dimensional (3D) effects and ensuring flow visual-
 195 ization. The laser is output through a -25mm convex cylindrical lens (from Edmund Optics Ltd).

196 Magnetic blockers (with an adjustable gap in the middle) positioned between the laser output and
197 tank wall transform the laser light into a thin sheet, orientated along the centre of the tank (see
198 Fig.1).

199 Images are captured in either single-frame or double-frame mode⁵³, depending on the expected
200 magnitude of particle displacements (Tab. I). In single-frame mode, each image consists of a
201 single frame that records the emission of two laser pulses with an exposure time of 42 μ s. The
202 shortest time interval that can be achieved between subsequent frames in single-frame mode is
203 restricted by the maximum camera frame rate of 15 frames per second. However, double-frame
204 recordings allow for shorter time intervals. In double-frame mode, each image is composed of
205 two frames separated by an interval Δt , which is achieved via control of the camera shutter. This
206 also defines the separation between the two laser pulses, so that the first frame captures the first
207 emission, and the second frame captures the second pulse. The exposure of the first frame is
208 defined by Δt , whereas the exposure of the second frame cannot be controlled and may capture
209 more ambient light. For this reason, double-frame recordings are performed with the overhead
210 room lights turned off. The recording settings for each experiment are provided in Tab. I in the
211 Supplementary Material.

212 Prior to running an experiment, the camera is first focused on fluoresced particles in a tank
213 full of seeded water, with the laser and camera in position. A calibration procedure is then per-
214 formed (within the Davis software) where images are taken of a calibration board positioned in
215 the imaging plane. Pixels are automatically converted to material coordinates in subsequent DaVis
216 operations. The calibration procedure is conducted prior to running an experiment, and requires
217 that the camera and laser positions are kept at a fixed position relative to the tank and imaging
218 plane – this is ensured via the supportive frame.

219 **D. Data processing**

220 *1. Post-processing PIV data*

221 Erroneous PIV velocity vectors were removed in ‘Vector Post-Processing’ in DaVis⁵⁰ accord-
222 ing to a threshold set by the correlation value r_c (the degree of confidence in the statistical corre-
223 lation procedure). A threshold of $r_c = 0.2$ eliminated vectors lying outside of the seeded fracture
224 flow.

225 Velocity data were exported from DaVis as a series of csv files (one for each time step), con-
226 taining the velocity components u_x, u_z and the corresponding spatial coordinates x, z . All further
227 analysis was performed in Matlab⁵⁹. Data were imported using the `readtable` function, and each
228 variable was converted to a 2D grid and processed with median filtering and Gaussian smoothing
229 functions (`medfilt2` and `smoothdata`). Velocity data collected in single-frame mode were time
230 averaged over an interval representing 5% of the experimental duration, resulting in an averaged
231 velocity profile and standard deviation. Double-frame velocity data were not time averaged due to
232 the time separation between two successive images being greater than 5%. All processing scripts
233 are available in an accompanying data publication⁶⁰.

234 **2. Tracking fracture geometry**

235 The fracture outline was extracted from the raw images by cropping around the illuminated
236 particles. Images were first converted to binary using `im2bw` (from the Image Processing
237 Toolbox⁶¹), and an appropriate pixel intensity threshold is selected to distinguish black from
238 white. The binary image was reduced in size and outliers removed, before applying the `rangesearch`
239 function (in the Statistics and Machine Learning Toolbox⁶²) to remove individual pixels
240 with fewer than a specified number of neighbours. The boundary function is applied to detect
241 the bounding shape of the reduced set of pixels, which is then converted to material coordinates.
242 An ellipse was fitted to the boundary points (with function `fitellipse`⁶³). Fracture length L and
243 width W were calculated from the length and width of the fitted ellipse at its centre point.

244 **3. Representative velocities**

245 Different representative velocities are used to characterize the fracture at a given time, includ-
246 ing tip velocities and internal flow velocities. Tip velocities consist of the fracture propagation rate
247 in the vertical and horizontal directions – u_{tip} and u_W respectively. Representative flow velocities
248 include the spatially-averaged mean velocity u_{mean} , a representative jet velocity u_{jet} , a representa-
249 tive downwards velocity u_{down} , and a circulation velocity⁴¹ u_{circ} . The latter represents the degree
250 of internal flow circulation: $u_{circ} = (u_{jet} - u_{down})/u_{jet}$. A value of $u_{circ} = 1$ means that there is
251 zero downwards flow and no circulation, whilst $u_{circ} = 2$ corresponds to the downwards velocity
252 being of equal magnitude to the upwards velocity, indicating strong circulation.

253 To obtain values for u_{jet} , vectors within the jet region were systematically cropped in each
 254 experiment and averaged in this area (within the 65-90 percentile range); a comparison with full
 255 velocity contours confirms that this method gives a velocity value that is representative of the
 256 jet. A similar method was applied to get u_{down} , where the data were instead cropped near the
 257 lateral fracture margins, and filtered according to $u_z < 0$. Tip velocities $u_{tip} = dL/dt$ and $u_W =$
 258 dW/dt were calculated by first fitting third-order polynomials to the temporal evolution of L and
 259 W , before integration over t . Error terms for u_{jet} , u_{mean} and u_{down} were calculated using the
 260 corresponding velocity standard deviation terms (note that these errors could be obtained for the
 261 non-time-averaged double-frame experiments W2-W4). Errors for u_{tip} and u_W were approximated
 262 by considering the difference between the temporal geometry data (L, W) and their fitted curves.

263 **4. Calculating fluid forces and pressure scales**

264 For each experiment, the viscous and inertial forces (7) can be calculated using the measured
 265 velocity data $\mathbf{u}_{i,j}$. Here, $\mathbf{u}_{i,j}$ denotes a single PIV grid measurement where indices i and j represent
 266 the spatial location in the x and z directions respectively. Adjacent grid points are separated by a
 267 constant distance Δx in both directions. Two adjacent points are denoted by $\mathbf{u}_{i,j}$ and $\mathbf{u}_{i+1,j}$ in the
 268 x axis, and $\mathbf{u}_{i,j}$ and $\mathbf{u}_{i,j+1}$ in the z axis. First and second order velocity derivatives are calculated
 269 using a finite difference method⁶⁴, and substituted into (7) to obtain approximations of inertial and
 270 viscous forces at a given time. These forces have a horizontal and vertical component (in 2D),
 271 such that $\mathbf{F}_v = (F_{Vx}, F_{Vz})$ and $\mathbf{F}_I = (F_{Ix}, F_{Iz})$. The numerical (finite difference) approximations of
 272 the force terms (denoted with a $\hat{\cdot}$ notation) are defined as:

$$\hat{F}_{Vxi,j} \approx \frac{\mu}{\Delta x^2} (u_{xi-1,j} + u_{xi+1,j} + u_{xi,j-1} + u_{xi,j+1} - 4u_{xi,j}), \quad (18)$$

$$\hat{F}_{Vzi,j} \approx \frac{\mu}{\Delta x^2} (u_{zi-1,j} + u_{zi+1,j} + u_{zi,j-1} + u_{zi,j+1} - 4u_{zi,j}), \quad (19)$$

$$\hat{F}_{Ixi,j} \approx \rho_f \frac{\partial u_{xi,j}}{\partial t} + \frac{\rho_f}{\Delta x} (u_{xi,j} (u_{xi+1,j} - u_{xi,j}) + u_{zi,j} (u_{xi,j+1} - u_{xi,j})), \quad (20)$$

$$\hat{F}_{Izi,j} \approx \rho_f \frac{\partial u_{zi,j}}{\partial t} + \frac{\rho_f}{\Delta x} (u_{xi,j} (u_{zi+1,j} - u_{zi,j}) + u_{zi,j} (u_{zi,j+1} - u_{zi,j})). \quad (21)$$

273 These terms are spatially-averaged across the full fracture profile to approximate the average

274 (absolute) forces (viscous \bar{F}_V and inertial \bar{F}_I) at a given time:

$$|\bar{F}_V| = \frac{1}{N-1} \sum_{i=2}^{N-1} |\hat{F}_{Vi,j}| \quad (22)$$

$$|\bar{F}_I| = \frac{1}{N-1} \sum_{i=2}^{N-1} |\hat{F}_{Ii,j}|, \quad (23)$$

275 where N is the total number of grid points and $|\hat{F}_{Vi,j}|$ and $|\hat{F}_{Ii,j}|$ denote the modulus of the numer-
276 ical viscous and inertial force terms at a single point.

277 Estimates of the viscous and inertial pressure scales are then obtained by multiplying the aver-
278 age force by the fracture length:

$$\Delta \hat{P}_V \approx |\bar{F}_V| L \quad (24)$$

$$\Delta \hat{P}_I \approx |\bar{F}_I| L. \quad (25)$$

279 E. Reynolds number calculations

280 The Reynolds number summarizes the bulk flow behavior as a single parameter, yet in reality
281 flow can be spatially and temporally variable, with a range of characteristic velocity and length
282 scales. In addition to $Re_0 = \frac{\rho_f Q}{\mu W}$ (14), we explore several alternative Reynolds numbers using
283 different characteristic velocities and length scales, that may potentially better represent the force
284 balance during fracture flow. We define four alternative Reynolds numbers: 1) the inlet Reynolds
285 number, Re_{in} , 2) the tip Reynolds number, Re_{tip} , 3) the jet Reynolds number, and 4) Re_{jet} , and the
286 mean Reynolds number Re_{mean} . The flow at the source of fluid injection is characterized by Re_{in} ,
287 and is known prior to running an experiment. Flow Reynolds numbers Re_{tip} , Re_{jet} and Re_{mean}
288 represent the internal flow during fracture propagation, and require measured velocity values (see
289 Sec. III D 3).

290 1. Inlet Reynolds number, Re_{in}

291 Re_{in} represents the fluid force balance at the inlet, and does not require any information about
292 the fracture flow or geometry. This is defined as:

$$Re_{in} = \frac{\rho_f u_{ind} d}{\mu}. \quad (26)$$

293 **2. Tip Reynolds number, Re_{tip}**

294 Re_{tip} uses the vertical fracture tip velocity u_{tip} as the characteristic velocity scale, and H as the
295 length scale:

$$Re_{tip} = \frac{\rho_f u_{tip} b}{\mu}. \quad (27)$$

296 **3. Jet Reynolds number, Re_{jet}**

297 Re_{jet} represents the internal flow behavior, with the characteristic velocity defined as the jet
298 velocity u_{jet} :

$$Re_{jet} = \frac{\rho_f u_{jet} b}{\mu}. \quad (28)$$

299 **4. Mean Reynolds number, Re_{mean}**

300 Re_{mean} also represents the internal flow, but instead uses the mean internal velocity u_{mean} as the
301 characteristic value:

$$Re_{mean} = \frac{\rho_f u_{mean} b}{\mu}. \quad (29)$$

302 **5. Fracture thickness measurements H**

303 The fracture thickness H is required to calculate the representative flow Reynolds numbers,
304 Re_0 , Re_{mean} , Re_{jet} , Re_{tip} . We conducted experiments to approximate H for low-viscosity, water
305 fractures and high-viscosity silicone oil fractures respectively. The fracture evolution was instead
306 recorded in the $y - z$ plane, either using a seeded fluid or a seeded gelatine (with no fluid seeding,
307 see Sec. III in the Supplementary Material). The representative thickness was approximated via
308 manual image inspection within the DaVis software.

309 **IV. RESULTS**

310 In total, eleven experiments were completed with Re_{in} ranging from 0.009 to 633 (Tab. I). All
311 experiments produced a broadly penny-shaped fluid-filled crack that grew and eventually erupted
312 at the surface. Nine experiments measured fluid velocities in the $x - z$ plane (five silicon oil (S)
313 and four water (W) injections) and two experiments measured a representative H for the different

314 fluid injections (one for silicone oil (SH), and one for water (WH)) (see Tab. II). All silicone oil
315 injections have $Re_{in} < 1$, whilst the water injections all have $Re_{in} > 1$. The fracture thickness H
316 varies with height, and measurements of H may be affected by optical distortions related to out-of-
317 plane fracture growth and mismatching refractive indices between the fluid and gelatine (see Sec. II
318 in the Supplementary Material). We therefore report a conservative range of H values: $H \approx 3 - 5$
319 mm for water, and $H \approx 7 - 15$ mm for silicone oil (S experiments have a high refractive index
320 mismatch and a larger error margin, as a closer refractive index matching leads to more accurate
321 measurements). These variations in H have a minor effect on the overall change in Reynolds
322 number, and the average H value was used in Re calculations.

323 In the following sections we present the fluid flow profiles, temporal fracture evolution, and
324 governing force balance results. Note that in these descriptions our main focus is the developed
325 flow pattern (the initial flow development is presented in Fig. 3 in the Supplementary Material).
326 To allow for a direct comparison between different experiments, time is normalized as $t^* = (t -$
327 $t_0)/t_{erupt}$, where t_0 is the time when $L = 10$ cm and t_{erupt} is the time interval between t_0 ($t^* = 0$)
328 and fluid eruption ($t^* = 1$).

329 **A. Flow profiles — central jet and recirculation**

330 Every experiment produced a high-velocity central jet that increased in width with height, be-
331 fore transitioning to a recirculating flow on either side of the jet, around two stagnant points (Fig. 2,
332 $t^* = 0.5$). There is a strong spatial velocity variation, with u_{jet} being the most dominant character-
333 istic flow velocity and ranging from 0.379 to 236 mm/s across all experiments (Tab. II, $t^* = 0.5$).
334 u_{jet} is at least one order of magnitude greater than the tip velocities in every experiment (which
335 range from 0.021 to 3.676 mm/s at $t^* = 0.5$).

336 Whilst all experiments have the same overall pattern of a jet and recirculating flow, there are
337 some clear differences between the $Re_{in} < 1$ and $Re_{in} > 1$ experiments (Figs. 2 and 3). In the
338 $Re_{in} < 1$ experiments (S1-S5), flow is mostly localized in the central jet region, and the downward
339 flow velocities are significantly lower than in the jet. In these experiments, the jet terminates
340 prior to reaching the vertical fracture tip, and an increase in Re_{in} correlates with a decrease in the
341 jet height. Conversely, in all $Re_{in} > 1$ experiments (W1-W4) the jet reaches the vertical fracture
342 tip and recirculates along the upper boundary, distributing high velocities throughout the fracture
343 profile. An increase in Re_{in} leads to a stronger degree of downwards flow and recirculation.

TABLE I. Solid and fluid material parameters for experiments S1-5 and SH (injecting silicone oil), experiments W1-W4 and WH (injecting water). Young’s modulus E (Pa), gelatine concentration c_g (wt%), volumetric flux Q (m³/s), fluid viscosity (μ), fluid density ρ_f (kg/m³), inlet diameter d (mm), thickness H (mm) (including range and associated experiment), inlet Reynolds number Re_{in} .

	E	Q	μ	ρ_f	d	H	Re_{in}
S1	4337	4.37×10^{-8}	0.450	998	1	11.5 (7-15, SH)	0.009
S2	4098	7.20×10^{-8}	0.450	998	1	11.5 (7-15, SH)	0.015
S3	4170	1.34×10^{-7}	0.450	998	1	11.5 (7-15, SH)	0.027
S4	4506	2.60×10^{-7}	0.450	998	2	11.5 (7-15, SH)	0.046
S5	4214	4.07×10^{-7}	0.450	998	1	11.5 (7-15, SH)	0.092
W1	4309	4.68×10^{-7}	0.001	998	1	4 (3-5, WH)	36.650
W2	2593	3.15×10^{-6}	0.001	998	2	4 (3-5, WH)	250.088
W3	3591	6.97×10^{-6}	0.001	998	1	4 (3-5, WH)	633.005
W4	2278	6.97×10^{-6}	0.001	998	1	4 (3-5, WH)	633.005
SH	3946	4.07×10^{-7}	0.450	998	1	11.5 (7-15)	0.092
WH	3591	6.97×10^{-6}	0.001	998	1	4 (3-5)	633.005

344 Along the central jet region, all experiments show an increase in velocity with height up to
 345 the normalized location of the velocity maximum, \hat{z}_{max} . Above \hat{z}_{max} , the velocity decreases with
 346 height as it approaches u_{tip} (Fig. 3A, $t^* = 0.5$). All $Re_{in} < 1$ experiments have a particularly steep
 347 velocity increase up to \hat{z}_{max} , which decreases from approximately 0.4 to 0.1 for experiments S1 to
 348 S5 respectively, thus indicating the decrease in the jet height with increasing Re_{in} for $Re_{in} < 1$. As
 349 the $Re_{in} > 1$ experiments have a longer jet than the $Re_{in} < 1$ experiments, their velocity maximum
 350 is higher at $\hat{z}_{max} \approx 0.5$ and the profiles are more or less symmetric about \hat{z}_{max} . (for all $Re_{in} > 1$
 351 experiments).

352 The normalized horizontal velocity line profiles highlight the focused flow around the jet region
 353 — which has the same relative thickness for all experiments — and the local velocity minima on
 354 either side. They collapse onto one another in the central jet region at $t^* = 0.5$ (Fig. 3B). The
 355 secondary velocity peaks represent the downwards circulating flow, which is significantly stronger
 356 for $Re_{in} > 1$. Whilst the central jet velocity dominates the flow in all experiments, W3 and W4

TABLE II. Flux-driven fracture thickness and characteristic velocity results for each experiment, at a dimensionless time of $t^* = 0.5$: inlet velocity u_{in} (mm/s), thickness H (mm) (including range and associated experiment), vertical tip velocity u_{tip} (mm/s), horizontal tip velocity u_W (mm/s), mean velocity u_{mean} (mm/s), jet velocity u_{jet} (mm/s), downflow velocity u_{down} (mm/s), $\frac{\partial u_{mean}}{\partial t}$ (mm/s²), $\frac{\partial u_{jet}}{\partial t}$ (mm/s²).

	u_{in}	u_{tip}	u_W	u_{mean}	u_{jet}	u_{down}	$\frac{\partial u_{mean}}{\partial t}$	$\frac{\partial u_{jet}}{\partial t}$
S1	3.975	0.023	0.021	0.050	0.379	-0.053	6.6×10^{-7}	4.4×10^{-5}
S2	6.551	0.040	0.023	0.084	0.686	-0.088	5.7×10^{-7}	5.3×10^{-5}
S3	12.209	0.076	0.054	0.124	1.229	-0.119	5.8×10^{-6}	3.3×10^{-4}
S4	10.347	0.258	0.308	0.247	1.925	-0.183	3.3×10^{-4}	0.0032
S5	41.430	0.147	0.182	0.233	1.591	-0.106	1.1×10^{-4}	6.6×10^{-4}
W1	36.724	0.236	0.171	2.432	6.907	-2.573	5.5×10^{-4}	-0.0042
W2	125.295	1.418	1.480	30.601	135.641	-58.565	0.052	-0.19
W3	634.273	2.539	2.929	108.404	211.867	-153.974	-0.60	1.36
W4	634.273	3.676	3.517	82.717	235.844	-124.827	-1.49	-0.80

reach particularly high downwards velocities of around 60% of the maximum value.

B. Characteristic velocities

Characteristic flow velocities ($u_{jet}, u_{down}, u_{mean}$) are nonlinear functions of Q , exhibiting a unique relationship for the two sets of experiments $Re_{in} < 1$ and $Re_{in} > 1$. In both cases, internal velocities initially increase with Q , before appearing to reach a limiting value. This is in contrast to the vertical tip velocity u_{tip} , which is a linear function of Q (Fig. 4, $t^* = 0.5$). The approximate inlet velocity u_{in} does not increase linearly with Q for all experiments due to differences in the size of the injection needle. However, when comparing u_{in} and u_{tip} , it's clear that u_{tip} is less than 1% of u_{in} at any given time (see Tab. II, and Fig. 2 in the Supplementary Material). Although u_{tip} is a linear function of Q overall, S4 has greater tip velocities than S5 at all times (Fig. 5A,B), despite having a lower Q . S4 contained a trapped air bubble at the fracture tip, which we interpret to have enhanced its overall propagation rate.

When $Re_{in} < 1$, u_{mean} and $|u_{down}|$ have similar values to u_{tip} , which all lie in the range 0.02 – 0.25 mm/s (Fig. 4 and Tab. II). The simple velocity approximation Q/WH is also very similar

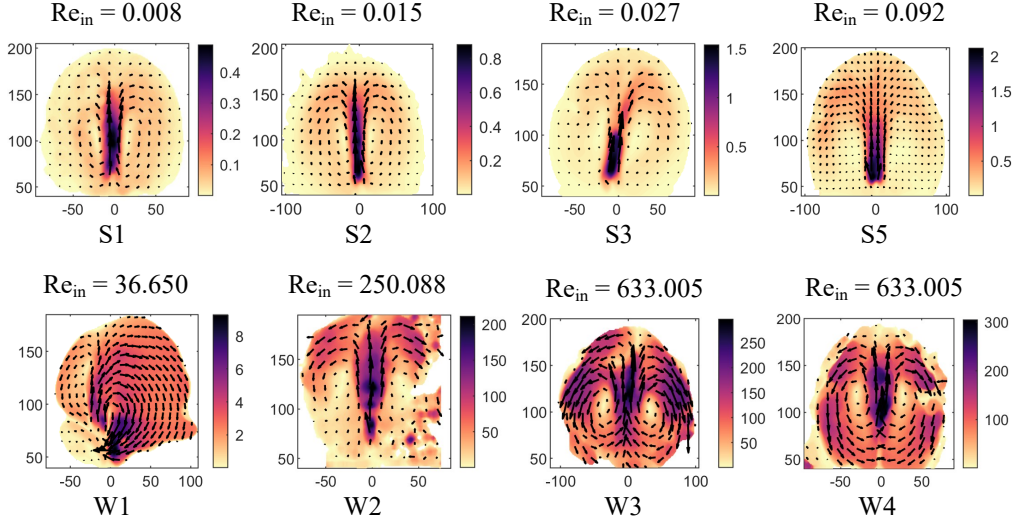


FIG. 2. Filled contours of velocity magnitude (mm/s) and vectors of flow direction (black arrows) for flux-driven fracture experiments, at a normalized time of $t^* = 0.5$. Four of the high-viscosity $Re_{in} < 1$ silicone oil experiments (S1-S3,S5) are shown on the top row, and the low-viscosity $Re_{in} > 1$ experiments (W1-W4) are on the bottom row. For each row, the experiments are ordered in terms of increasing Q and Re_{in} . The vectors show the flow direction, and their size represents the velocity magnitude, scaled up by a factor of two. Only every third vector is plotted (horizontally and vertically), whilst the filled contours show the full resolution of the flow velocity magnitude.

371 to u_{tip} , u_{mean} and u_{down} . When $Re_{in} > 1$, u_{mean} and u_{down} are significantly closer in value to u_{jet}
 372 than u_{tip} . The simple velocity approximation Q/WH is larger than u_{tip} when $Re_{in} > 1$, but still
 373 significantly under-predicts the mean internal flow.

374 C. Temporal evolution of characteristic velocities

375 Overall, all experiments exhibit a similar pattern in terms of the temporal behavior of different
 376 characteristic velocities (Fig. 5). The nine experiments have a widely dispersed range of fracture
 377 tip velocities in the vertical (u_{tip}) and horizontal (u_W) directions (Fig. 5A and 5B), however, when
 378 normalized they show consistent behavior. In all cases, u_{tip} initially decreases, then reaches a
 379 short-lived state of steady propagation around $t^* = 0.5$ (Fig. 5C) and finally increases to eruption.
 380 In contrast, u_W decreases rapidly until near-eruption when it approaches a steady value (Fig. 5D).
 381 In the initial stage of propagation, $u_{tip}/u_W < 1$ for all experiments except S4 (Fig. 5E). The two
 382 velocities then approach one another (at a different t^* value for each experiment), after which u_{tip}

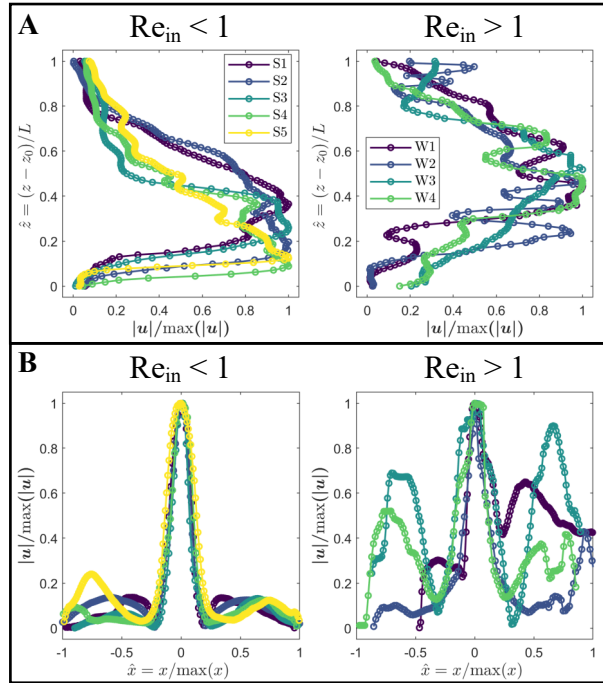


FIG. 3. Normalized velocity magnitude profiles at $t^* = 0.5$: (A) along the central vertical line above the injector and (B) along a horizontal line 60 mm above the injector, for silicon oil (S1-S5, $Re_{in} < 1$) and water (W1-W4, $Re_{in} > 1$) experiments. Height has been normalized so that $\hat{z} = 0$ is just above the injector, and $\hat{z} = 1$ is the tip location. In the vertical profiles, the pentagon denotes the fracture tip velocity. The horizontal distance has been normalized so that the $\hat{x} = 0$ corresponds to the location of the maximum velocity and the horizontal extent is between $\hat{x} = \pm 1$.

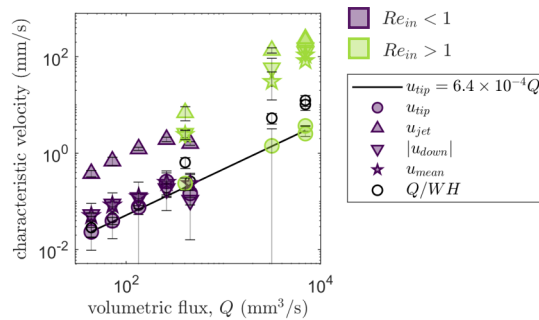


FIG. 4. Characteristic fracture velocities (tip velocity u_{tip} , jet velocity u_{jet} , absolute downwards velocity $|u_{down}|$, mean velocity u_{mean} , mm/s) relative to the volumetric flux Q (mm^3/s) at a dimensionless time of $t^* = 0.5$. Experiments with $Re_{in} < 1$ are shown in purple, and $Re_{in} > 1$ experiments are shown in green. The line of best fit between u_{tip} and Q is shown for all experiments ($u_{tip} = 6.4 \times 10^{-4} Q$).

Up, down, and round again: the circulating flow dynamics of flux-driven fractures

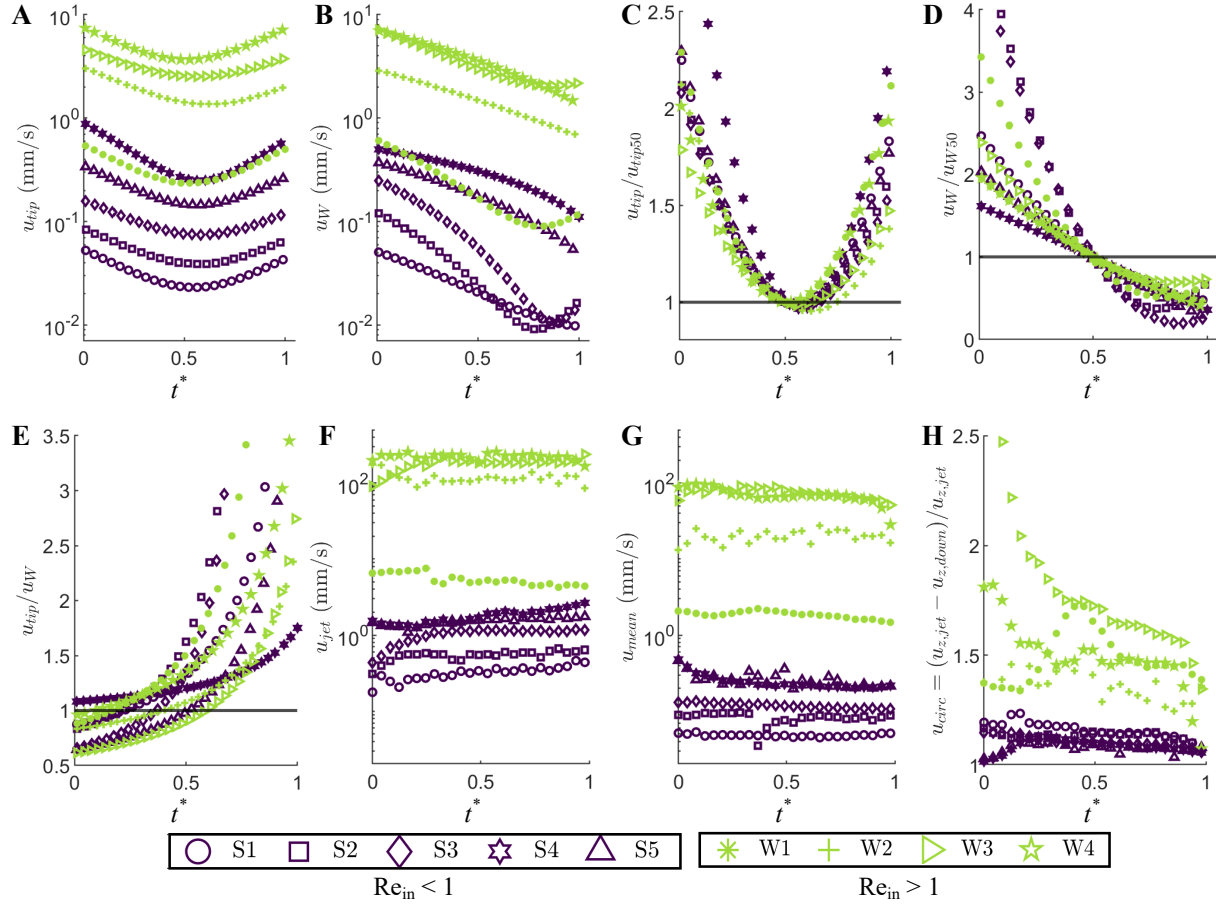


FIG. 5. Evolution of characteristic fracture velocities (purple = silicon oil, green = water) with normalized time t^* . A) Fracture tip velocity in the vertical direction, u_{tip} , B) Velocity of crack breadth increase (horizontal tip velocity, u_W), C) Normalized u_{tip} (according to u_{tip50} , the mean velocity at $t^* = 0.5$), D) Normalized u_W (according to u_{W50} , the mean horizontal tip velocity at $t^* = 0.5$), E) Ratio of vertical to horizontal tip propagation velocities (u_{tip}/u_W), F) Internal jet flow velocities (from PIV) u_{jet} , G) Mean absolute velocities (from PIV) u_{mean} , and H) Flow circulation velocities (from PIV) $u_{circ} = (u_{jet} - u_{down})/u_{jet}$, where $u_{circ} = 2$ indicates strong circulation.

383 significantly exceeds u_W up until eruption.

384 Compared to temporal variations in tip velocities, temporal variations in internal velocities
 385 are generally insignificant (Fig. 5F,G, Tab. II). Depending on the experiment, u_{mean} and u_{jet} are
 386 either approximately constant in time (e.g. S1 and S2 where $\partial u_{jet}/\partial t \approx O(10^{-7})$ and S2), or vary
 387 slowly in time (e.g. S4 where $\partial u_{jet}/\partial t \approx 0.0032$ mm²/s). Accelerations were approximated as
 388 the gradient of the linear curve fitted to the temporal velocity data plotted in Fig. 5F,G (reported in
 389 Tab. II).

TABLE III. Characteristic Reynolds numbers for each experiment: Re_{in} , Re_0 , Re_{tip} , Re_{mean} , Re_{jet}

	Re_{in}	Re_0	Re_{tip}	Re_{mean}	Re_{jet}
S1	0.009	6.05×10^{-4}	5.91×10^{-4}	0.001	0.010
S2	0.015	9.70×10^{-4}	9.93×10^{-4}	0.002	0.017
S3	0.027	0.002	0.002	0.004	0.030
S4	0.046	0.004	0.007	0.007	0.048
S5	0.092	0.006	0.004	0.006	0.039
W1	36.65	2.401	0.973	9.964	28.680
W2	250.088	19.991	6.541	131.081	549.470
W3	633.005	38.205	10.284	439.670	845.773
W4	633.005	46.563	17.713	339.39	941.492

390 Experiments with $Re_{in} > 1$ (W experiments) exhibit a strong initial degree of circulation ($u_{circ} \approx$
391 2, and $u_{circ} > 2$) that decreases in intensity over time (Fig. 5H). In the early stages of experiment
392 W3, $u_{circ} > 2$ and the downwards flow is faster than the jet flow. When $Re_{in} < 1$ (S experiments),
393 u_{circ} is small for the entire duration of the experiment (with a maximum value of 1.1 to 1.2 in the
394 early stages that very gradually decreases over time)

395 D. Force balance during fracture propagation

396 Re_{jet} is consistently higher than the alternative Re definitions (Tab. III), as u_{jet} is the largest
397 characteristic velocity (Tab. II). Re_{jet} is most similar to Re_{in} , despite u_{in} being significantly greater
398 than u_{jet} (Tab. II). In all experiments, Re_{tip} is one order of magnitude smaller than Re_{jet} and Re_{in} ,
399 reflecting the small u_{tip} values compared to u_{jet} (Fig. 4). When $Re_{in} < 1$, Re_{mean} is very similar to
400 Re_{tip} . Conversely, when $Re_{in} > 1$, Re_{mean} is the same order of magnitude as Re_{jet} and Re_{in} . When
401 $Re_{in} < 1$, Re_{tip} , Re_{jet} and Re_{mean} reach limiting values (all < 0.05) with increasing Re_{in} . When
402 $Re_{in} > 1$, Re_{tip} , Re_{jet} and Re_{mean} do not (yet) reach a limiting value with increasing Re_{in} .

403 Variations of the mean viscous and inertial forces ($|\bar{F}_V|$ and $|\bar{F}_I|$, see Equation (7)) with respect
404 to Re_{in} are shown in Fig. 7. As expected, viscous forces dominate over inertial when $Re_{in} < 1$
405 ($|\bar{F}_V| > |\bar{F}_I|$), and conversely, inertial forces dominate over viscous when $Re_{in} > 1$ ($|\bar{F}_I| > |\bar{F}_V|$).
406 When $Re_{in} \ll 1$, $|\bar{F}_V|$ scales linearly with increasing Re_{in} , whilst $|\bar{F}_I|$ scales with Re_{in}^2 . At the

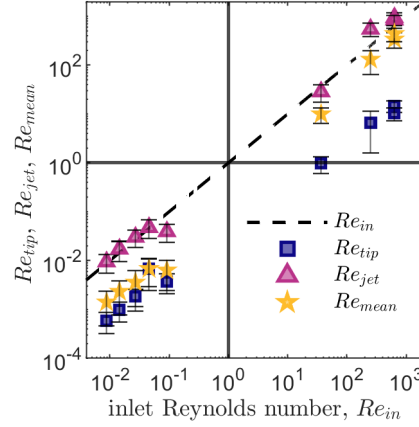


FIG. 6. The relationship between alternative flow Reynolds number (Re_{tip} (blue squares), Re_{jet} (pink triangles) and Re_{mean} (yellow stars)) and the inlet Reynolds number, Re_{in} . The dashed line depicts Re_{in} , and the solid lines show $Re = 1$. All calculations are made using velocity measurements at $t^* = 0.5$. Error bars are shown, which incorporate the error from velocity and H measurements.

407 highest Re_{in} value less than one (experiment S5), $|\bar{F}_V|$ and $|\bar{F}_I|$ deviate from this pattern. Naturally,
 408 for $Re_{in} \approx 1$, the two forces are expected to be of similar magnitude. Therefore, in the transitional
 409 region between $Re_{in} < 1$ and $Re_{in} > 1$, curves of $|\bar{F}_V|$ and $|\bar{F}_I|$ will cross over — $|\bar{F}_V|$ decreases,
 410 whilst $|\bar{F}_I|$ continues to increase. However, there are not enough data points in this region to
 411 determine how the forces behave during the transition. When $Re_{in} \gg 1$, $|\bar{F}_V|$ and $|\bar{F}_I|$ scale linearly
 412 with Re_{in} . At the lowest Re_{in} value greater than one (experiment W1), $|\bar{F}_I|$ and $|\bar{F}_V|$ deviate from
 413 the linear scaling law.

414 The fracture pressure P_f is the largest resistive pressure in all experiments, which decreases with
 415 time (Fig. 8, showing S3 and W4 as representative low and high Re_{in} experiments respectively).
 416 The viscous pressure drop $\Delta\hat{P}_V$ (approximated numerically, Equation (24)) has a similar magnitude
 417 for S3 and W4, despite these experiments having very different Re_{in} values. For S3 ($Re_{in} = 0.027$),
 418 the lubrication theory approximation of the viscous pressure drop (ΔP_V , Equation (9)) is similar
 419 to the numerical profile, and both increase with t^* (and L). For W4 ($Re_{in} = 633.005$), ΔP_V is con-
 420 siderably larger than $\Delta\hat{P}_V$, which remains approximately constant over time. The inertial pressure
 421 drop $\Delta\hat{P}_I$ (Equation (25)) is negligible for experiment S3, yet it is of the same order of magnitude
 422 as P_f in W4. Before $t^* = 0.5$, $\Delta\hat{P}_I$ increases slightly in W4 before decreasing after this (at a faster
 423 rate than P_f).

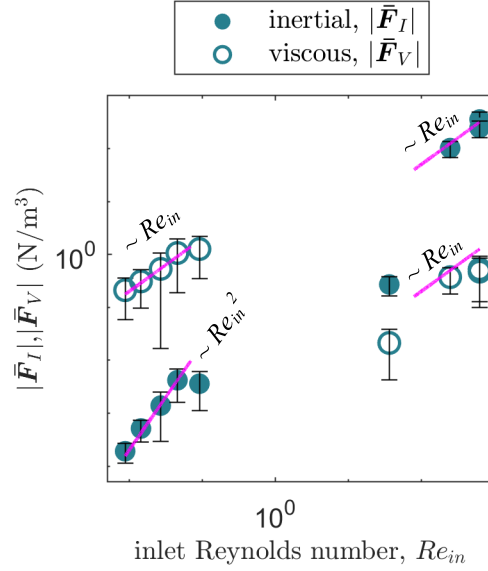


FIG. 7. Mean fluid forces (inertial $|\bar{\mathbf{F}}_I|$ (filled circles) and viscous $|\bar{\mathbf{F}}_V|$ (empty circles)) as a function of Re_{in} , at a dimensionless time of $t^* = 0.5$. The error bars represent one standard deviation from the mean (across the full 2D fracture profile). Lines indicate the power law scaling (linear or quadratic) of the forces with respect to Re_{in} .

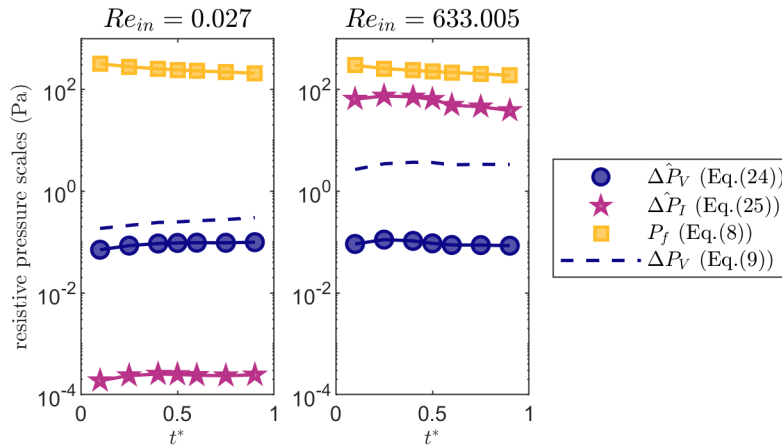


FIG. 8. Resistive pressure scales against dimensionless time t^* , for representative experiments S3 ($Re_{in} = 0.027$) and W4 ($Re_{in} = 633.005$). Numerical approximations of viscous $\Delta \hat{P}_V$ and inertial $\Delta \hat{P}_I$ pressure drops are depicted by blue circles and pink stars respectively. The fracture P_f is shown by yellow squares, and the blue dashed line represents the viscous pressure scale derived from lubrication theory (ΔP_V).

424 **V. DISCUSSION**425 **A. Self-similar flow in flux-driven fractures: central jet and recirculating zones**

426 Our experiments show that a central, localized jet and recirculating flow are consistent features
 427 of Newtonian, flux-driven fractures for a wide range of inlet Reynolds numbers ($0.009 \leq Re_{in} \leq$
 428 633) and internal flow velocities ($0.3 \leq u_{jet} \leq 235$ mm/s). Similar flux-driven fracture experi-
 429 ments to ours (injecting a Newtonian fluid into gelatine) have been shown to exhibit the same flow
 430 structure, with a narrow range of internal flow velocities ($u_{jet} \approx 5 - 10$ mm/s)^{32,40,41}. This char-
 431 acteristic flow pattern also occurs in different jet flow problems, as first shown in the pioneering
 432 experiments of Zauner (1985)⁶⁵ where fluid was injected into a tank filled with the same fluid. The
 433 resultant jets increased in thickness with height due to entrainment from the outer flow (which we
 434 also observe). For low Re ($Re \approx 10$), the jet terminated at a finite distance from the injector and
 435 transitioned into regions of re-circulatory flow (also known as viscous toroidal eddies⁶⁵). Using
 436 asymptotic analysis on jets with $Re > 1$, Schneider (1985)⁶⁶ showed that momentum flux decays
 437 with increasing distance from the injector, primarily due to convection at the interface between the
 438 jet and the outer flow (i.e. momentum within the jet is transferred to the outer flow). This analysis
 439 suggested that viscous stresses do not contribute to the momentum flux decay, and showed that
 440 jet termination and re-circulatory flow is induced when the momentum flux becomes very small.
 441 Further examples of where this flow pattern occurs are inside a balloon being inflated with air^{67,68},
 442 and in the ‘stable, recirculatory flow’ stage of cavity formation in a porous soil due to an increasing
 443 flow rate^{69,70}. Here, we explain why our experiments exhibit this characteristic flow pattern.

444 In flux-driven fractures, fluid is injected at a higher rate than the fracture can propagate
 445 (Tab. IV). The resultant flow is a complex coupling of a jet flow and a solid-fluid boundary
 446 problem, where viscous effects are fundamental. Viscous forces are proportional to velocity gra-
 447 dients ((7)), so that viscous effects are always important in shear layers, even when viscosity is
 448 negligible in the main flow⁷¹. Shear layers comprise localized regions of rotating fluid elements,
 449 aka vorticity $\boldsymbol{\omega} = \nabla \times \mathbf{u}$. For a 2D flow, $\boldsymbol{\omega} = (0, 0, \omega)$ has one non-zero component:

$$\omega = \frac{\partial u_z}{\partial x} - \frac{\partial u_x}{\partial z}. \quad (30)$$

450 Inertial forces convect vorticity towards shearing boundary layers, whereas viscous forces act to
 451 diffuse vorticity away from these boundaries. Diffusive viscous flow at boundary layers controls
 452 the dynamics in the main flow⁷¹.

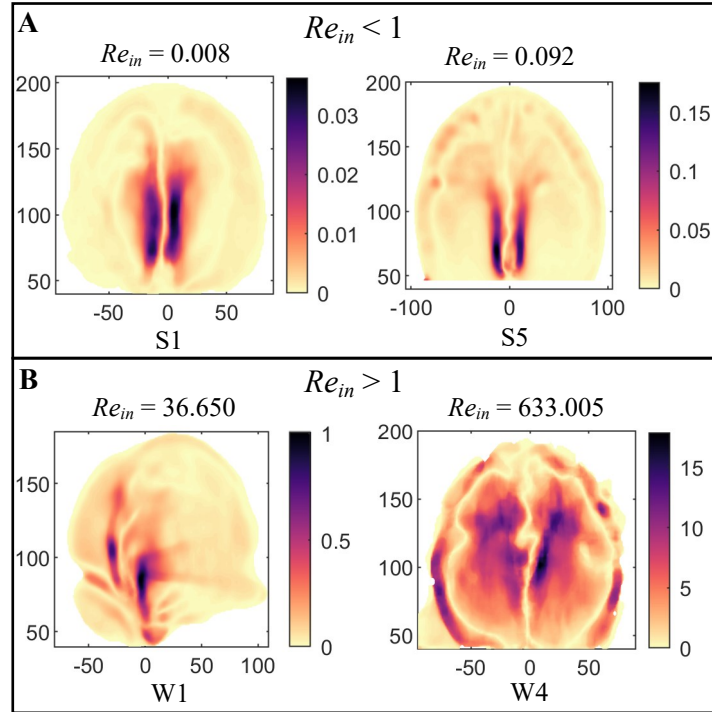


FIG. 9. Filled contours of vorticity (1/s) at a normalized time of $t^* = 0.5$ for A) low Re_{in} experiments S1 and S5 and B) high Re_{in} experiments W1 and W4. Velocity gradients were calculated according to the finite difference method⁶⁴.

453 In our experiments, measured velocity profiles show that there are two regions of high shear
 454 where the fluid velocity rapidly changes value: at the interface between the jet and the main
 455 flow, and the no-slip boundary (Figs. 2,3). Localized vorticity is created at the jet margins and
 456 convected with the flow, where the degree of convection depends on Re_{in} (Fig. 9). Combined with
 457 the integral no-slip condition at the solid fracture boundary⁶⁶, viscous diffusion of vorticity from
 458 the main jet leads to a recirculating vortex on either side of it. The relative strengths of convection
 459 and diffusion of vorticity vary with Re_{in} , and lead to variations of the characteristic flow pattern
 460 (see Sec. V C). In summary, the characteristic jet and recirculating flow pattern is controlled
 461 by viscous shear layers, and we propose that this is a consistent feature of Newtonian flux-driven
 462 fractures. We expect that this flow pattern is unique to flux-driven Newtonian fractures: buoyancy-
 463 driven fractures can achieve greater tip velocities and exhibit unidirectional flow profiles⁴¹, and
 464 non-Newtonian fluids have a shear rate-dependent viscosity that would likely result in markedly
 465 different flow patterns.

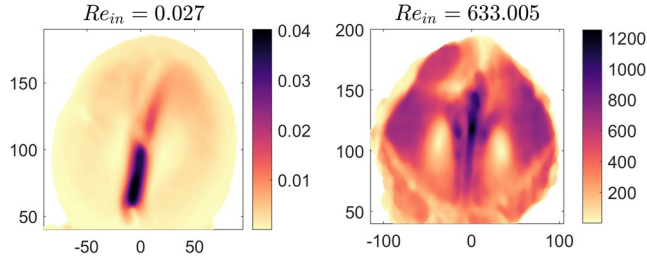


FIG. 10. Filled contours of local flow Reynolds numbers for experiments S3 ($Re_{in} = 0.027$) and W3 ($Re_{in} = 633.005$), at a normalized time of $t^* = 0.5$. Local Re values were obtained at each measurement point throughout the 2D fracture profile, using the local velocity magnitude $|\mathbf{u}|$ and constant values of ρ_f , μ and H .

466 B. Characteristic Reynolds numbers of fracture flow

467 Fluid velocity in a flux-driven fracture is strongly spatially variable, which challenges the mean-
 468 ing of assigning a single Reynolds number to characterize fracture flow. The Reynolds number
 469 varies locally — this is highlighted by the range in alternative Re values for a single experiment
 470 (Fig. 6), and also by profiles of local Reynolds numbers. Fig. 10 exhibits filled contours of local
 471 Re for experiments S3 ($Re_{in} < 1$) and W3 ($Re_{in} > 1$), showing the range of different flow regimes
 472 that can arise within a single fracture at a snapshot in time. In experiment S3, local Re reaches
 473 0.04 in the jet, yet is of the order 10^{-3} throughout the majority of the profile (outside the jet). This
 474 is reflected in the alternative characteristic Re definitions (for S3), where $Re_{jet} = 0.03$ is signifi-
 475 cantly higher than $Re_{mean} = 0.004$ (Tab. III). Conversely, W4 has high local Re values throughout
 476 the fracture profile, and $Re_{jet} (\approx 845)$ is much closer in value to $Re_{mean} (\approx 440)$. Although the Re
 477 distribution is relatively uniform in W4, local Re contours also show spatial variation. Regions of
 478 both $Re > 1000$ (in the jet) and $Re < 1000$ potentially indicate simultaneous turbulent and laminar
 479 regimes.

480 Although Re varies locally (Fig. 10), it remains useful to characterize internal fracture flow
 481 with a single Reynolds number estimate. $Re_{mean} = \frac{\rho_f w u_{mean}}{\mu}$ characterizes the overall flow well
 482 for both $Re_{in} < 1$ and $Re_{in} > 1$. When $Re_{in} < 1$, the high-velocity (and high Re) jet region is
 483 concentrated to a relatively small area. Momentum flux near the fluid inlet is rapidly dissipated,
 484 and Re_{mean} is significantly smaller than Re_{in} and Re_{jet} . Re_{mean} is in fact approximately equal to
 485 Re_{tip} and Re_0 (Fig. 6 and Tab. III). When $Re_{in} > 1$, momentum flux is distributed throughout
 486 the fracture, and Re_{mean} is of a similar value to Re_{in} and Re_{jet} . In practical applications, u_{tip}

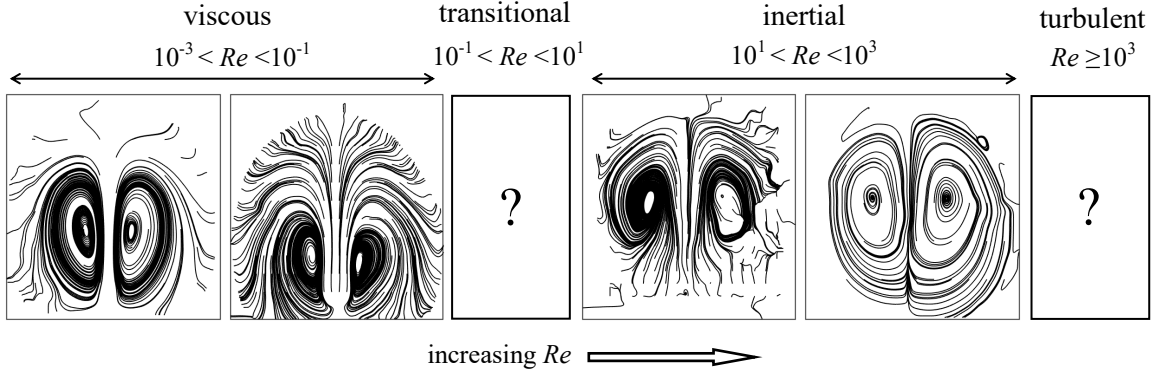


FIG. 11. Flow streamlines showing the different patterns that occur for different flow regimes (at dimensionless time $t^* = 0.5$), based on the inlet Reynolds number. From left to right, these streamlines represent experiments S1, S5, W3 and W5. Each fracture shape has been normalized by its the maximum length and breadth. Streamlines were generated using the Matlab⁵⁹ streamlines function.

487 and u_{in} can be measured^{15,72,73} whereas internal fluid velocities u_{mean} and u_{jet} cannot. Therefore,
 488 we propose that Re_{tip} or Re_0 provides an appropriate characteristic Reynolds number for slow,
 489 viscosity-dominated fractures whilst Re_{in} is more appropriate for fractures with important inertial
 490 effects. Note that calculating Re_{in} also requires knowledge of the area of the injection source –
 491 this is straightforward in analogue experiments, but not necessarily in nature.

492 C. Flow regimes in flux-driven fractures

493 We propose that flux-driven fracture flow can be split into four regimes (Fig. 11) according
 494 to the inlet Reynolds number: viscous ($Re_{in} < 10^{-1}$), transitional ($10^{-1} \leq Re_{in} \leq 10^1$), inertial
 495 ($10^1 < Re_{in} < 10^3$) and turbulent ($Re_{in} \geq 1000$). These regimes have been identified based on
 496 internal flow patterns and the behaviour of average fluid forces across the Re_{in} range. Whilst
 497 fracture flow is characterized by a localized jet and recirculation, our experiments show that this
 498 pattern can vary significantly within the range $0.009 \leq Re_{in} \leq 633$. However, our experiments do
 499 not cover the range $0.1 \leq Re_{in} \leq 37$ — which is where we suggest a transitional regime between
 500 viscous and inertia dominated flow exists. The dynamics of flux-driven fracture flow in transitional
 501 and turbulent regimes have proved challenging to explore experimentally and therefore should be
 502 the subject of future research.

503 **1. Viscous regime**

504 In the viscous regime ($Re_{in} < 0.1$), the jet always terminates before reaching the vertical fracture
505 tip. Increasing Re_{in} leads to shorter jets with higher velocities near the inlet (Figs. 2,3A), and
506 a higher magnitude of vorticity over a smaller region (Fig. 9). Viscous diffusion of vorticity
507 also reduces, and the recirculatory zones become smaller yet more intense (i.e. contain higher
508 velocities and vorticity) with increasing Re_{in} (Fig. 11). A decrease in jet height with increasing
509 Re is the opposite of what occurs for unconfined jets in a fluid tank with no upper boundary⁶⁵, for
510 $Re > 1$. Those jets decrease in height with decreasing Re , as more momentum is dissipated by
511 viscous forces.

512 Viscous forces are greater than inertial forces and scale linearly with Re_{in} initially, whilst in-
513 inertial forces increase at a faster rate and scale with Re_{in}^2 (Fig. 7). This is expected from simple
514 order of magnitude estimates of the fluid force terms (see Equations (11) and (12)), where $|\bar{\mathbf{F}}_V|$
515 scale with characteristic velocity \mathcal{U} and $|\bar{\mathbf{F}}_I|$ scales with \mathcal{U}^2 (note that the Reynolds number rep-
516 resents a velocity scale). Higher inertial forces lead to higher velocities and vorticity in the jet,
517 yet the simultaneous increase in viscous forces inhibits the jet from increasing in length. As Re_{in}
518 approaches unity, the increase in both $|\bar{\mathbf{F}}_I|$ and $|\bar{\mathbf{F}}_V|$ slows down, indicating that the flow is near
519 the onset of the transitional regime.

520 **2. Inertial regime**

521 For $10^1 < Re_{in} < 10^3$, the jets do not terminate prior to reaching the upper solid boundary
522 (Fig. 3A). Vorticity is convected with the jet flow and along the fracture margins (Fig. 9). For the
523 highest $Re_{in} = 633$ (experiments W3 and W4), convection of vorticity towards the solid boundary
524 dominates over viscous diffusion away from it, and a layer of high vorticity is confined to the entire
525 fracture boundary. This leads to flow circulation throughout the entire fracture. With decreasing
526 Re_{in} , vorticity is convected some distance along the upper fracture boundary before the flow loses
527 momentum and vorticity is diffused from the boundary into the main flow. This results in vortices
528 that are located closer to the upper fracture tip (as opposed to $Re_{in} < 1$ where the vortices are
529 located near the injector Fig. 11). Based on numerical simulations of an air-inflated balloon, it is
530 expected that at higher Re , each vortex will split into multiple smaller scale vortices⁶⁸. However,
531 this pattern would be altered significantly by turbulent flow.

532 Scaling arguments (see Equations (11) and (12)) suggest that $|\bar{\mathbf{F}}_I|$ should increase with the
533 square of the velocity (and therefore the Reynolds number). Whilst this is observed within the
534 viscous regime, when $Re_{in} > 1$ inertial forces increase linearly with Re_{in} (Fig 7). Unlike in the
535 viscous regime where velocities vary smoothly over larger length scales (Fig. 3), in the inertial
536 regime, flow is strongly spatially variable with finer-scale flow structures. Therefore, an average
537 approximation of the inertial force may not be fully representative. Our results suggest that as
538 Re_{in} decrease and approaches one, $|\bar{\mathbf{F}}_I|$ and $|\bar{\mathbf{F}}_V|$ deviate from their scaling laws. This potentially
539 indicates that the flow is near the transitional regime, at the higher end of the Re_{in} range.

540 **3. Transitional regime**

541 Our experiments do not span the transitional range $0.1 \leq Re_{in} \leq 36$, which is challenging to
542 achieve experimentally. When injecting silicone oil ($Re_{in} < 0.1$), we reached the maximum pos-
543 sible Re_{in} ($= 0.092$) that could be achieved experimentally. For the water experiments, it was not
544 possible to inject fluid at a lower rate than presented here, without potential settling of the tracer
545 particles. Therefore it is currently unclear how internal fracture flow behaves in the transition
546 from viscous to inertial flow. However, the flow behaviour in the viscous and inertial regimes
547 gives some insight into what occurs during the transition.

548 Fracture tip velocities are very similar for experiments that lie at the transitional margins, which
549 suggests that the transitional regime has a narrow range in tip velocities (Fig. 12). This region of
550 approximately constant tip velocity coincides with a shift in the behaviour of the fluid forces
551 (Fig.7). Unlike in the viscous and inertial regimes $|\bar{\mathbf{F}}_V|$ and $|\bar{\mathbf{F}}_I|$ do not appear to be a simple
552 function of Re_{in} . During the transition from viscous to inertial flow, $|\bar{\mathbf{F}}_V|$ must decrease from the
553 high values in the $Re_{in} < 1$ experiments to the lowest values in the $Re_{in} > 1$ experiments. Similarly,
554 $|\bar{\mathbf{F}}_I|$ must increase across the transition, although Fig.7 indicates that this increase is non-linear
555 and $|\bar{\mathbf{F}}_I|$ potentially plateaus before increasing again. However, there are not enough data points
556 to determine how the fluid forces evolve across the transitional regime. When $Re \approx 1$, inertial
557 and viscous forces become similar in magnitude and are of equal importance. More experiments
558 are needed to understand how fluid forces evolve across the transitional regime, and determine the
559 exact Re values at which the transition occurs. Future experiments could use tracer particles with a
560 lower density, or a Newtonian fluid with a higher density and viscosity than water (yet less viscous
561 than silicone oil).

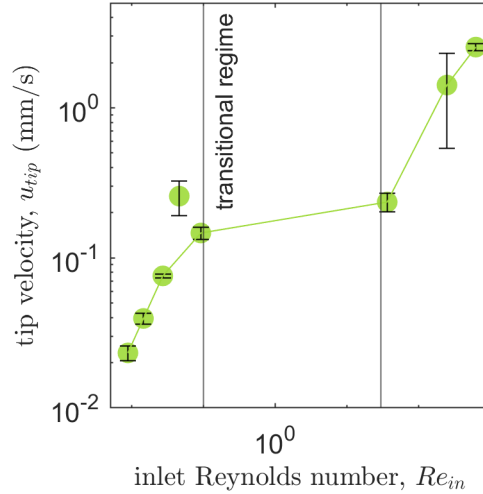


FIG. 12. Vertical fracture tip velocity u_{tip} (mm/s) against Re_{in} at a dimensionless time of $t^* = 0.5$. The line connects all experiments except S4, which is an outlier due to a trapped bubble and additional buoyancy effects.

562 4. Turbulent regime

563 The onset of turbulence in fractures is commonly assumed to occur at $Re \approx 1000^{1,3,74}$. Ex-
 564 periments W3 and W4 have regions of flow with both local $Re > 1000$ and $Re < 1000$ (Fig.10),
 565 potentially indicating simultaneous laminar and turbulent regimes. However, the uniform struc-
 566 ture of the jet (Fig. 4 in the Supplementary Material) and consistent, non-chaotic flow behavior
 567 suggest that fracture flow is not turbulent in our experiments⁷⁵. We injected fluid at the fastest rate
 568 achievable with our injection equipment, yet could not achieve turbulence when injecting water.
 569 Note that, it could be possible to achieve turbulence by injecting liquids with a lower viscosity than
 570 water, yet these liquids also needs to satisfy the condition of having a similar density to gelatine
 571 and being able to hold tracer particles in suspension. An appropriate liquid may be challenging to
 572 identify, and this could be explored in future work. Thus, the dynamics of turbulent fracture flow
 573 remain an open question. Local Re contours (Fig.10) suggest that the central jet would be the first
 574 region of flow to become turbulent, but it is currently unclear whether the characteristic circulating
 575 flow structure would persist at higher Reynolds numbers. Note that 2D flow profiles do not reveal
 576 how the fluid is behaving in the third out-of-plane dimension — 3D imaging is required to know
 577 if the flow is chaotic across the fracture thickness.

578 **D. Controls on fracture propagation**

579 Pressure scale estimates (Fig 8) suggest that the initial deceleration in u_{tip} is due to an increase
580 in either the viscous or inertial resistive pressure as the fracture grows, depending on Re_{in} . The
581 resistance to fracture (characterized by P_f) decreases with fracture length, causing the fracture to
582 accelerate towards the free surface⁷⁶. The horizontal fracture growth, u_W consistently decelerates,
583 which doesn't coincide with any pressure scale. However, this does coincide with a decrease in
584 the circulation velocity over time (Fig. 5H), suggesting that a reduction in the downwards flow
585 velocity leads to a reduction in u_W .

586 The tip velocity u_{tip} is a linear function of the flux Q (Fig. 4), yet for a given tip velocity, there
587 is a wide range of potential fluid behavior within. Across the transitional regime, injection rates
588 and tip velocities are very similar, yet Re_{in} ranges from approximately 0.1 to 30 (Fig. 12). This
589 suggests that a constant proportion of the driving pressure (due to fluid injection) contributes to-
590 wards fracture propagation, regardless of the internal fluid behavior. The remainder of this applied
591 pressure is distributed via different combinations of the inertial and fluid forces (Fig. 8), producing
592 different internal flow patterns (Figs.2 and 11). Although the fluid injection rate controls fracture
593 propagation, we expect that the internal fluid dynamics have a subtle but potentially significant
594 effect on the coupled solid host deformation. Further experiments focusing on solid displacement
595 measurements are needed to investigate this.

596 Through a theoretical analysis of a 2D flux-driven fracture propagating in an infinite, elastic
597 medium, Emerman et al. (1986)¹⁰ found the tip velocity to be a linear function of the inlet velocity,
598 with $u_{tip} \approx 0.45u_{in}$. This is a markedly different relationship to our experiments, where u_{tip} is
599 consistently less than 1% of u_{in} (Fig. 2 in the Supplementary Material). This difference is likely
600 due to the 2D plane-strain model assumption with an infinite fracture width, as opposed to the point
601 source injection method in our experiments. In propagating, non-buoyant fractures, the fracture
602 width is always expected to exceed the length scale of the fluid inlet due to radial fracture growth.
603 However, the size of the injection area affects the inlet Reynolds number and will have a strong
604 effect on flow dynamics. This was investigated in numerical simulations of the flow within an
605 inflating balloon⁶⁸: smaller inlets led to longer and more focused jets, yet jet formation and flow
606 circulation always occurred. Future experimental work could investigate the effect of the fluid
607 inlet area on flux-driven fracture dynamics.

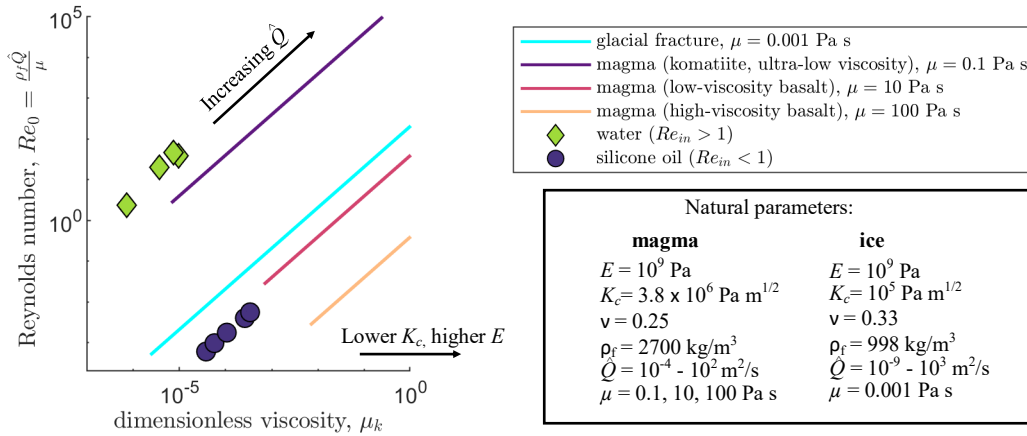


FIG. 13. Dimensionless parameter space for experiments and natural geophysical examples, defined by the dimensionless viscosity μ_k and the Reynolds number $Re_0 = \frac{\hat{Q}\rho_f}{\mu}$. The water and silicone oil experiments are depicted by green diamonds and purple circles respectively. Each line represents the potential range of μ_k and Re_0 for a range of $\hat{Q} = Q/W$ values, with all other parameters constant. The light blue line depicts ice fractures, whilst the pink and purple lines represent dykes with for three different magma viscosities.

608 E. Application to magmatic and glacial systems

609 Experimental, flux-driven fractures in gelatine are an idealized analogue of natural, geophysical
 610 flux-driven fractures. The dimensionless viscosity μ_k and inertia R_k are $\ll 1$ (Tab. IV), confirming
 611 that the analogue fractures propagate in the toughness regime. Although fracture toughness dom-
 612 inates overall, the ratio of viscous and inertial forces varies significantly across the experiments.
 613 We now consider the dynamic similarity between the analogue experiments and natural glacial
 614 and magmatic systems by comparing the dimensionless parameter space defined by μ_k and the
 615 Reynolds number. For the latter, we use the definition $Re_0 = \frac{\rho_f \hat{Q}}{\mu}$, where $\hat{Q} = \frac{Q}{W}$ is the flux per unit
 616 width, in order to directly compare with nature. This requires appropriate estimates of magmatic
 617 and glacial parameters.

618 The applicability of the experiments to natural systems is limited by the model assumptions
 619 (e.g. elastic solid, Newtonian fluid). Glacial ice is not strictly linear elastic, but it is accepted to
 620 behave in an elastic way under fracture⁷⁷. In magmatic systems the assumption of elasticity is
 621 only applicable to the lithosphere⁷⁸. Appropriate E values range from $10^8 - 10^{10}$ Pa for glaciers⁷⁹
 622 and $10^9 - 10^{10}$ Pa for the elastic crust⁸⁰. Water in glacial crevasses is Newtonian ($\mu \approx 10^{-3}$
 623 Pa s), whereas the rheology of magma depends on the relative proportions of crystals, melt and

bubbles⁸¹. Newtonian magmas are relatively crystal-poor with no bubbles, representing a primitive
mafic (low-silica) magma: μ ranges from $10^1 - 10^2$ Pa s for basaltic magma^{82,83}, yet can be as low
as 10^{-2} for ultramafic magmas such as komatiite or carbonatite²⁹. Numerical models of basaltic
dykes with a constant flux suggest that Q can range from $1 - 1000$ m³/s⁸⁴, whereas glacial fractures
exhibit a wider range of Q values — from $O(10^{-5})$ m³/s in thin fracture networks⁷³, to $O(10^3)$
m³/s in rapid drainage events¹⁵. Fracture lengths, widths and thicknesses have a wide variety of
potential values, and we consider a range of fracture sizes. The natural parameter estimates are
summarized in Tab. IV.

According to the dimensionless parameter space defined by μ_k and Re_0 (Fig. 13), our experi-
ments represent the lowest end of the μ_k spectrum for natural magmatic and glacial fractures. The
natural parameter space is depicted as a series of linear lines, each representing a different fluid
viscosity, and a wide range of \hat{Q} values. All other parameters are assumed to be constant: whilst ρ_f
and v have little effect on the overall parameter space, E and K_C do act to shift the dimensionless
viscosity range significantly. Here we have selected values that best represent the experiments in
this parameter space, whilst being in the valid range specified above (the upper and lower ends
of the E and K_C ranges respectively, see Fig. 13 inset box). Both sets of experiments represent
natural injections with a low flux (per unit width). Recall that the rate of injection is limited by
the fluid viscosity (silicone oil could not be injected at a higher rate than achieved here). The
silicone oil experiments ($Re_{in} < 1$) are fairly well representative of glacial fractures and basaltic
dykes, whereas the water experiments ($Re_{in} > 1$) are more representative of a very low viscosity
magma, such as a primitive komatiite. Future work could explore fractures with higher μ_k , and fill
in the gaps in our understanding of flux-driven fluid dynamics across the natural parameter space.
A higher μ_k could be achieved by using other fluids with different viscosities, and injecting them
at a range of rates.

VI. CONCLUSIONS

Analogue experiments of flux-driven fractures have shown that internal fracture flow has a
self-similar pattern of a high-velocity central jet with a zone of fluid recirculation on either side,
consistent across a range of regimes. We have utilized PIV velocity data to identify four po-
tential regimes: viscous, inertial, transitional, and turbulent. Viscous and inertial regimes were
produced experimentally (with some experiments perhaps bordering the transitional regime) for

	Experiments	Magma (basalt)	Ice
E	2278 - 4337	$1 \times 10^9 - 10^{10a}$	$10^8 - 10^{10i}$
K_C	66-93	$1.4 \times 10^6 - 3.8 \times 10^{6b}$	$10^5 - 20^{5jk}$
ν	0.5	0.1-0.5 ^a	0.3 ⁱ
Q	$4.4 \times 10^{-8} - 7.0 \times 10^{-6}$	1- 10^{3c}	$10^{-5} - 10^3$ ^{lm}
μ	0.001 - 0.45	10^1 - 10^{3de}	10^{-3g}
ρ_f	998,1040	2700 ^p	998 ^g
H	0.003 - 0.015	0.1 - 10^{df}	$10^{-3} - 1$ ⁿ
W	0.14 - 0.18 ^o	$10^1 - 10^{4h}$	$1 - 10^{4h}$
Re_0	$6 \times 10^{-4} - 5 \times 10^1$	$10^{-4} - 10^4$	$10^{-4} - 10^8$
μ_k	$5 \times 10^{-7} - 3 \times 10^{-4}$	$10^{-5} - 10^{10}$	$10^{-8} - 10^{10}$
R_k	$4 \times 10^{-10} - 9 \times 10^{-6}$	-	-

TABLE IV. Characteristic parameters of flux-driven fractures in the current experiments and magmatic and glacial settings: Young’s modulus E (Pa), fracture toughness K_C ($\text{Pa m}^{\frac{1}{2}}$), Poisson’s ratio ν (dimensionless), volumetric flux Q (m^3/s), viscosity μ (Pa s), fluid density ρ_f (kg/m^3), fracture thickness H (m), fracture width W (m), Reynolds number $Re_0 = \frac{Q\rho_f}{B\mu}$, dimensionless viscosity $\mu_k = \frac{12\mu Q}{E} \left(\frac{E}{K}\right)^4$, dimensionless inertia $R_k = \frac{\rho_f E^{5/3} Q^{5/3}}{K^{8/3} t^{1/3}}$. R_k was calculated at $t^* = 0.5$ for the experiments, and not estimated in nature. References for natural values: a) Heap et al. (2020)⁸⁰, b) Balme et al. (2004)⁸⁵, c) Traversa et al. (2010)⁸⁴, d) Wada et al. (1994)⁸², e) Roman et al. (2021)⁸³, f) Rubin (1995)³, g) values for water, h) assuming a range of sizes, i) Vaughan (1995)⁷⁹, j) Fischer et al. (1995)⁸⁶, k) Rist et al. (1999)⁸⁷, l) Das et al. (2008)¹⁵, m) Fountain et al. (2005)⁷³, n) Holdsworth et al. (1969)⁸⁸, o) range in experimental values at $t^* = 0.5$, p) typical basalt value⁸⁹.

654 inlet Reynolds numbers spanning $O(10^{-3}) \leq Re_{in} \leq O(10^3)$. In the viscous regime, the jet and
655 adjacent vortices shrink with increasing Re yet become more intensely localized near the jet. To
656 our knowledge, this is the first experimental insight into the behavior of jets at $Re < 1$. In the
657 inertial regime, the jet length always exceeds the fracture length, and an increase in Re_{in} leads to a
658 greater degree of flow circulation. Although data are lacking for the transitional regime ($Re \approx 1$)
659 due to experimental limitations, we propose that the average fluid forces have a complex relation-
660 ship with Re_{in} , yet fractures propagate at similar tip velocities within this regime. Despite the

661 complexity of the internal flow, the propagation velocity is a linear function of the flux Q . These
662 results have important implications for interpreting natural data on propagating fractures, and de-
663 veloping better numerical models to predict them. A key advantage of our experimental model is
664 that the solid transparency allows for measurements of fracture and flow dynamics in real-time.
665 Furthermore, the model scales appropriately with natural flux-driven fractures, as shown by the
666 dimensionless parameter space defined by Re and the dimensionless viscosity μ_k . However, there
667 remains a knowledge gap regarding transitional and turbulent flow in fractures. Model simplifi-
668 cations also restrict our analysis to fractures in elastic solids injected by Newtonian fluids with a
669 constant viscosity and density. These assumptions are most restrictive in the application to vol-
670 canology, where hot rocks can deform inelastically, and crystal and bubble content can lead to
671 variations in magma viscosity and density. Future experiments (using different fluid and solid
672 properties) are required to understand the complete range of flow regimes in flux-driven fractures,
673 across the full natural parameter space. Experimental measurements in 3D would bring further
674 advancement to our understanding of fracture dynamics.

675 **VII. SUPPLEMENTARY MATERIAL**

676 Supplementary Material to this article is provided online, containing further details on the ex-
677 perimental methodology and additional results visualizations.

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691 **DATA AVAILABILITY STATEMENT**

692 The data that support the findings of this study are openly available in the NERC EDS National
693 Geoscience Data Centre at <https://doi.org/10.5285/92383dba-8b6e-43da-9d78-e8784b97124c>.

694 **Appendix A: Nomenclature**

Experimental parameters	
A	fluid inlet area
c_g	gelatine concentration
d	fluid inlet diameter
E	Young's modulus
K_C	fracture toughness
μ	dynamic viscosity
ν	Poisson's ratio
ρ_f	fluid density
Q	volumetric flux
\hat{Q}	flux per unit width
t_0	time when $L = 10$ cm
t_{erupt}	time between t_0 and eruption
u_{in}	inlet velocity

Geometry	
H	fracture thickness
L	fracture length
W	fracture width
x	horizontal axis
y	out of plane axis
z	vertical axis
\hat{x}	normalized horizontal coordinate
\hat{z}	normalized vertical coordinate
\hat{z}_{max}	normalized height of velocity maximum

Mathematical formulation	
$\partial/\partial t$	partial time derivative
∇^2	Laplace operator
\mathbf{F}_I	inertial force, $\mathbf{F}_I = (F_{Ix}, F_{Iz})$
\mathbf{F}_V	viscous force, $\mathbf{F}_V = (F_{Vx}, F_{Vz})$
\mathbf{n}	unit normal
p	dynamic pressure
P_F	fracture pressure scale
ΔP_V	viscous pressure scale
ΔP_I	inertial pressure scale
t	time, $t_0 \leq t \leq t_{max}$
\mathbf{u}	velocity vector, $\mathbf{u} = (u_x, u_z)$
$\boldsymbol{\omega}$	vorticity, $\boldsymbol{\omega} = \nabla \times \mathbf{u}$
ω	vorticity magnitude

Characteristic velocities	
u_{in}	inlet velocity
u_{tip}	vertical tip velocity
u_{tip50}	vertical tip velocity when $t^* = 0.5$
u_W	horizontal tip velocity
u_{W50}	horizontal tip velocity when $t^* = 0.5$
u_{mean}	mean absolute velocity
u_{down}	representative downwards (downflow) velocity
u_{circ}	circulation velocity
u_{jet}	jet velocity

PIV analysis	
r_c	correlation value
Δt	time increment
τ_r	particle relaxation time
U_g	Stokes particle velocity
Δx	grid spacing

Dimensionless numbers	
μ_k	dimensionless viscosity
Re_0	flux Reynolds number
Re_{in}	inlet Reynolds number
Re_{mean}	mean Reynolds number
Re_{jet}	jet Reynolds number
Re_{tip}	tip Reynolds number
R_k	dimensionless inertia
St	Stokes number
t^*	dimensionless time, $t^* = (t - t_0)/t_{erupt}$

Numerical (finite difference) approximations	
\hat{F}_I	inertial force
\hat{F}_V	viscous force
\bar{F}_I	average inertial force
\bar{F}_V	average viscous force
$\Delta\hat{P}_V$	viscous pressure scale
$\Delta\hat{P}_I$	inertial pressure scale

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